



# Airborne high spectral resolution lidar observation of pollution aerosol during EUCAARI-LONGREX

S. Groß<sup>1</sup>, M. Esselborn<sup>1,\*</sup>, F. Abicht<sup>1,\*\*</sup>, M. Wirth<sup>1</sup>, A. Fix<sup>1</sup>, and A. Minikin<sup>1</sup>

<sup>1</sup>Deutsches Zentrum für Luft- und Raumfahrt (DLR), Institut für Physik der Atmosphäre (IPA), Münchner Str. 20, Oberpfaffenhofen, 82234 Weßling, Germany

\* now at: European Southern Observatory, Karl-Schwarzschild-Str. 2, 85748 Garching, Germany

\*\* now at: Max-Born-Institut, Max-Born-Str. 2a, 12489 Berlin, Germany

Correspondence to: S. Groß (silke.gross@dlr.de)

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**Abstract.** Airborne high spectral resolution lidar observations over Europe during the EUCAARI-LONGREX field experiment in May 2008 are analysed with respect to the optical properties of continental pollution aerosol. Continental pollution aerosol is characterized by its depolarisation and lidar ratio. Over all, the measurements of the lidar ratio and the particle linear depolarization ratio of pollution aerosols provide a narrow range of values. Therefore, this data set allows for a distinct characterization of the aerosol type “pollution aerosol” and thus is valuable both to distinguish continental pollution aerosol from other aerosol types and to determine mixtures with other types of aerosols.

of aerosols on the Earth’s climate system, long-term observations and coordinated observations during intensive field campaigns (such as closure studies, Quinn et al., 1996) are crucial. Furthermore, the presence of aerosols in the lower atmosphere plays a significant role in air quality and health considerations.

The European integrated project on Aerosol Cloud Climate and Air Quality Interactions (EUCAARI) was supported by the European Commission under the Sixth Framework Programme with the aim of investigating the role of aerosols on climate and air quality (Kulmala et al., 2009). Embedded in the EUCAARI intensive observational period, the aircraft field experiment EUCAARI-LONGREX (Long Range Experiment) was conducted in May/June 2008 to investigate physical and chemical properties of atmospheric aerosols within the free troposphere and the boundary layer on European scale. Essential activities of this experiment included a comprehensive characterization of main aerosol types over Europe and their exchange between boundary layer and free troposphere. Spatial and temporal gradients of aerosol loading were investigated to evaluate aerosol sources and sinks. Moreover, the experiment included a closure study of aerosol optical thickness observed from ground, sky and space (Kulmala et al., 2011). During EUCAARI-LONGREX two research aircraft were based at the airport of Oberpfaffenhofen, Germany. The Falcon 20 jet, operated by the German Aerospace Centre (DLR), equipped with a nadir viewing HSRL (High Spectral Resolution Lidar) and a comprehensive in-situ instrumentation generally focused on the upper troposphere and occasionally guided the low flying British

## 1 Introduction

Aerosol particles are a main component of the atmosphere and affect the Earth’s climate system in two different ways. They scatter and absorb the solar and terrestrial radiation, and they act as cloud condensation nuclei and for this reason influence cloud properties. According to the findings of the IPCC (Foster et al., 2007) the current level of scientific understanding of the impact of aerosols on the global climate system is considered as medium to low. The effect of aerosols depends on the spatial and vertical distribution of aerosols, on their chemical and microphysical properties, as well as on the reflectance of the underlying surface and the presence of clouds. As aerosols show a large spatial and temporal variability their impact on regional scale can be quite different (IPCC, 2007). To improve our knowledge on the impact

BAe-146 (FAAM, Facility for Airborne Atmospheric Measurements) into specific aerosol layers. Above ground observation sites and other regions of particular interest the Falcon flew stacked vertical profiles in order to sample the atmospheric column from high altitude down to the boundary layer.

Airborne lidar instruments are particularly suitable for observing the horizontal and the vertical distribution of aerosol structures on a regional scale. The advantage of HSRL over normal backscatter-only lidar is the ability to directly measure the extinction-to-backscatter ratio (lidar ratio) of aerosols (Shiple et al., 1983; Shimizu et al., 1983). The value of this quantity depends on the physical properties of aerosols, namely the size distribution, the complex index of refraction and the morphology (Evans, 1988). It does not depend on its concentration. Current spaceborne lidars such as the Cloud Aerosol Lidar and Infrared Pathfinder Satellite Observation (CALIPSO) instrument (Winker et al., 2007) provide global coverage but only measure the attenuated backscatter signal. Hence the retrieval of vertical aerosol backscatter and extinction profiles relies on a-priori assumptions on the lidar ratio. Therefore, the accuracy of the retrieved extinction profile is determined by the assumed lidar ratio (e.g. Wandinger et al., 2010). The results of airborne HSRL measurements can be helpful to improve the informative content of spaceborne measurements. This paper focuses on the analysis of the airborne HSRL measurements of the aerosol lidar ratio and the particle linear depolarisation ratio taken during the first half of EUCAARI-LONGREX (6 May to 14 May). Both quantities only depend on the aerosol type and can be used to classify different types of aerosol (Groß et al., 2011, 2012; Burton et al., 2012).

