

Quantification of tidal grounding line migration using Sentinel-1 observations

Sindhu Ramanath^{1,2}, Michael Engel², Lukas Krieger¹, Dana Floricioiu¹ and Jan Wuite³

¹Remote Sensing Technology Institute, German Aerospace Center, Weßling, Germany

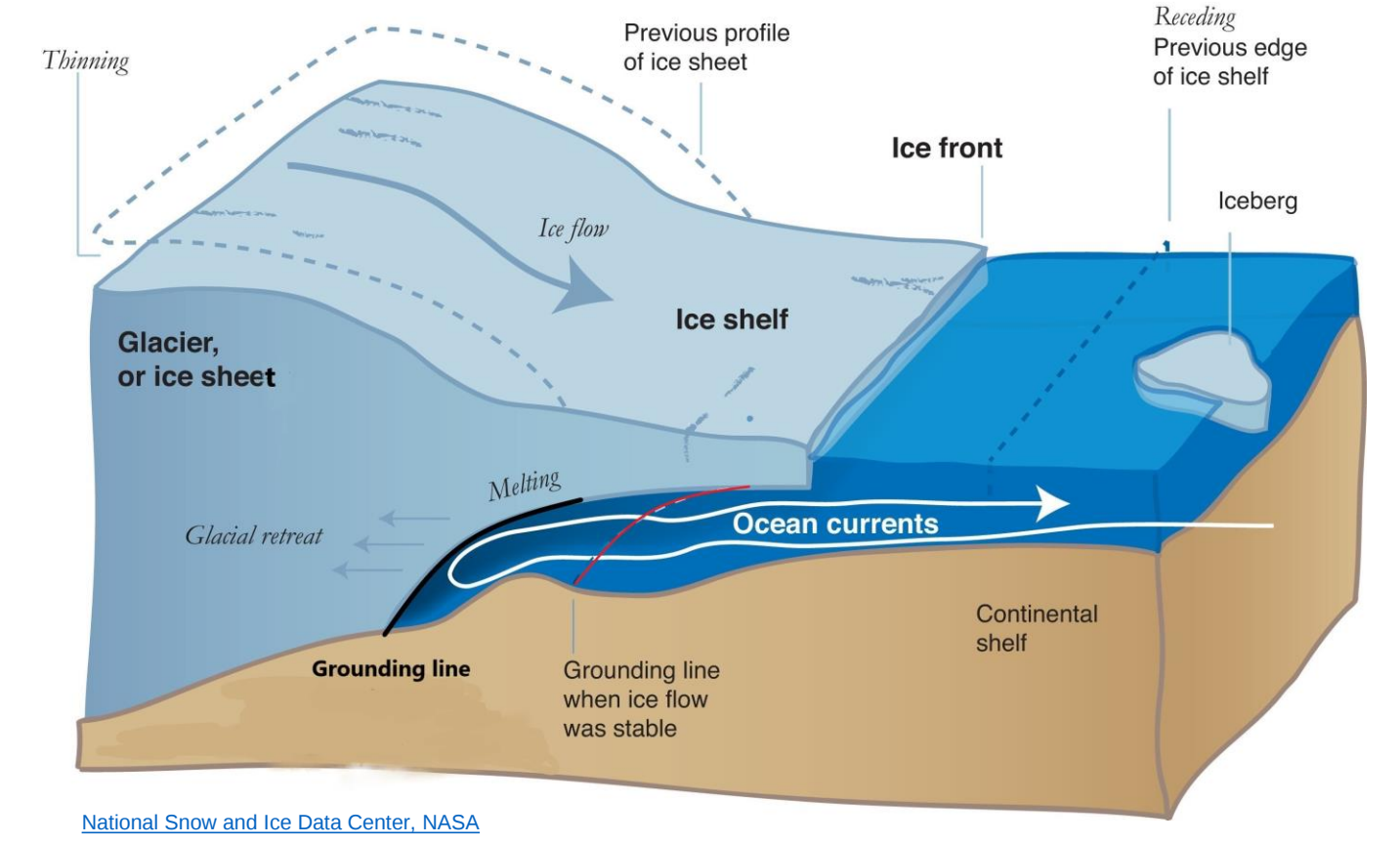
²School of Engineering and Design, Technical University of Munich, Germany

