

APPLICATION OF DUAL-FREQUENCY EISCAT MEASUREMENTS TO DETERMINE ION-NEUTRAL COLLISION FREQUENCIES WITH THE DIFFERENCE SPECTRUM METHOD

21ST INTERNATIONAL EISCAT SYMPOSIUM, TROMSØ, NORWAY

02 AUGUST 2024

Florian Günzkofer

Gunter Stober

Johan Kero

David R. Themens

Yasunobu Miyoshi

Dimitry Pokhotelov

Claudia Borries

Institute for Solar-Terrestrial Physics, DLR

Institute of Applied Physics, University of Bern

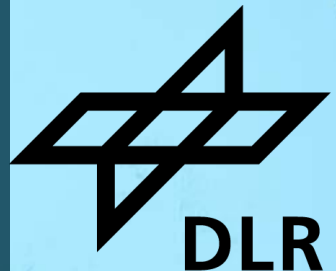
Swedish Institute for Space Physics

University of Birmingham

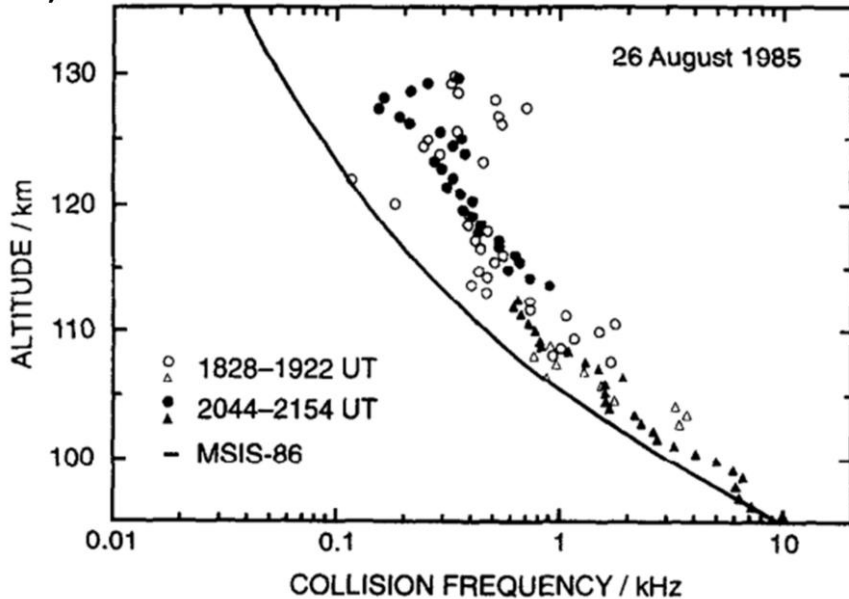
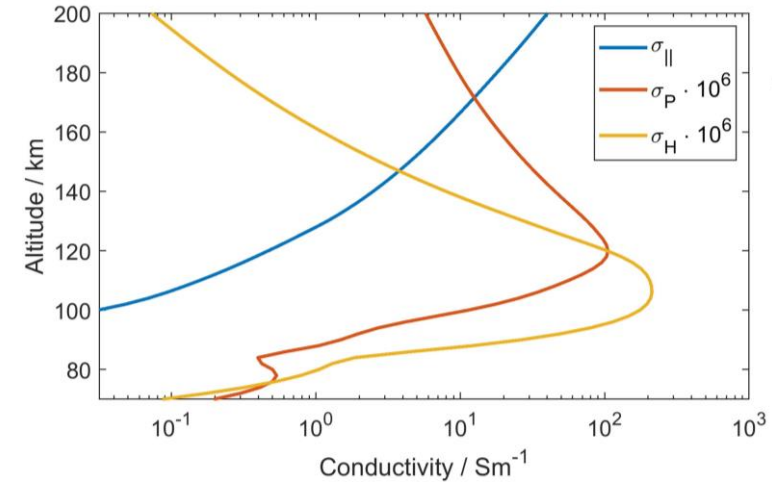
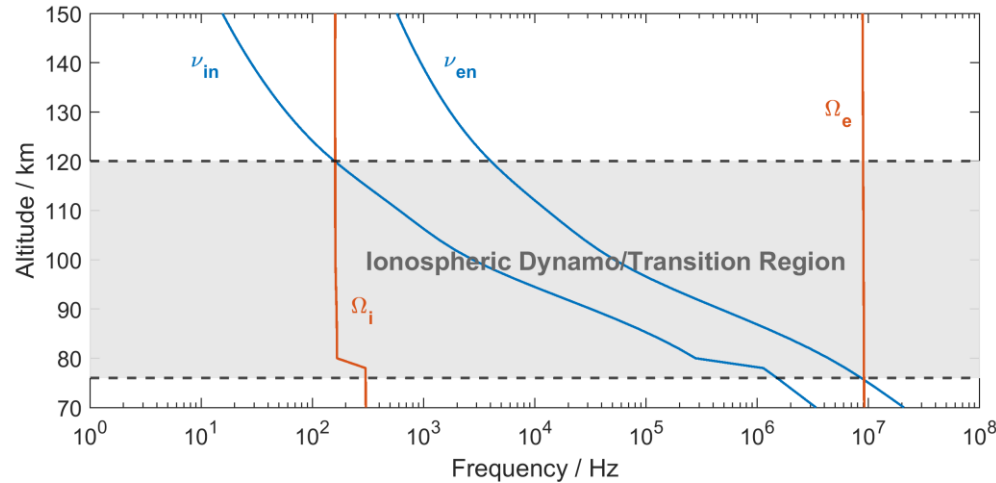
Kyushu University