## 2 Methods

### 2.1 DLR Falcon HSRL

During the EUCAARI-LONGREX field campaign a HSRL system was operated on board the DLR Falcon research aircraft. This system was originally developed as an airborne demonstrator for a possible future spaceborne multi-wavelength differential absorption lidar (DIAL) for water vapour measurements at 935 nm (WALES – Wirth et al., 2009). In addition to the H<sub>2</sub>O-channels, the system provides backscatter and depolarisation channels at 1064 nm and 532 nm for aerosol characterization. At 532 nm a molecular channel is implemented for the HSRL measurements. The HSRL receiver channel was already successfully deployed during former field campaigns. A detailed technical description of the HSRL system and a presentation of its measurements can be found by Esselborn et al. (2008) and Esselborn et al. (2009). Typically, the system allows signal integration times as short as a few seconds with an acceptable statistical error of less than 5 % in the backscatter data. The vertical

resolution of the profiles is 15 m for all backscatter and depolarisation data. For the extinction and lidar ratio profiles the vertical resolution is 360–700 m. Relative systematic errors of the backscatter measurements are typically 4 % to 8 %, statistical errors are typically < 1 % and therefore negligible. For the extinction measurements the relative systematic errors are usually less than 5 %, so that statistical errors of typically 5 % to 20 % are dominant. Relative systematic errors of the depolarisation measurements are primarily due to mechanical accuracy of the calibration measurement and result in a relative error of the particle linear depolarisation ratio of the order of 10 % to 16 %. For detailed information on the error calculation see Esselborn et al. (2008) and Freudenthaler et al. (2009). Because of eye safety regulations in the considerably busy airspace over Europe, HSRL measurements at 532 nm were restricted to places which were cleared by air traffic control.

### 2.2 Meteorological conditions during EUCAARI

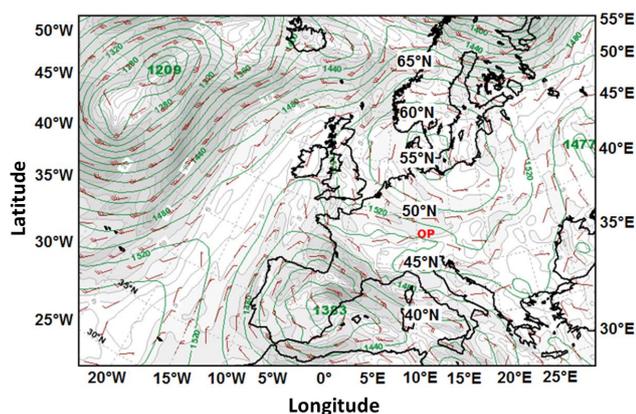
The synoptic situation of the first half of the EUCAARI-LONGREX field campaign, from 6–14 May 2008, was dominated by a persistent high-pressure blocking system, remaining almost stationary over southern Sweden and Denmark (Fig. 1). Due to the stability of the system during the main part of the campaign, a constant east-west transport of air-mass could be observed in Central Europe. The absence of precipitation during this period favoured the accumulation of pollution aerosol in the continental boundary layer. Figure 1 shows the typical meteorological situation during the first half of the campaign exemplarily for the geopotential and wind at 850 hPa on 10 May 2008.

From 11 May to 14 May the high-pressure system moved westward, north-east of the British Islands and started to decay on 14 May. The second half of the field campaign was characterized by advection mostly from northerly directions. A detailed description of the meteorological situation during the EUCAARI-LONGREX campaign can be found by Hamburger et al. (2011).

Figure 2 shows a backward trajectory analysis for the five Falcon flight missions during the first half of the field campaign. The backward trajectories show a circulated flow pattern and advection from eastern directions. The trajectories were calculated with the Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPLIT) model (Draxler and Rolph, 2012) and the NCEP Global Data Assimilation System (GDAS) data, and start at the times and places the HSRL measurements were analysed. The duration is five days backward for all trajectories.

### 2.3 Flight missions during EUCAARI

A total of 48 flight hours were performed during 15 flight missions (Hamburger et al., 2011, 2012). Due to the persistent high pressure system transporting air masses in westerly



**Fig. 1.** Geopotential and wind field at 850 hPa on 10 May 2008. OP marks the airport of Oberpfaffenhofen.

directions, the flights were restricted to areas north of the Alps. Figure 3 shows a compilation of all Falcon flight missions performed during EUCAARI-LONGREX.