³ENVEO IT GmbH, Innsbruck, Austria

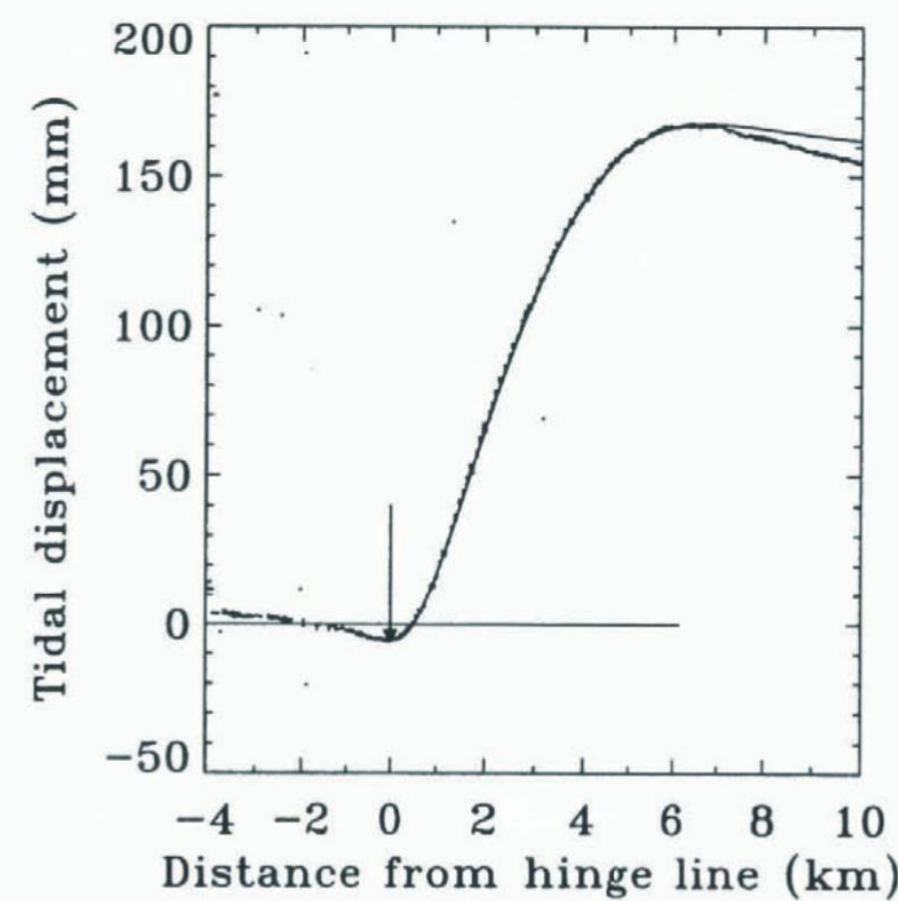
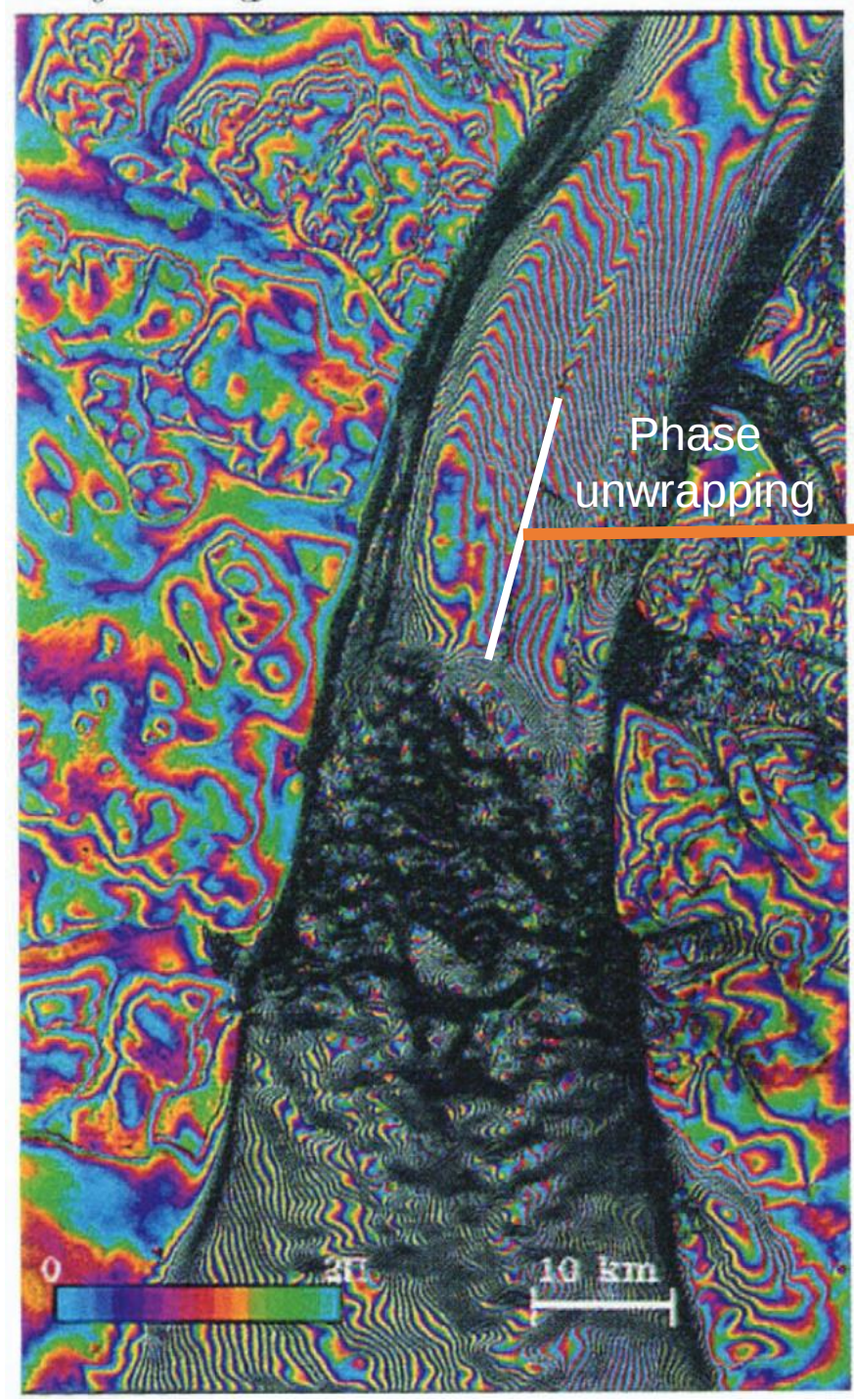
Motivation

Grounding lines (GLs) are subsurface geophysical features that represent the boundary between grounded ice and floating ice shelves. GLs derived from tidal remote sensing methods such as Differential Interferometric SAR, laser and radar altimetry contain an ephemeral displacement in addition to their true location. Previous works have demonstrated that grounding lines migrate with distances ranging from a few hundred meters to several kilometers heterogeneously and out of phase with ocean tides [1] – [2], implying that the tidal component does not diminish in an interannual time series.