Institute of Physics, University Greifswald

Institute for Solar-Terrestrial Physics, DLR



The role of the ion-neutral collision frequency

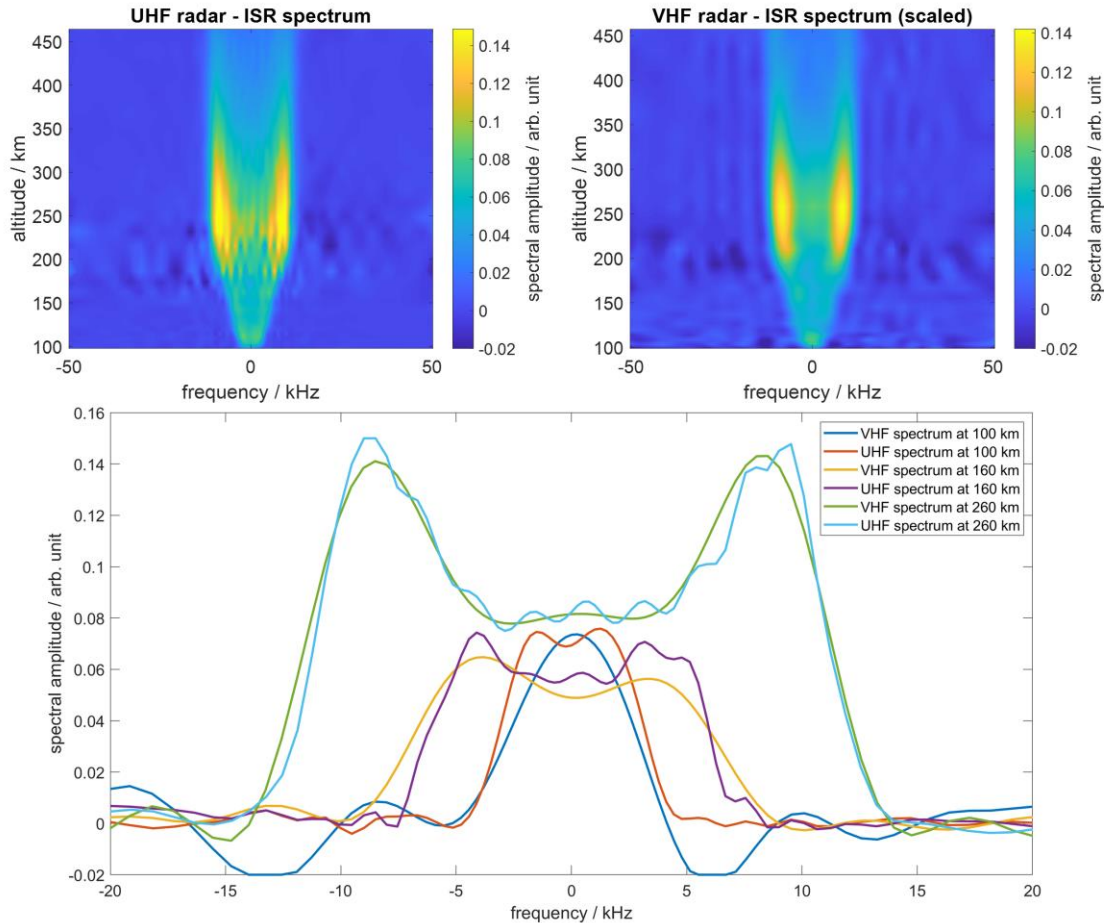


- ion-neutral collision frequency ν_{in} determines the transition from a collisional to a collision-less ionosphere
 - determines maxima of ionospheric conductivities σ_P and σ_H
- ν_{in} impacts the ISR spectrum due to the ion adiabatic coefficient γ_i
 - $\gamma_i T_i$ term is ambiguous, assumptions ($T_i = T_e$) required
- ν_{in} determines the drag force between ion and neutral particles
 - vertical ion momentum equation can be applied ($U_z = 0, E \gtrsim 20 \text{ mV/m}$)
- previous measurements indicated **considerable differences between climatology (MSIS) and measurement ν_{in} profiles** (Nygrén, 1996; Oyama *et al.*, 2012)

[Nygrén, *Adv. Space Res.*, **18**, 79-82, (1996).]

The difference spectrum method

[Grassmann, *J. Atmos. Terr. Phys.*, **55**, 573-576, (1993).]



[Günzkofer *et al.*, *Atm. Meas. Tech.*, **16**, 5897-5907, (2023).]