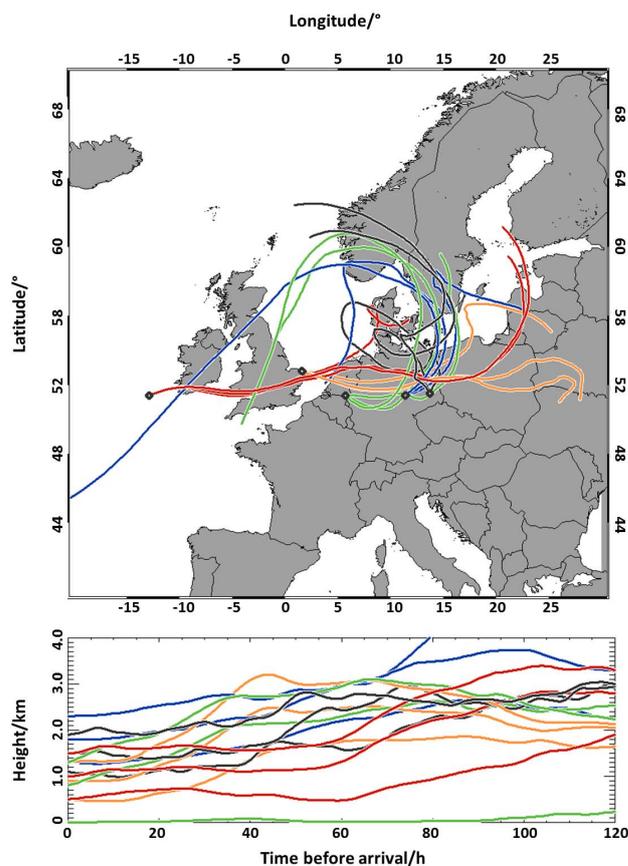
During EUCAARI-LONGREX the German Falcon operated out of Oberpfaffenhofen airport (48.08° N, 11.27° W). The flight patterns of the Falcon were designed for lidar observation at high altitude and subsequent in situ profiling over the instrumented ground stations and other specific regions identified by lidar.

### 3 Results and discussion

Due to the meteorological conditions, only measurements during the first half of the campaign are used for the characterization of pollution aerosols. This restriction is necessary to make sure that pollution aerosol was the dominant, if not the only, aerosol type within the observed aerosol layers, and to ensure high aerosol load within these aerosol layers to achieve high accuracy of the measurements with resulting small uncertainties. We present a case study on 14 May 2008 which shows an exemplary analysis of the measurement data since this day has extensive data analyses in terms of a closure study, i.e. lidar measurements, vertical profiling of in-situ measurements, and satellite (PARASOL) overpass (Kulmala et al., 2011).

#### 3.1 Case study of pollution aerosol on 14 May 2008

The flight mission on 14 May 2008 was dedicated to coordinated observations of a pollution aerosol plume west of Ireland. To ensure maximum flight endurance west of the coast of Ireland, the Falcon aircraft was transferred to Shannon/Ireland the day before and performed a local flight on 14 May 2008. (The Falcon flight track is shown in Fig. 3 in the yellow line). The location west of the Irish coast was chosen due to a forecasted pollution transport event. Backward trajectories (Fig. 2 red line) show that the observed



**Fig. 2.** 5-day backward trajectories for the locations of flight missions during the first half of the EUCAARI-LONGREX field campaign. The trajectories start in height levels identified by the lidar measurements and considered as pollution aerosols (also see Table 1).

air masses were transported from easterly directions, crossing over Northern Germany, the Netherlands, Southern England and Southern Ireland. The mean flight altitude during this flight leg was 10 km. The colour coding in Fig. 4 indicates the column integrated aerosol optical depth (AOD) between flight altitude and ground along the flight track. It is a measure of the column integrated aerosol load and indicates to which degree aerosols prevent the transmission of light through the atmosphere by scattering and absorption. An AOD of less than 0.1 indicates clear sky conditions whereas an AOD of 1 indicates hazy conditions. Along the flight track the AOD ranged between  $\sim 0.1$  near the Irish Coast and  $\sim 0.5$ . The AOD was significantly higher in the south-eastern part of the flight leg.

The lidar backscatter ratio (Fig. 5) indicated the aerosol layers along the flight path, in the marine boundary layer and in the free troposphere up to about 4.5 km altitude. The aerosol structure was inhomogeneous with several elevated layers.

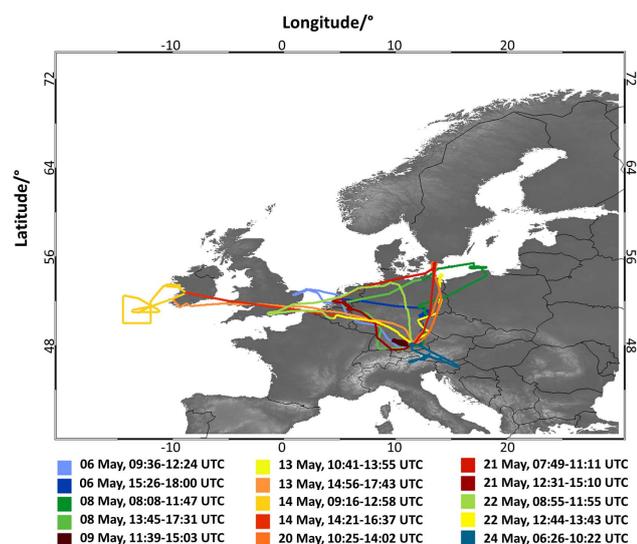


Fig. 3. Falcon flight tracks during EUCAARI-LONGREX.