We explore the use of SAR Differential Range Offsets Tracking (DROT) to provide insights into tidal migration of the grounding line. We used a times series from 2019 of LOS offsets from 6-day repeat cycle Sentinel-1 acquisitions over Larsen C Ice Shelf. GL positions from the offset profiles were derived by inverting the 1D elastic beam model using the Cross Entropy-based Importance Sampling for Bayesian Updating (CEBU) algorithm [3].



GL detection from InSAR phase [4]



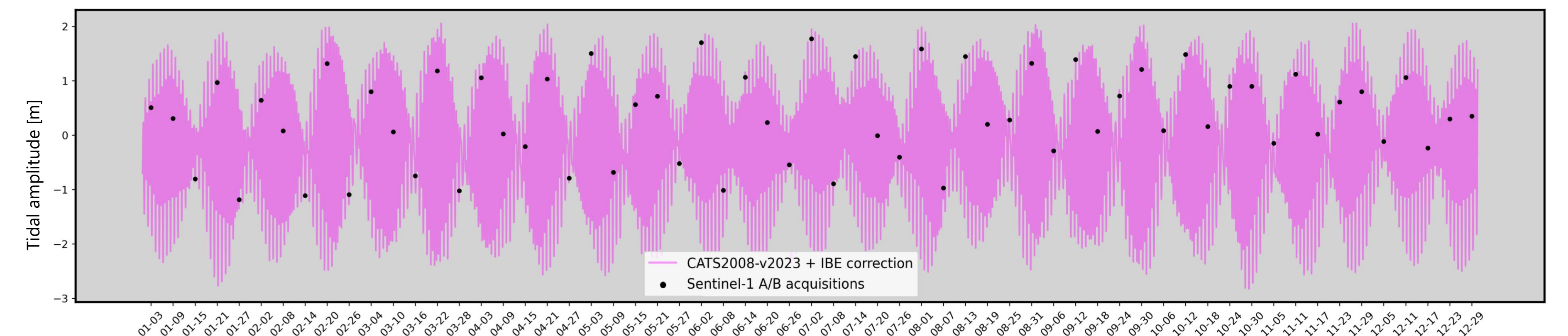
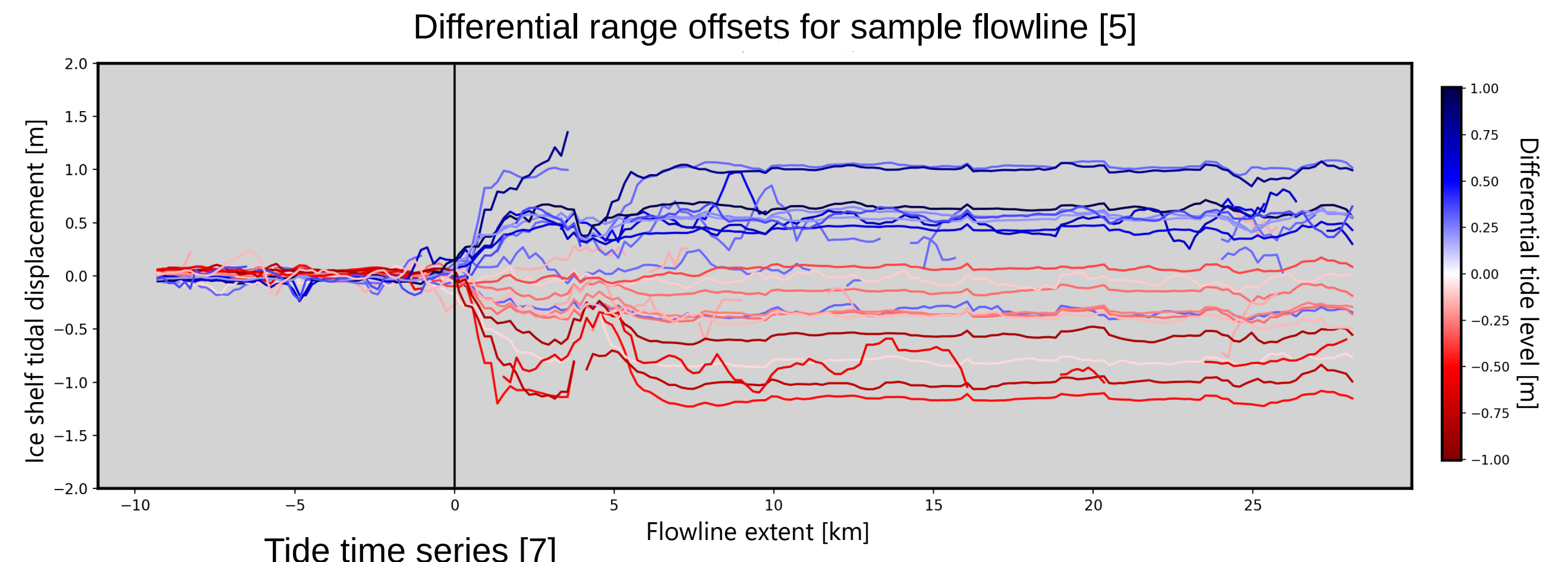
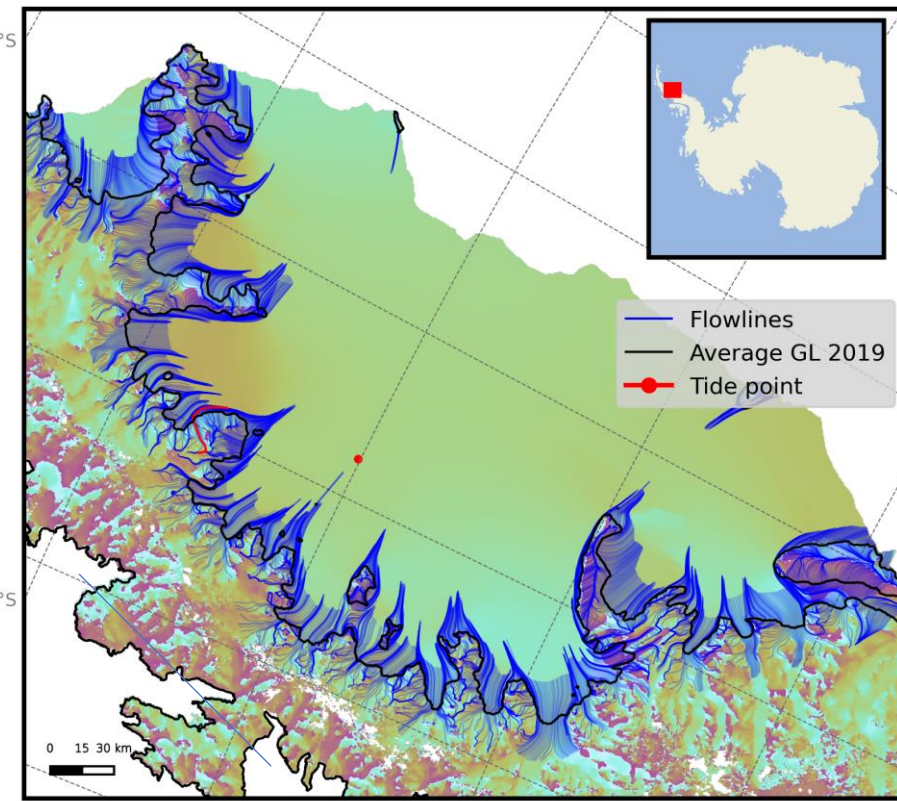
$w(x) = A_0(t)[1 - e^{-\beta x}(\cos \beta x + \sin \beta x)]$