Measured difference function:

$$D(\omega_{UHF} + \delta\omega) = S(\omega_{UHF} + \delta\omega) - \beta \cdot \tilde{S}(\omega_{UHF} + \delta\omega)$$

β determined for $D(\omega_{UHF} + \delta\omega) = 0$ (F region)

scaled VHF spectrum:
$$\zeta = \frac{\omega_{UHF}}{\omega_{VHF}} = 4.15$$

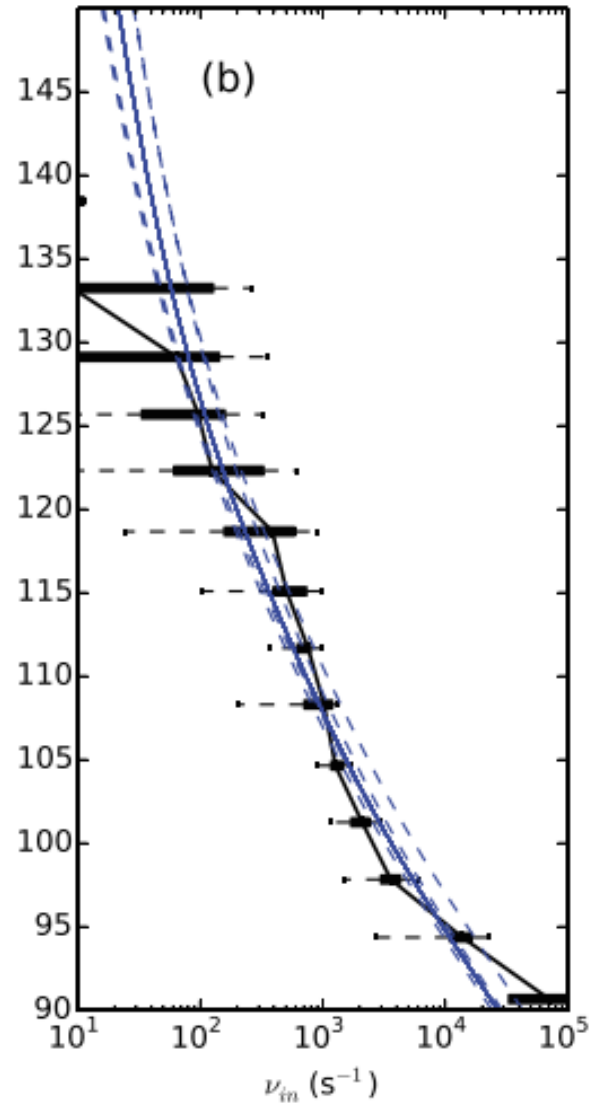
$$\tilde{S}(\omega_{UHF} + \delta\omega) = \begin{cases} \zeta^2 \cdot S(\omega_{VHF} + \delta\omega) \\ S(\omega_{UHF} + \delta\omega, \zeta^2 \cdot N_e, T_e, T_i, \zeta \cdot \mathbf{v}_{in}, v_i) \end{cases}$$

Theoretical difference function:

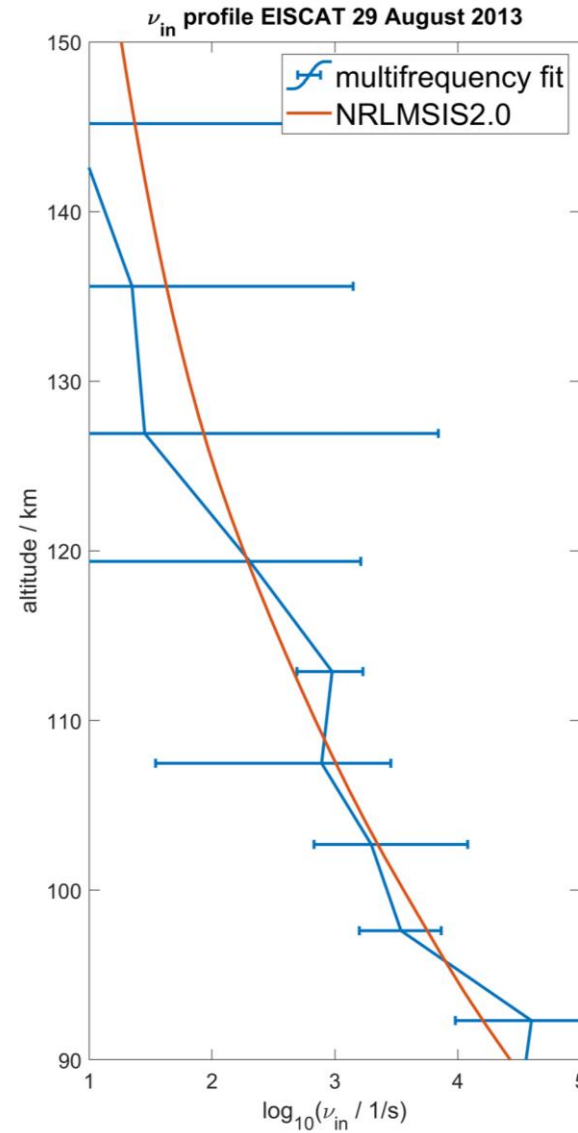
$$d(\omega_{UHF} + \delta\omega, N_e, T_e, T_i, \mathbf{v}_{in}, v_i) = s(\omega_{UHF} + \delta\omega, N_e, T_e, T_i, \mathbf{v}_{in}, v_i) - \beta \cdot s(\omega_{UHF} + \delta\omega, \zeta^2 \cdot N_e, T_e, T_i, \zeta \cdot \mathbf{v}_{in}, v_i)$$

Comparison to Nicolls et al., 2014

08/29/2013 0703-1054 UT



[Nicolls et al., *Geophys. Res. Lett.*, **41**, 8147-8154, (2014).]