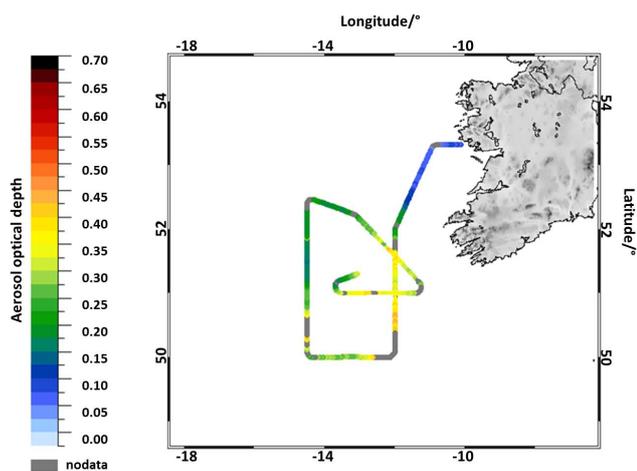


Fig. 4. AOD at 532 nm along the Falcon flight track south-west of Ireland on 14 May 2008. “Nodata” marks areas in which the calculation of AOD was prevented by clouds, or by biased measurements caused by spiralling.

In order to allow the analysis of air masses of pure continental origin having not being influenced by marine boundary layer and/or advection we used a combination of back-trajectories, calculated with the Lagrangian Analysis Tool (LAGRANTO) (Wernli and Davies, 1997; Wernli, 1997), and European Center of Medium-Range Weather Forecast (ECMWF) data. For the definition of the source strength of anthropogenic emission along each trajectory the Emissions Database for Global Atmospheric Research (EDGAR) was used. A detailed description of this procedure as well as an application of trajectory analysis to in-situ observations on board the FALCON aircraft can be found by Hamburger et

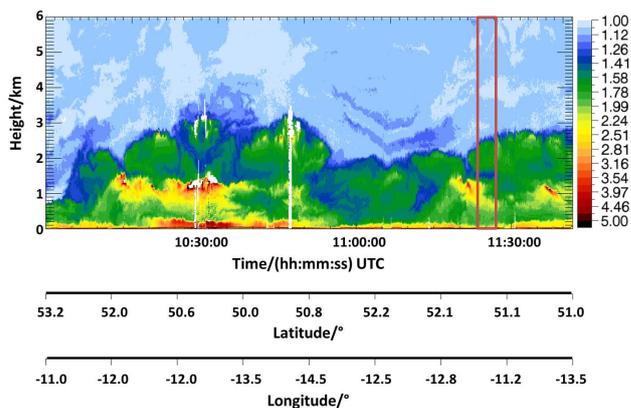


Fig. 5. Aerosol distribution shown as cross-section of the backscatter ratio at 532 nm along the flight track on 14 May 2008 shown in Fig. 4.

al. (2012). The HSRL data was filtered according to the following criteria:

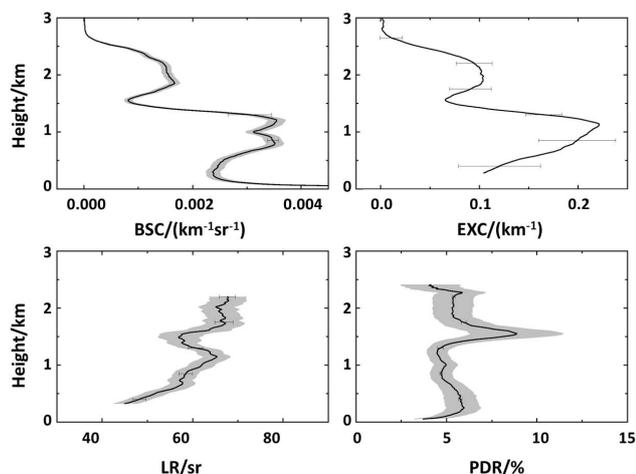
1. The residence time of the aerosol over the European Continent without Scandinavia is larger than 36 h.
2. During these 36 h the trajectories stayed below 2.5 km (750 hPa).
3. During the last 12 h before measurements the trajectories have been above 0.8 km (930 hPa) over ocean.

Furthermore, we choose measurement periods with sufficient aerosol load and vertical homogeneity of the layering to avoid uncertainties due to inhomogeneity in the averaged data. Accordingly, a part in the middle of the flight region was chosen for a detailed study, which is indicated in Fig. 5 by the red rectangle. The mean profiles of the particle backscatter coefficient, the particle extinction coefficient, the particle lidar ratio and the particle linear depolarisation ratio are shown in Fig. 6.

