

### REVIEW ARTICLE OPEN



# Optical observations of thunderstorms from the International Space Station: recent results and perspectives

Torsten Neubert 1 → Francisco J. Gordillo-Vázquez and Heidi Huntrieser 3

The International Space Station (ISS) is in the lowest available orbit at ~400 km altitude, bringing instruments as close to the atmosphere as possible from the vantage point of space. The orbit inclination is 51.6°, which brings the ISS over all the low- and mid-latitude regions of the Earth and at all local times. It is an ideal platform to observe deep convection and electrification of thunderstorms, taken advantage of by the Lightning Imaging Sensor (LIS) and the Atmosphere Space Interaction Monitor (ASIM) experiments. In the coming years, meteorological satellites in geostationary orbit (~36,000 km altitude) will provide sophisticated cloud and lightning observations with almost complete coverage of the Earth's thunderstorm regions. In addition, Earth-observing satellite instruments in geostationary- and low-Earth orbit (LEO) will measure more atmospheric parameters at a higher resolution than we know today. The new infrastructure in space offers an opportunity to advance our understanding of the role of thunderstorms in atmospheric dynamics and climate change. Here, we discuss how observations from the ISS or other LEO platforms with instruments that view the atmosphere at slanted angles can complement the measurements from primarily nadiroriented instruments of present and planned missions. We suggest that the slanted viewing geometry from LEO may resolve the altitude of electrical activity and the cloud structure where they occur, with implications for modelling thunderstorms' effects on the atmosphere's radiative properties and climate balance.

npj Microgravity (2023)9:12; https://doi.org/10.1038/s41526-023-00257-4

#### INTRODUCTION

Thunderstorms develop primarily at low- and mid-latitudes, where the solar energy input is the largest. The atmosphere over land is heated unequally depending on the underlying surface, and thermal bubbles develop<sup>1</sup>. Some may rise as deep convection to the upper tropopause, occasionally even into the lower stratosphere<sup>2</sup>. Deep convection carries high amounts of water vapour, dust, aerosols, and trace gases from the polluted boundary layer that may reside at high altitudes for times much longer than the duration of the storm (days versus hours). Spreading over extended regions (~100-1000 km), they perturb the radiative properties of the upper troposphere and lower stratosphere (UTLS) region<sup>3–5</sup>. Lightning affects trace species' concentrations by specific chemical reactions in the heated lightning channel<sup>6–8</sup>. It can cause deaths and injuries, crop and property damage, and may ignite wildfires<sup>9</sup> that release huge quantities of trace species (including greenhouse gases) into the atmosphere. Locally and temporally these amounts might exceed anthropogenic emission of such gases<sup>10</sup>. The emissions may be injected into the stratosphere<sup>11,12</sup> and are important contributors to global warming and climate change <sup>13</sup>. Concurrently, climate change increases the frequency of hot and dry weather situations that fuel wildfires. Studies on lightning activity in a warmer climate suggest that the average global activity may decrease because of a diminishing amount of hail in thunderstorms 14, whereas regional activities may increase<sup>15,16</sup>. Especially in the high Arctic region, where wildfires are easily induced, a drastic rise in lightning activity has been observed<sup>17</sup>, causing a rapid release of trace species with limited possibilities to quench the fires. For these reasons, the World Meteorological Organisation included lightning as an essential climate variable 18. It is of interest, then, to understand how thunderstorms and lightning activity, on the one hand, affect the climate balance and, on the other hand, how they are affected by a changing climate. Here, the ISS is in a well-suited orbit, passing over all the major thunderstorm regions of the Earth within  $\pm 51.6^{\circ}$  latitude, as shown in Fig. 1.

In the following, we give a brief account of the development of optical observations of lightning activity from space. We discuss some aspects of how thunderstorms may affect the radiative properties of the atmosphere and point out that the altitude of perturbations to atmosphere of greenhouse gases is important. We end by discussing the opportunities for using the ISS, or another platform in LEO, for future observations with slanted viewing geometry that allows for measurements of thunderstorm activity with resolution in altitude, complementing the primarily nadir-viewing measurements of current and planned missions.