x : horizontal axis
 $w(x)$: tidal deflection
 $A_0(t)$: level of ice shelf if floating in isostatic equilibrium
 $\beta = \sqrt{\frac{3\rho_{sea}g(1-\nu^2)}{Eh^3}}$
 h : ice thickness
 ν : Poisson's ratio
 E : Young's modulus
 g : acceleration due to gravity
 ρ_{sea} : density of sea water

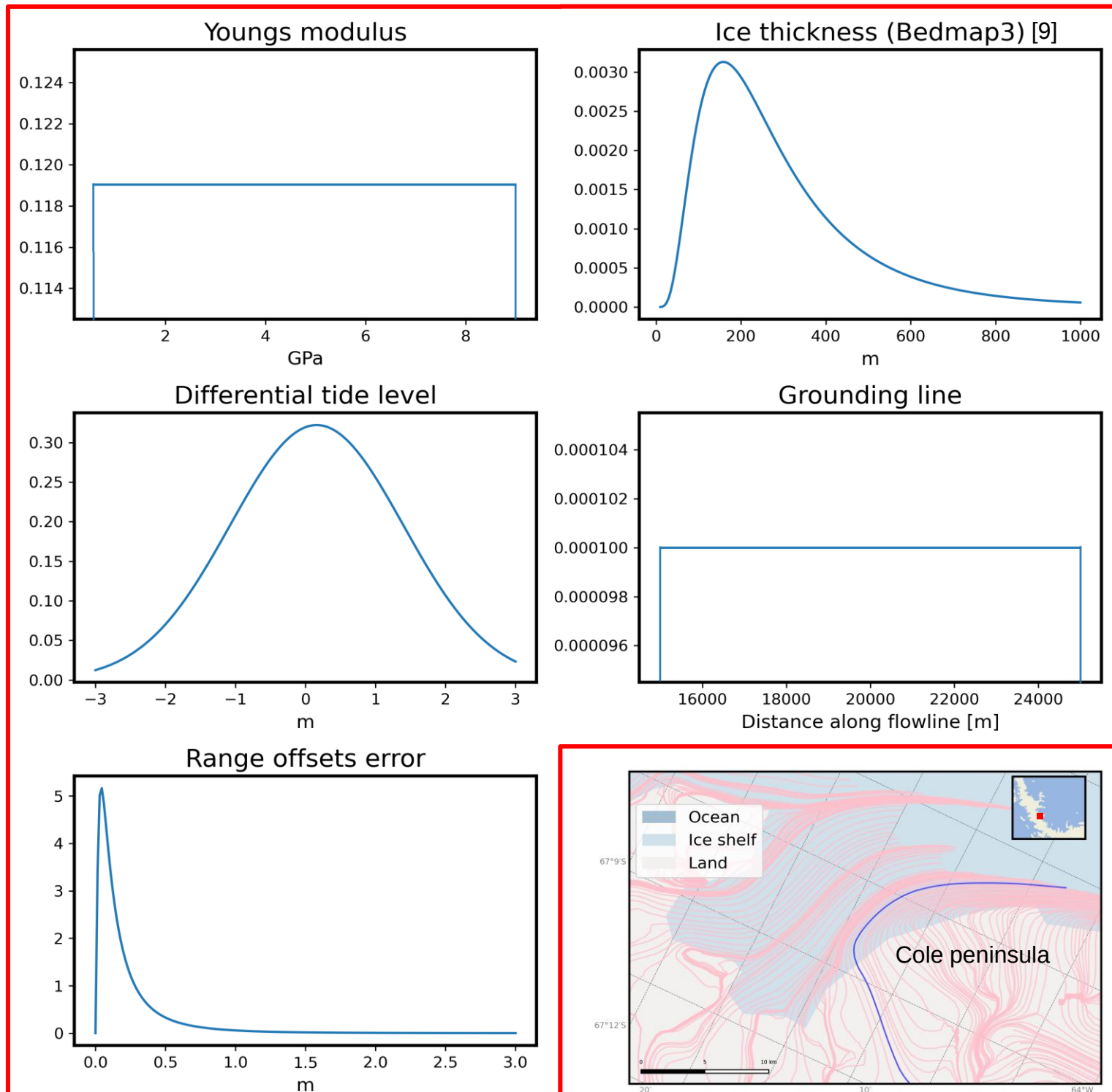
Large variation in E (0.1 – 10 GPa)
 Uncertainties in h and tide elevation!