From these profiles, a two layer structure of the selected time period is observable which is also supported by backward trajectories (not shown here). The lower layer, which is mainly influenced by polluted aerosols from source regions over England, reaches up from about 0.9 to about 1.4 km and shows a mean lidar ratio of  $60 \pm 5$  sr and a mean particle linear depolarisation ratio of  $5 \pm 1$  %, the upper layer (from about 1.7 km to 2.5 km) which is mainly influenced from European continental aerosols shows a mean lidar ratio of  $67 \pm 5$  sr and a mean value of the particle linear depolarisation ratio of about  $6 \pm 1$  %. Below about 0.9 km the lidar ratio strongly decreases. This decrease may be caused by marine influenced air masses. The mean particle extinction coefficient in the lower layer is about  $0.2 \text{ km}^{-1}$ , whereas the maximum value of the particle extinction coefficient in the upper layer is only  $0.1 \text{ km}^{-1}$ . The latter results are in very good agreement with the vertical profiles of the extinction

**Table 1.** Mean lidar ratio (LR) and particle linear depolarisation ratio (PDR) at 532 nm for the analysed pollution layers shown in Figs. 6 and 7.  $\pm$ -values indicate the mean systematic uncertainties.

Date	Location	Time (UTC)	Altitude Range [km]	LR [sr]	PDR [%]
6 May	Southeast Coast England	10:56–11:06	0.9–1.7	$56 \pm 6$	$5 \pm 1$
6 May	Melpitz, Germany	16:08–16:15	1.8–2.3	$57 \pm 5$	$6 \pm 1$
8 May	Cabauw, Netherlands	14:45–14:56	0.5–1.3	$52 \pm 4$	$9 \pm 1$
13 May	Cottbus, Germany	11:18–11:26	1.3–1.6	$53 \pm 7$	$6 \pm 1$
14 May	Irish Coast	11:23–11:26	0.9–1.4	$60 \pm 5$	$5 \pm 1$

**Fig. 6.** HSRL vertical profiles of backscatter coefficient (BSC) (upper left panel), extinction coefficient (EXC) (upper right panel), lidar ratio (LR) (lower left panel) and particle linear depolarisation ratio (PDR) (lower right panel) at 532 nm for the selected flight track south-west of Ireland on 14 May. The grey shaded areas mark the systematic error, the error bars the statistical errors.

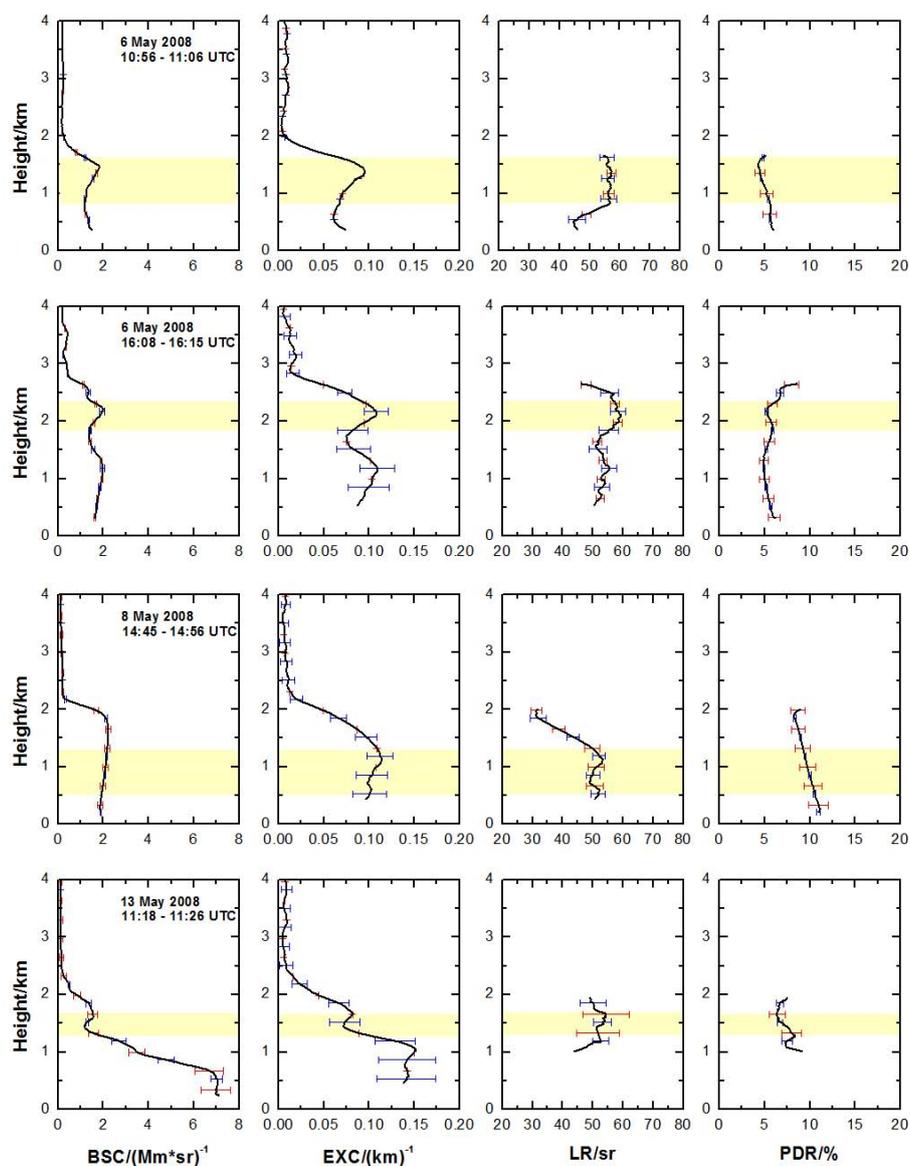
coefficient derived from the in situ particle size distribution measurements on board the Falcon aircraft (Kulmala et al., 2011).