[Günzkofer et al., *Atm. Meas. Tech.*, **16**, 5897-5907, (2023).]

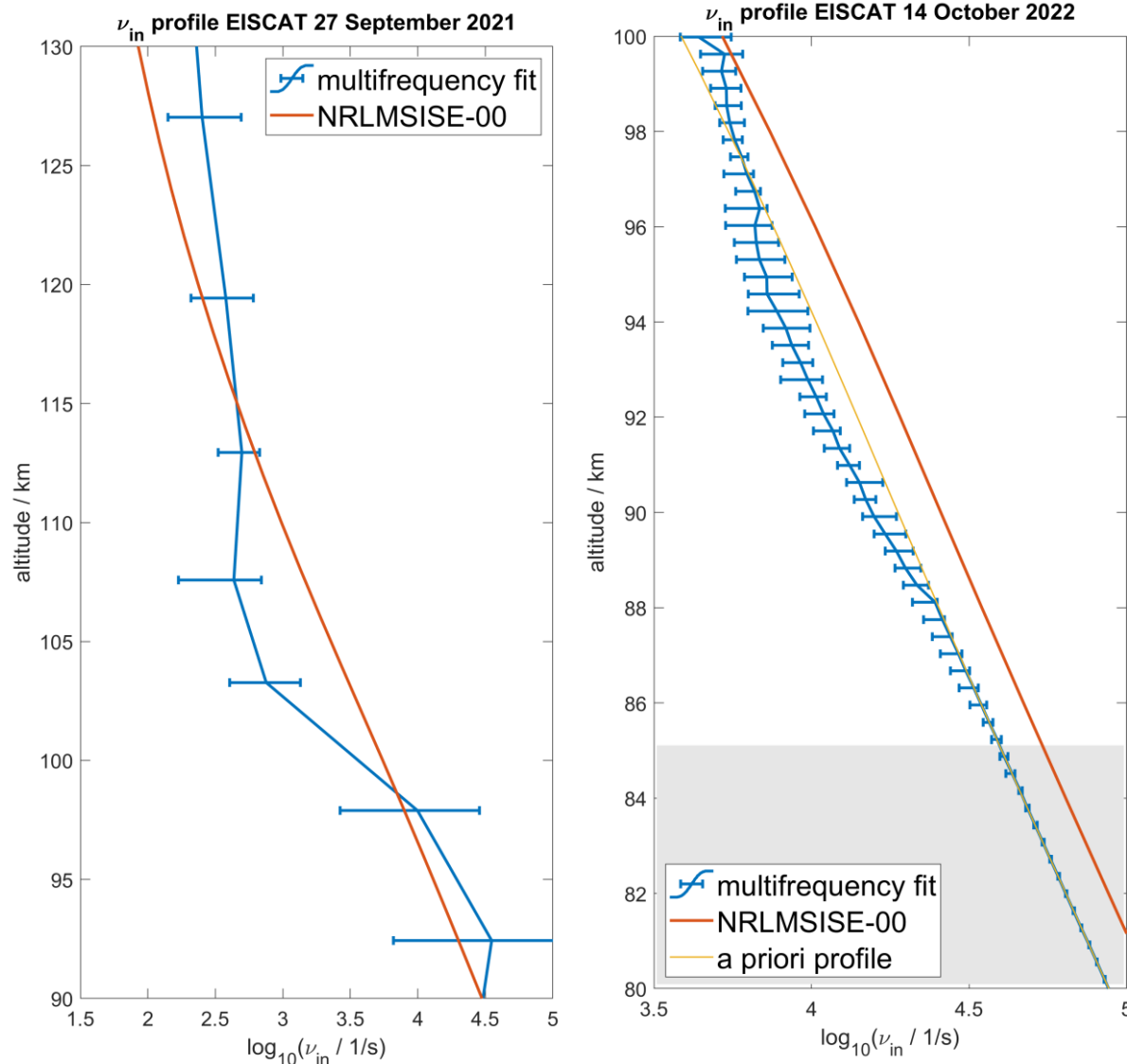
Nicolls et al., 2014:

- dual-frequency EISCAT campaign from 29 August 2013 analyzed following Grassmann, 1993b
- simultaneous fit of both spectra with combined error function applied
→ not possible with GUIDAP

Difference spectrum method:

- separate single-frequency analysis
→ analysis possible with GUIDAP
- obtained median profile strongly resembles Nicolls et al., 2014
→ interquartile errorbars are considerably increased