Data

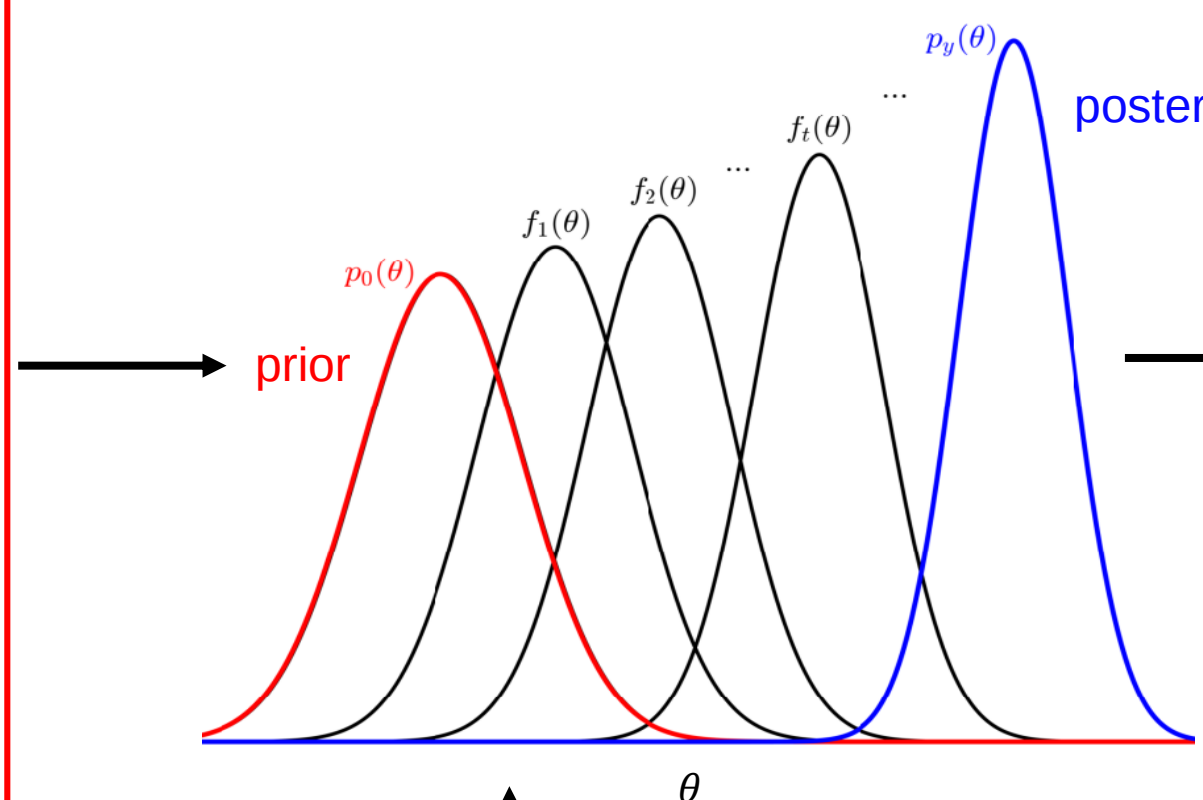
Ice flow direction [5] and average DROT GL [6]