### 3.2 General findings of all pollution cases

#### 3.2.1 Vertical layering

In the previous subsection we presented an analysis of the case study of 14 May 2008 as an example. Now we consider all analysed lidar measurements of pollution cases during the first half of the EUCAARI-LONGREX campaign. To get an overview of the temporal and spatial distribution of the pollution aerosols Fig. 7 shows vertical profiles of the backscatter coefficient, the extinction coefficient, the lidar ratio and the particle linear depolarisation ratio (all at 532 nm) for further four HSRL measurement. The profiles correspond to measurements at the southeast coast of England on 6 May (10:56–11:06 UTC), near Melpitz, Germany, on 6 May (16:08–16:15 UTC), near Cabauw, Netherlands, on

8 May (14:45–14:56 UTC), and near Cottbus, Germany, on 13 May (11:18–11:26 UTC). The backscatter coefficient for these measurements ranged between about  $0.001 \text{ km}^{-1} \text{ sr}^{-1}$  to about  $0.002 \text{ km}^{-1} \text{ sr}^{-1}$  within the pollution layers. The corresponding extinction coefficient in the pollution layer was about  $0.06\text{--}0.1 \text{ km}^{-1}$ . The mean values of the lidar ratio (52–60 sr) and the particle linear depolarization ratio (5–9 %) within the presumed pollution aerosol layers are summarized in Table 1. Considering the lidar ratio and its uncertainties within the corresponding height range we cannot see significant differences for the analysed case studies. However the linear particle depolarization ratio within the considered height range on 8 May (Fig. 7) is higher than the mean values found for the other days. Measurements of the relative humidity from the nearest soundings in space and time (not shown here) show that the relative humidity in the pollution layer on 8 May (14:45–14:56 UTC) was significantly lower (30–50 %) than on the other days (60–90 %). The influence of the relative humidity on the aerosol optical properties was e.g. discussed by Zieger et al. (2011). Theoretical model studies (Ackermann, 1998) revealed that low relative humidity leads to lower values of the lidar ratio of pollution aerosols. This result is in good agreement with the findings of the case study of 8 May (14:45–14:56 UTC), where one of the lowest mean values of the lidar ratio was found. However, whether the low relative humidity can explain the difference in the particle linear depolarization ratio or if it is caused by mixing with other aerosol types remains speculation. Simulations or measurements dealing with the dependence of the particle linear depolarization ratio of pollution aerosols on the relative humidity are, to our knowledge, unavailable so far. In the profiles of the intensive optical properties we furthermore see differences to the height ranges above and below the presumed aerosol layers. In the measurements performed on 6 May (10:56–11:06 UTC) north-east of the English coast and on 14 May (11:23–11:26 UTC) south-west of the Irish coast the lidar ratio beneath the pollution layer (below about 1 km) decreases with decreasing height. This is most probably an indication for the mixture with marine aerosols which have similar values for the particle linear depolarisation ratio (about 3 % – Groß et al., 2011), but significantly lower values (of about 18 sr – Groß et al., 2011)



**Fig. 7.** Vertical profiles of backscatter coefficient (BSC), extinction coefficient (EXC), lidar ratio and particle linear depolarisation ratio (PDR) at 532 nm for selected measurement periods during the first half of the EUCAARI-LONGREX campaign. The red error bars indicate the systematic errors; the blue error bars indicate the statistical errors. Height levels meeting the criteria for pollution influenced layers are indicated by the light yellow.

for the lidar ratio. On 8 May (14:45–14:56 UTC) the lidar ratio above the pollution layer significantly decreases. This decrease may be caused by mixing with less absorbing background aerosols (Tesche et al., 2007). A similar signature could also be observed by Hänel et al. (2012) above pollution aerosol layers over and near Beijing.