Analyzed dual-frequency campaigns



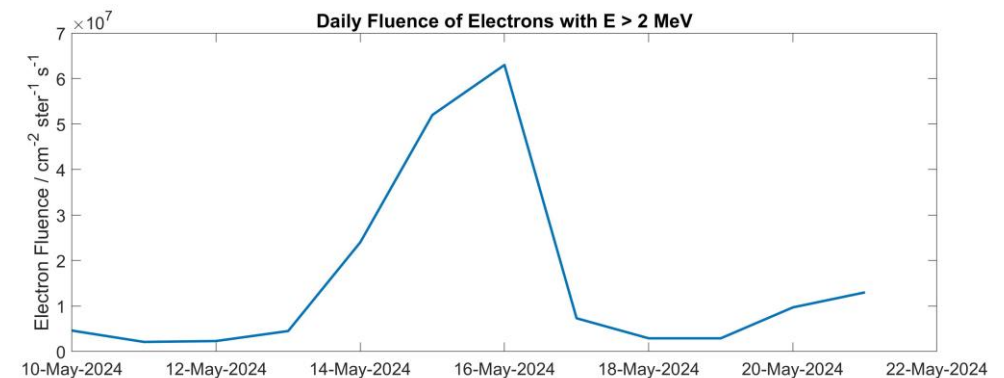
[Günzkofer et al., *Atm. Meas. Tech.*, **16**, 5897-5907, (2023).]

DLR EISCAT dual-frequency campaigns:

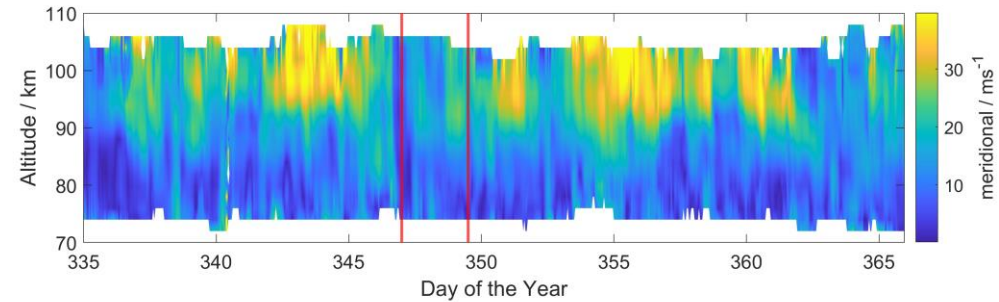
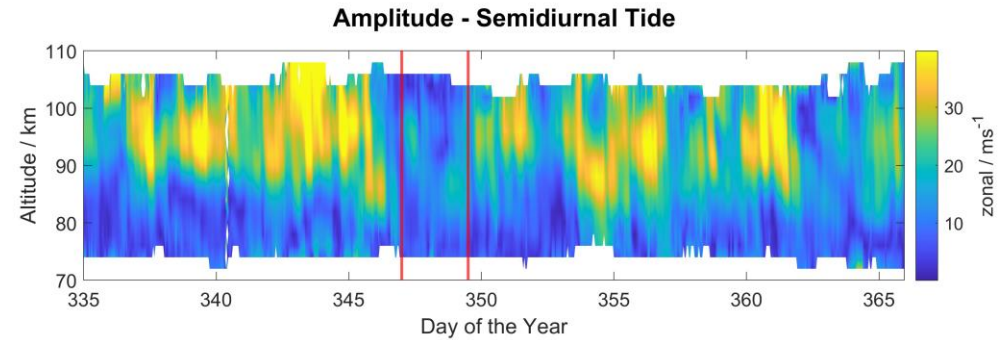
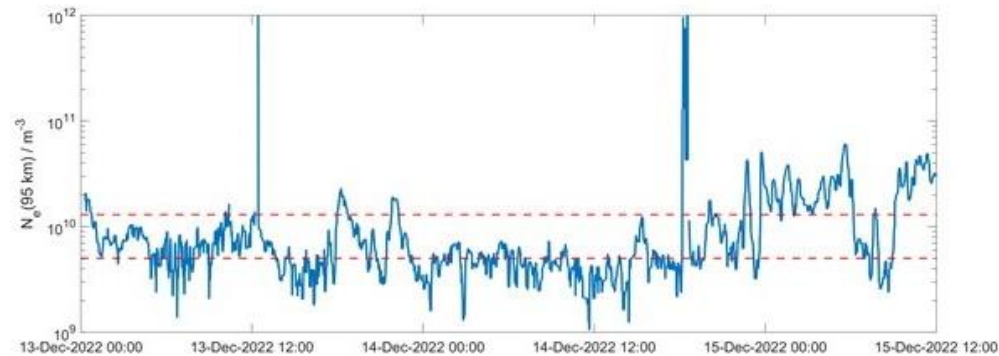
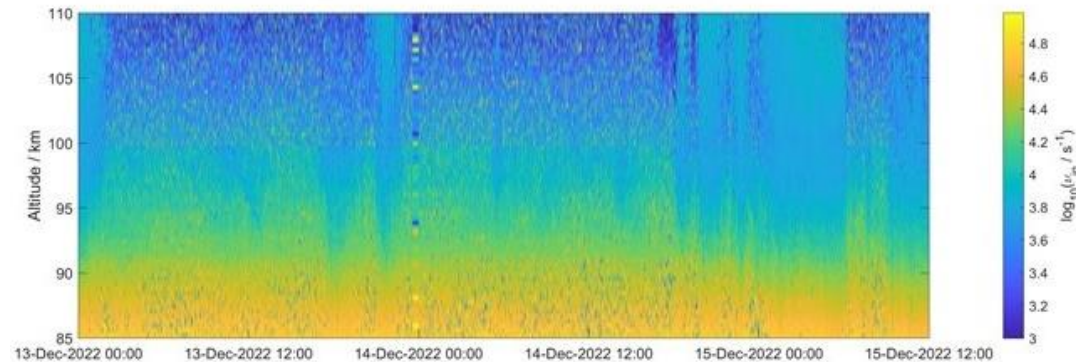
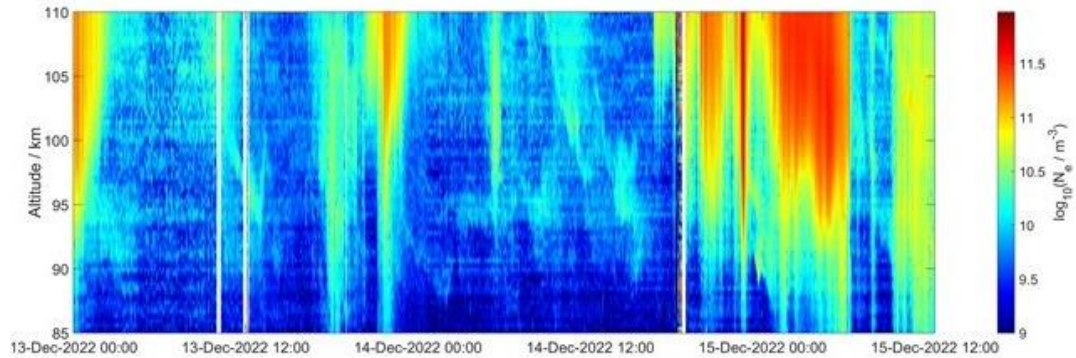
- 27 September 2021
beata, el 45°, az 180°
- 14 October 2022
manda zenith (CP 6)