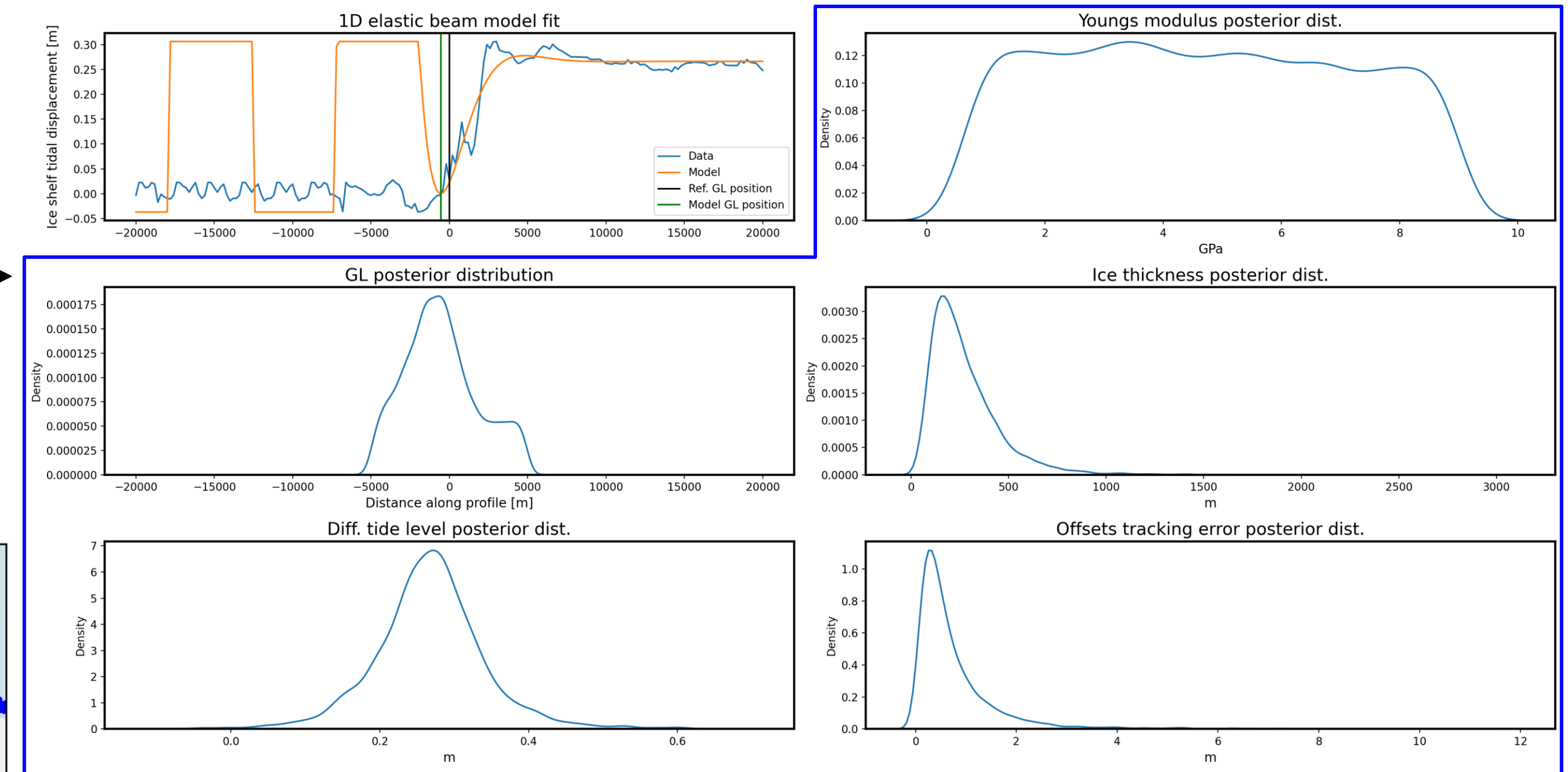
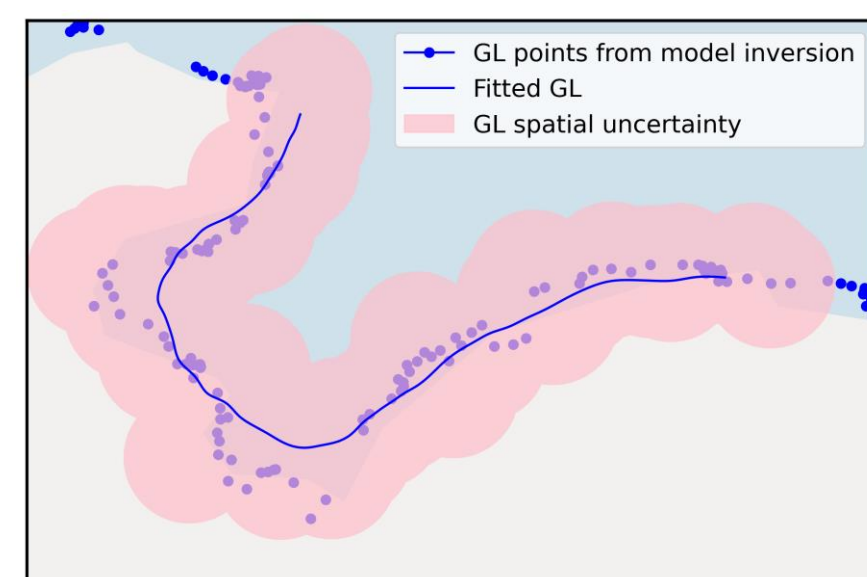
Workflow



CEBU [3]