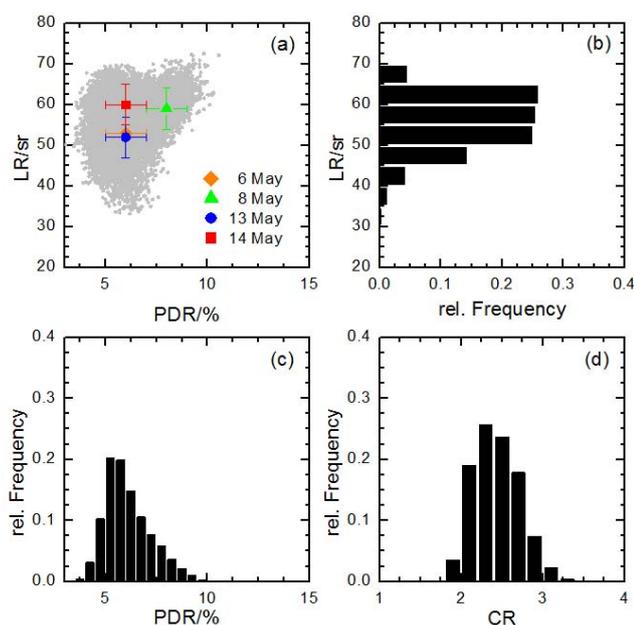
### 3.2.2 Optical properties

Considering all HSRL measurements during the first half of the EUCAARI-LONGREX campaign which achieve the criteria specified in Sect. 3.1 we found a mean lidar ratio of

about 56 sr and a mean particle linear depolarization ratio of about 6 % for anthropogenic pollution aerosols. The lidar ratio in the presumed pollution aerosol layers ranged between 33 sr and 72 sr, and the particle linear depolarisation ratio showed values between 3 % and 11 %. The mean values of both quantities and their statistical information for the different measurement days are summarised in Table 2 and illustrated in Fig. 8. Figure 8a shows the lidar ratio plotted against the particle linear depolarisation ratio for selected flight tracks according to the trajectory criteria described in Sect. 3.1. The grey dots show the results of all measurement

**Table 2.** Lidar ratio (LR) and particle linear depolarisation ratio (PDR) at 532 nm, and their statistical information determined from analysed Falcon flights.  $Q_{0.25}$  and  $Q_{0.75}$  indicate the 1st and 3rd Quartile.

Date	LR				PDR			
	mean $\pm$ std [sr]	Range [sr]	$Q_{0.25}$ [sr]	$Q_{0.75}$ [sr]	mean $\pm$ std [%]	Range [%]	$Q_{0.25}$ [%]	$Q_{0.75}$ [%]
6 May	53 $\pm$ 6	33–69	48	56	6 $\pm$ 1	3–8	5	6
8 May	59 $\pm$ 5	44–72	56	62	8 $\pm$ 1	6–11	7	8
13 May	52 $\pm$ 3	42–55	50	54	6 $\pm$ 1	5–10	6	7
14 May	60 $\pm$ 5	50–70	59	63	6 $\pm$ 1	5–9	6	7
Over all	56 $\pm$ 6	33–72	51	61	6 $\pm$ 1	3–11	5	7

**Fig. 8.** (a) Scatterplot of lidar ratio (LR) vs. particle linear depolarisation ratio (PDR) (both at 532 nm) for measurements within pollution aerosol layers during the first half of the EUCAARI-LONGREX campaign. The grey dots show all lidar measurement points within pollution layers, the colour-coded squares show the mean values within the pollution layers for the different measurement days summarized in Table 2. (b) and (c) show the frequency distribution of the lidar ratio (LR) and the particle linear depolarisation ratio (PDR) at 532 nm, respectively, for all considered measurements during the first half of the EUCAARI-LONGREX campaign. (d) Frequency distribution of the colour ratio (CR) for all considered measurements during the first half of the EUCAARI-LONGREX campaign.

points considered as pollution aerosols. The colour coded measurement points show the mean values for the individual measurement days.

Figure 8b and c show the corresponding histograms of the observed lidar ratio and the particle linear depolarisation ratio, respectively. Obviously, the most frequent value pairs are within a range of the particle linear depolarization ratio of

5–7 % and a range of the lidar ratio of 50–60 sr. 50 % of all measurement points show values of the particle linear depolarization ratio between 5 % and 7 % and a lidar ratio between 51 and 61 sr, and about 90 % of all measurement points show values between 4 % and 8 % of the particle linear depolarization ratio and between 48 sr and 65 sr of the lidar ratio. The median of the particle linear depolarization ratio is 6 %, of the lidar ratio 56 sr. Lidar ratios around and below 40 sr, found for about 5 % of all measurement points, may either indicate rural particles (Evans, 1988; Anderson et al., 2000) or mixtures with other aerosol types (e.g. clean continental aerosol or marine aerosol).