Other EISCAT dual-frequency campaigns:

- Geminids 13-15 December 2022 (Sweden)
manda zenith (CP 6)
- SEP Event 16 May 2024 (UK)
manda zenith (CP 6)



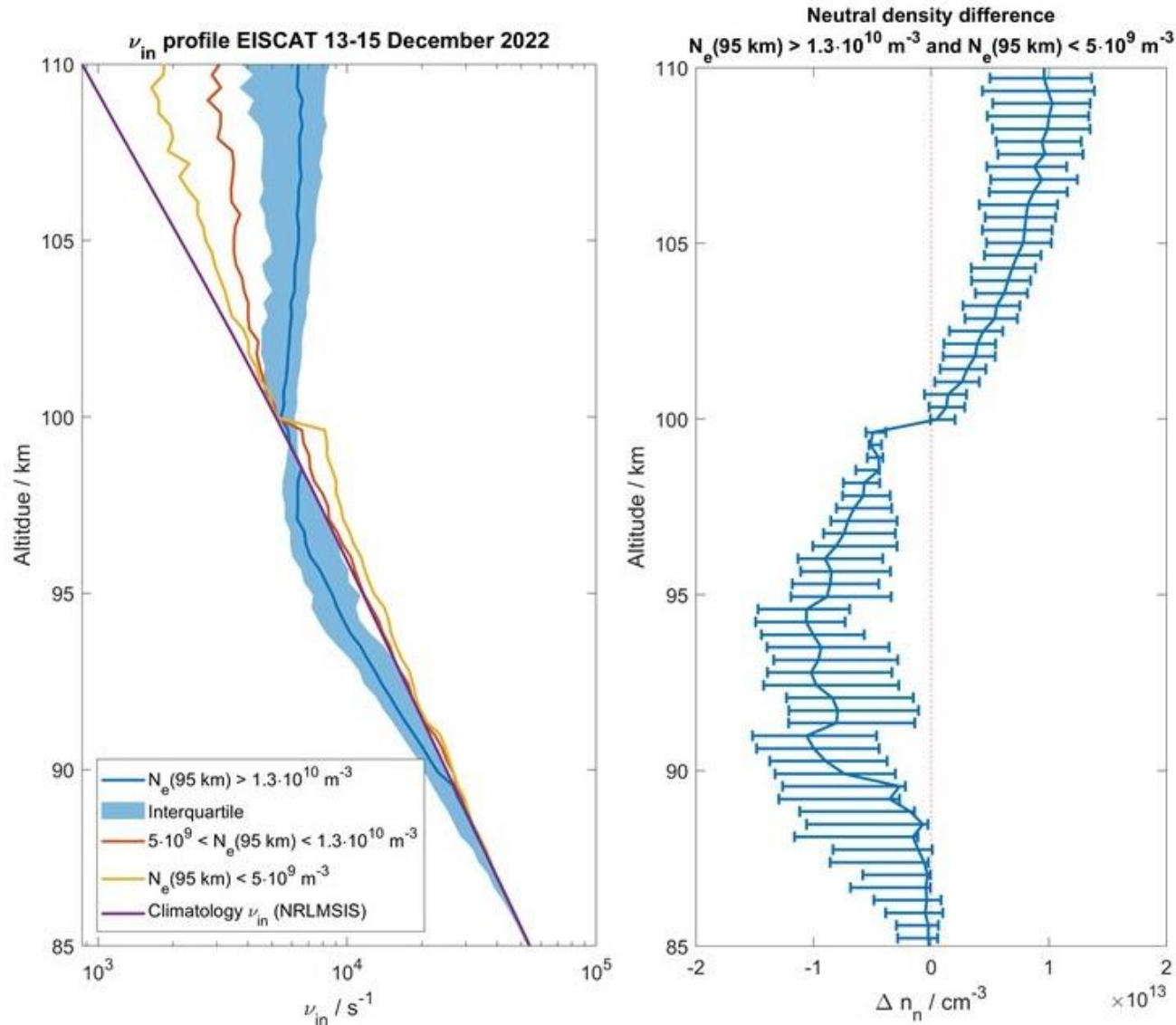
Impact of particle precipitation – Geminids campaign 2022