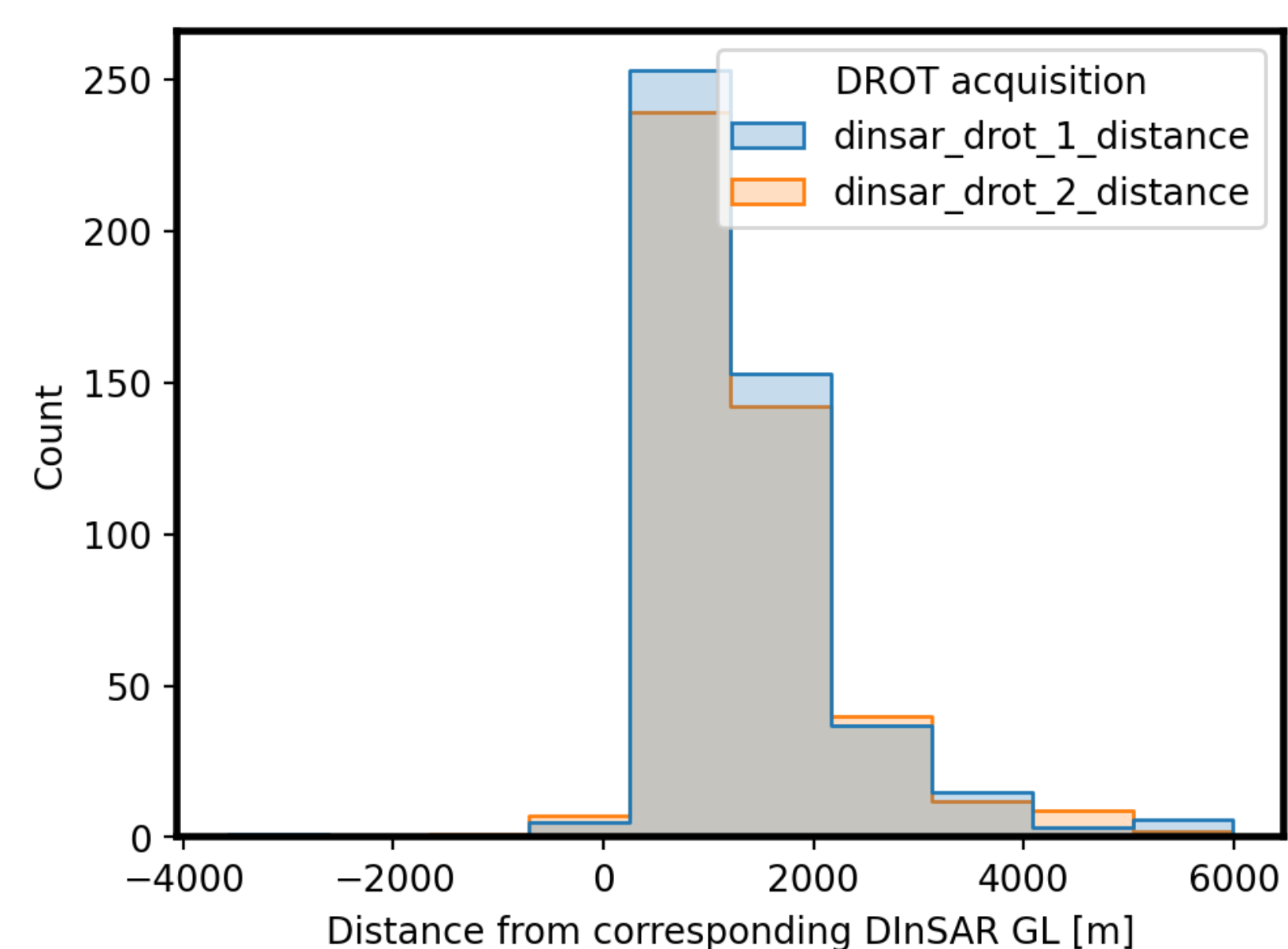
Differential range offsets for each flowline