But not only the lidar ratio and the particle linear depolarisation ratio are only dependent on the aerosol type. Also the colour ratio (CR), i.e. the ratio of the particle backscatter coefficients at 532 nm and at 1064 nm, is independent of the aerosol concentration. As it is directly related to the Angström exponent AE ( $AE = -\ln(CR)/\ln(532/1064)$ ) it provides information about the particle size. As a rule of thumb colour ratios around one indicate large particles (e.g. dust – Tesche et al., 2009), while colour ratios around two and higher are an indication for small particles compared to the lidar wavelength (e.g. smoke – Ansmann et al., 2009). For the layers of pollution aerosol observed during the first half of the EUCAARI-LONGREX campaign, a mean colour ratio of 2.4 was found, indicating small, submicron particles. More than 85 % of all data points show a colour ratio between 2.0 and 2.8, and less than 3.7 % show colour ratios lower than 2.0 (see Fig. 8d); an indication that no large particles were present.

### 3.3 Discussion

The characterization of polluted air masses is an active research topic concerning air quality and climate effect. Therefore a number of several papers report about measurements within polluted air masses. Lidar ratios of pollution aerosols over Central Europe were reported by Wandinger et al. (2002). They found values of 50–65 sr at 532 nm from Raman lidar measurements. Ansmann et al. (2000) found lidar ratios of 60–90 sr in polluted aerosol layers over

India, measured with a Raman lidar system during the field campaign INDOEX. Müller et al. (2001) reported about mean lidar ratios of about 60 sr for the same measurement experiment. During the ACE-2 field campaign Raman lidar measurements in polluted aerosol layers over Portugal showed lidar ratios of 50–70 sr (Ansmann et al., 2001). Xie et al. (2008) found values of  $61 \text{ sr} \pm 14 \text{ sr}$  in moderately polluted aerosol layers over Beijing, China. Hänel et al. (2012) found lidar ratios around 60 sr in pollution layers in and near Beijing, and Chen et al. (2012) found a mean value of the lidar ratio of about 58 sr in the Pearl River Delta. Using a  $180^\circ$  backscatter nephelometer lidar ratios of  $64 \text{ sr} \pm 4 \text{ sr}$  for pollution aerosol over Illinois (Anderson et al., 2000), and 60–70 sr for continental aerosols were observed (Doherty et al., 1999). Inversions from AERONET (Holben et al., 1998) sun-photometer measurements resulted in values of  $71 \text{ sr} \pm 10 \text{ sr}$  for industrial pollution over Europe, and over Middle and North America (Cattrall et al., 2005). Measurements over South-East Asia resulted in values of  $58 \text{ sr} \pm 10 \text{ sr}$  for industrial pollution aerosols (Cattrall et al., 2005). While a number of papers deal with the lidar ratio of industrial pollution or urban haze, measurements of the particle linear depolarisation ratio of this aerosol type are rare. Müller et al. (2007) reported about particle linear depolarisation ratios for urban haze of about or lower 5 % over Central and South-West Europe, and over North America. Burton et al. (2012) found values between 2–11 % over North America. In polluted air masses over Beijing, China, mean values of the particle linear depolarisation ratio of 7 % were found (Xie et al., 2008). Our values of  $56 \text{ sr} \pm 6 \text{ sr}$  and  $6\% \pm 1\%$  for pollution aerosols agree well with former findings.

#### 4 Summary and conclusion

During EUCAARI-LONGREX both the meteorological conditions and the flight strategy were best suited to study pollution aerosols with respect to location and age of the aerosols. The meteorological conditions were determined by a persistent high pressure system over the North Sea during the first half of the observing period and northerly advection during the second half. Especially during the first half high aerosol optical depth (AOD) values were observed due to the build-up of a pollution layer in the absence of precipitation. A total of 15 Falcon flight missions with 48 flight hours were performed. The measurements included observations with an airborne HSRL to characterize the optical properties and the spatial distribution of pollution aerosol. HSRL measurements are an excellent tool to characterize aerosols, as they provide high-resolved temporal and vertical information on the aerosol distribution and aerosol optical properties with a high accuracy. The main goal of this study was to determine lidar derived optical properties for European pollution aerosols, in particular the particle linear depolarization ratio and the lidar ratio. Both quantities are independent from the

aerosol load and, in combination, have the potential to distinguish different aerosol types as was recently shown by Groß et al. (2011, 2012) and Burton et al. (2012).

During all measurements included into this study, pollution aerosols over and from different locations across Europe and of different age could be observed (Hamburger et al., 2012). Over all, the measurements of the lidar ratio and the particle linear depolarization ratio of pollution aerosols showed a narrow range of values. The lidar ratio ranged between 48–65 sr in 90 % of the measurements points. For the particle linear depolarization ratio values between 4 and 8 % were found for 90 % of all measurement points. Based on our observations, the mean lidar ratio of the pollution aerosol is  $56 \text{ sr} \pm 6 \text{ sr}$ . The mean depolarisation ratio is  $6\% \pm 1\%$ . These results are in very good agreement to former findings. This is in particular valuable to define an aerosol type “pollution aerosol”, and to determine the right input parameter for aerosol retrievals, e.g. of satellite instruments such as on CALIPSO. Furthermore, the characterization of the aerosol type “pollution aerosol” is valuable to distinguish pollution aerosol from other aerosol types and to determine mixtures with other types of aerosols.

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