[Günzkofer *et al.*, in preparation]

- original idea: investigate impact of atmospheric tides, **but**: significantly low tidal amplitudes during campaign
- strong particle precipitation detected
- electron density at 95 km altitude $N_e(95 \text{ km})$ is applied as proxy for particle precipitation

Impact of particle precipitation – neutral upwelling

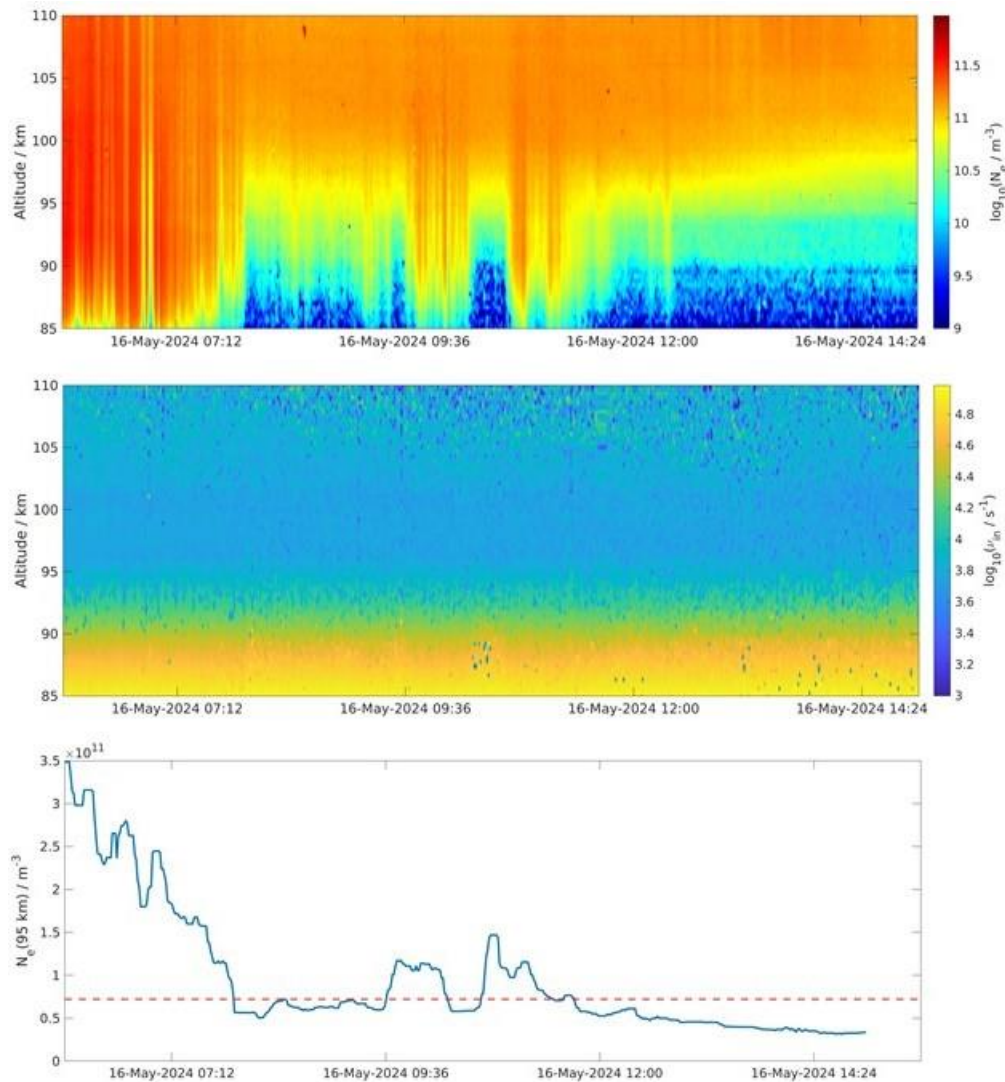


- collision frequency profiles binned for:

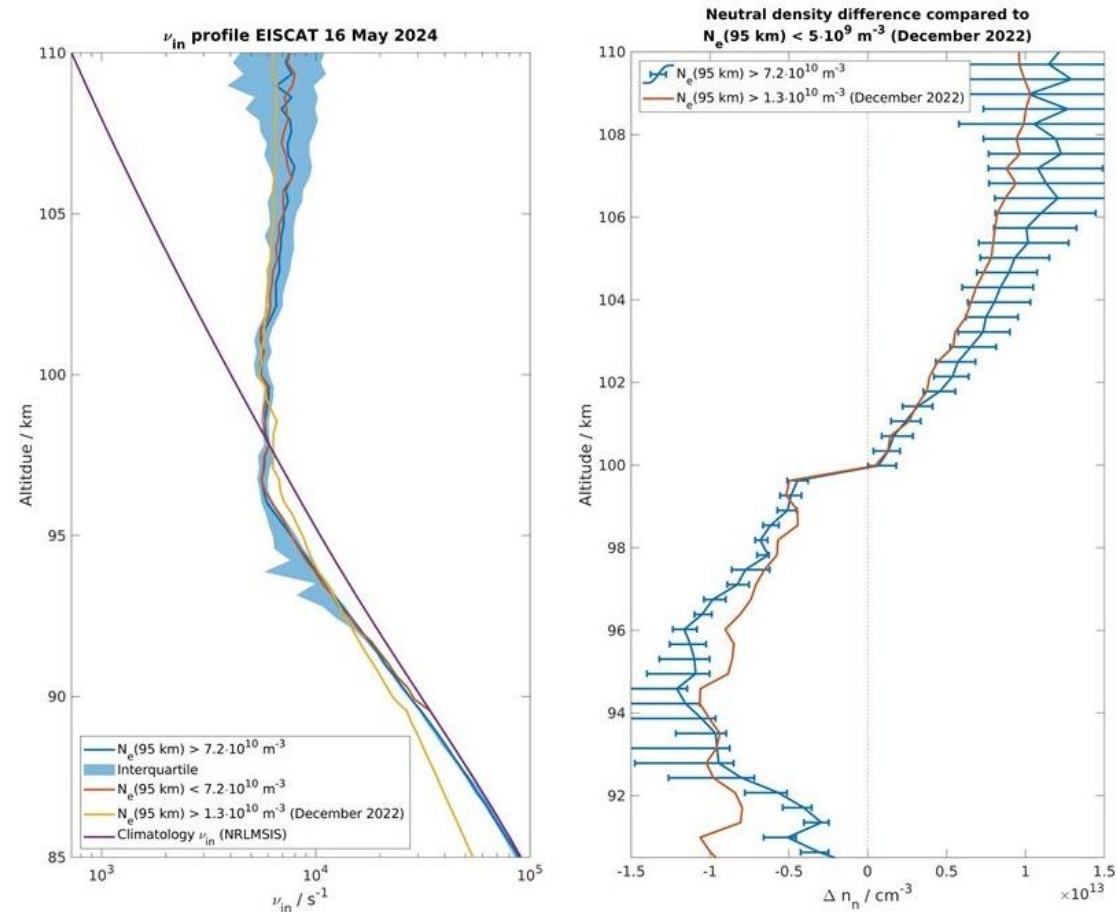
1. $N_e(95 \text{ km}) < 5 \cdot 10^9 \text{ m}^{-3}$
2. $5 \cdot 10^9 < N_e(95 \text{ km}) < 1.3 \cdot 10^{10} \text{ m}^{-3}$
3. $N_e(95 \text{ km}) > 1.3 \cdot 10^{10} \text{ m}^{-3}$

- neutral particle density profile calculated from collision frequency
- difference of neutral particle density profile for high and low particle precipitation calculated
- neutral atmosphere heating at 90 – 100 km altitude with consequent atmospheric upwelling

Impact of particle precipitation – SEP event May 16, 2024



[Günzkofer *et al.*, in preparation]



[Günzkofer *et al.*, in preparation]

- ➔ profiles similar to December 2022 measurement for $N_e(95 \text{ km}) > 1.3 \cdot 10^{10} \text{ m}^{-3}$
- ➔ neutral uplift becomes more pronounced for stronger particle precipitation

Summary



→ dual-frequency ISR campaigns allow for ν_{in} measurements without restrictions on the state of the ionosphere

Difference spectrum method

Advantages	Disadvantages
based on standard ISR analysis software (GUISDAP)	increased uncertainties due to two-step analysis (1. single-frequency ISR fit, 2. difference spectrum fit)
easily adaptable for different radar modes	β parameter required to compensate technical differences of the two radars

→ investigation of the neutral atmosphere in the MLT region possible since $n_n \sim \nu_{in}$

→ ionospheric conductivities and currents are strongly impacted by the collision frequency which therefore have a direct impact on the space weather

florian.guenzkofer@dlr.de

References:

Günzkofer *et al.*, *Atm. Meas. Tech.*, **16**, 5897-5907, (2023).

Nicolls *et al.*, *Geophys. Res. Lett.*, **41**, 8147-8154, (2014).

Oyama *et al.*, *J. Geophys. Res.*, **117**, A05308, (2012).

Nygrén, *Adv. Space Res.*, **18**, 79-82, (1996).

Grassmann, *J. Atmos. Terr. Phys.*, **55**, 573-576, (1993).