Results

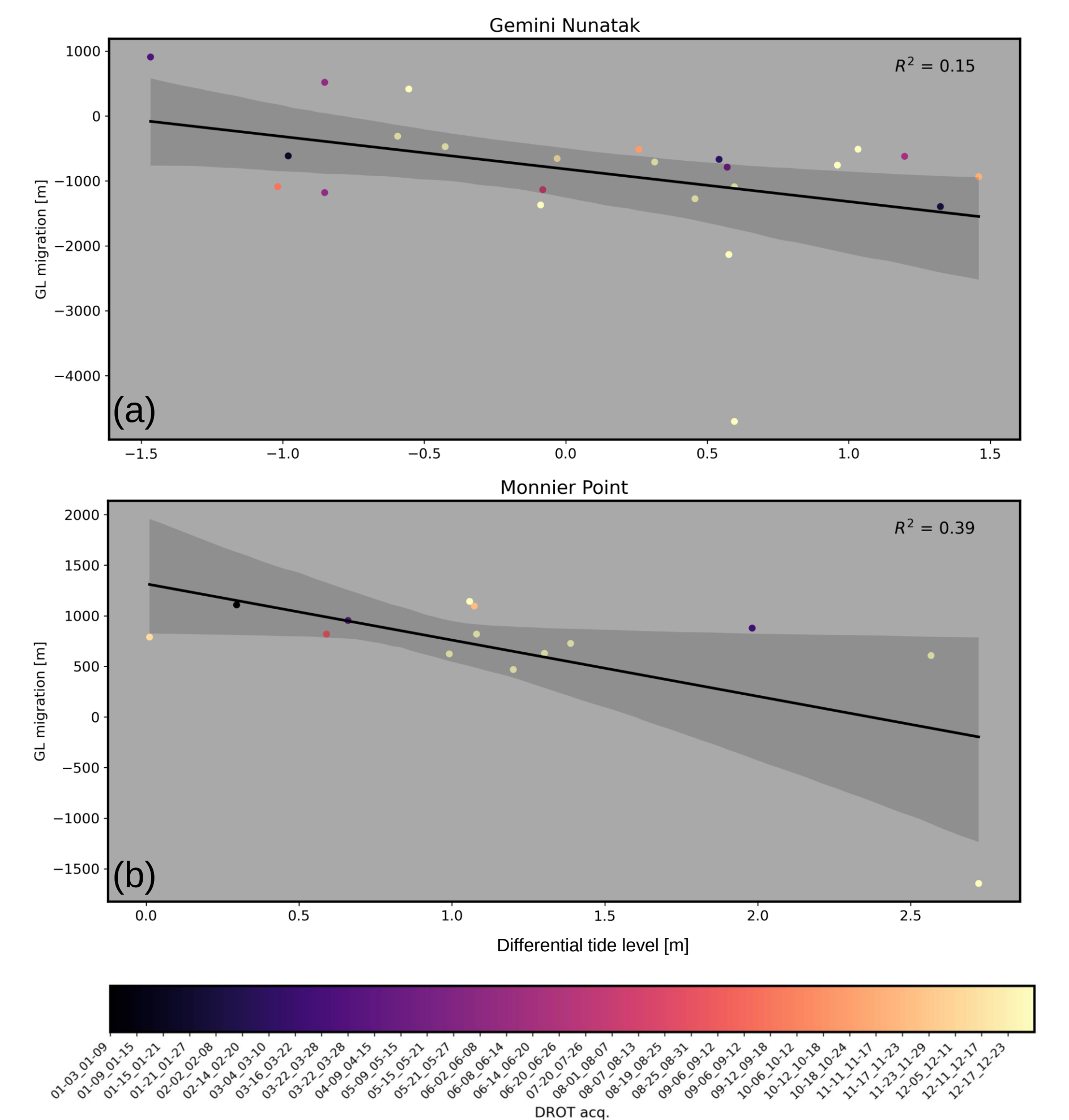
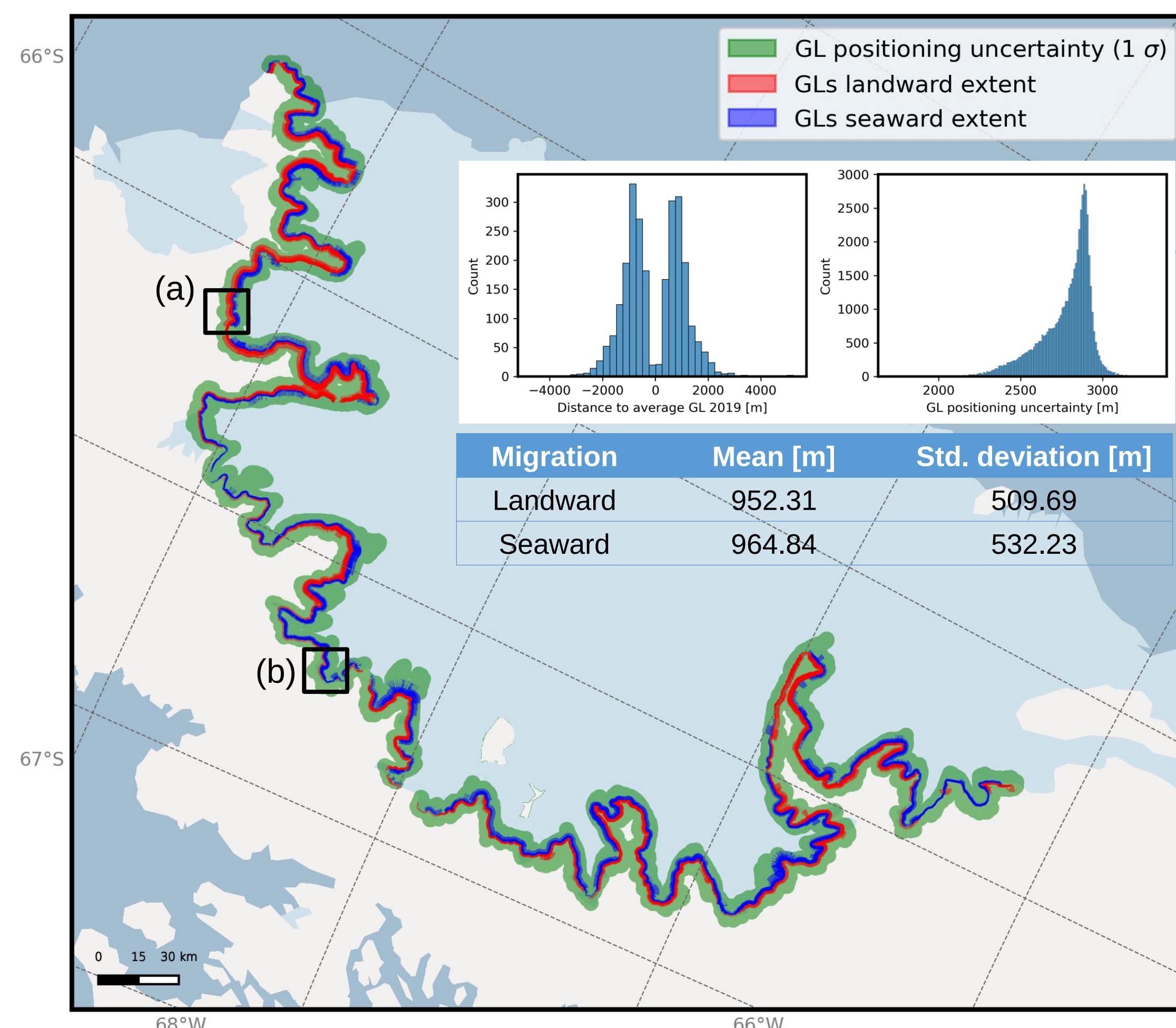
1. Comparison to DInSAR GLs

The distance between each DInSAR GL and the two DROT lines — one corresponding to each S1 acquisition pair used to form the DInSAR was measured.



DROT GLs are 1.35 ± 0.91 km seawards of DInSAR GLs, similar to the finding in [6].

2. GL migration w.r.t average GL



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[5] Nagler, T., Rott, H., Hetzenecker, M., Wuite, J., & Poin, P. (2015). The Sentinel-1 mission: New opportunities for ice sheet observations. *Remote Sensing*, 7(7), 9371–9389.

[6] Wallis, B. J., Hogg, A. E., Zhu, Y., & Hooper, A. (2024). Change in grounding line location on the Antarctic Peninsula measured using a tidal motion offset correlation method [Preprint]. *The Cryosphere*. <https://doi.org/10.5194/egusphere-2023-2874>

[7] Howard, S. L., Greene, C. A., Padman, L., Erofeeva, S., & Sutterley, T. (2024). "CATS2008_v2023: Circum-Antarctic Tidal Simulation 2008, version 2023" U.S. Antarctic Program (USAP) Data Center. doi: <https://doi.org/10.15784/601772>

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[9] Pritchard, H., et al., (2024). BEDMAP3 - Ice thickness, bed and surface elevation for Antarctica - gridding products (Version 1.0) [Data set]. NERC EDS UK Polar Data Centre. <https://doi.org/10.5285/2d0e4791-8e20-46a3-80e4-f5f6716029d2>