### **OBSERVATIONS**

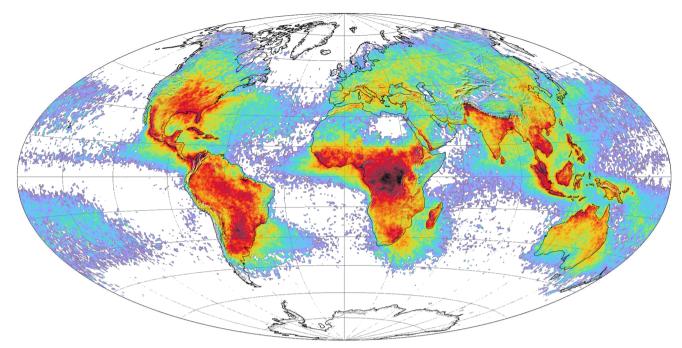
## The first instruments for optical lightning detection from space

Optical observations of lightning from low-earth orbit were first conducted by the Optical Transient Detector (OTD) on the MicroLab-1 satellite launched in 1995 into a 70° orbit at 740 km altitude<sup>19</sup>. The instrument represented a technological breakthrough because it could measure lightning flashes during the day in a background of light reflected from solar-illuminated clouds. This capability was achieved by selecting a narrow spectral band (~1 nm) around a strong line of the lightning spectrum (an atomic oxygen line at 777.4 nm), a pixel resolution that matches the cloud illumination by a flash (~10 km²), and a time resolution that fits the typical duration of optical pulses (2 ms). Excited atomic

<sup>&</sup>lt;sup>1</sup> National Space Institute, Technical University of Denmark (DTU Space), Elektrovej bld. 328, 2800 Kongens Lyngby, Denmark. <sup>2</sup>Instituto de Astrofísica de Andalucía (IAA - CSIC), Glorieta de la Astronomía s/n, 18008 Granada, Spain. <sup>3</sup>Institute of Atmospheric Physics, German Aerospace Center, Münchner Straße 20, 82234 Oberpfaffenhofen-Wessling, Germany. <sup>™</sup>Germany. <sup>™</sup>Germ







**Fig. 1 Global annual lightning flash density rates.** The colours mark flashes/km²/year from 0,1 (magenta) to 70 (black). The data are from the OTD and TRMM-LIS (1995-2003). Based on a NASA image by Marit Jentoft-Nilsen, with data provided by the Global Hydrology and Climate Center Lightning Team<sup>102</sup>.

oxygen is almost absent in the troposphere except in the high-temperature channels of lightning leaders, and is a unique signature of a lightning flash. The OTD gave the first global map of lightning climatology by a single sensor with good statistics of all the thunderstorm regions of the Earth at all local times and seasons<sup>19</sup>.

The Lightning Imaging Sensor (LIS), based on the OTD, was included in the payload of the Tropical Rainfall Measuring Mission (TRMM) satellite. TRMM was launched in 1997 into an orbit at 350 km altitude and 35° inclination. The orbit was raised to 402 km in 2001 and re-entered the atmosphere in 2015<sup>20,21</sup>. LIS data were used to detect deep convection without land-ocean bias, estimate the precipitation mass in the mixed-phase region of thunder-clouds, and differentiate storms with strong updrafts from those with weak vertical motion.

The OTD and LIS data provided a wealth of information on the climatology of lightning day and night, such as lightning flash rate variability with location over the globe<sup>22</sup>, hot-spot regions, and variations with local time and season<sup>19</sup> at geographic scales down to  $0.5^{\circ 23}$  and  $0.1^{\circ 24}$ , respectively. A study of lightning activity on a decadal-length time scale showed that the activity measured by LIS remained roughly constant, while a ground-based detection system found a ~13% reduction in the continental United States<sup>25</sup>. Other studies show that when considered on a global scale, the relationship between column-integrated precipitation ice mass and lightning flash density is almost linear and invariant between land, ocean and coastal regimes in contrast to liquid precipitation<sup>26</sup>.

Furthermore, imaging instruments such as OMI on Earth observation satellites can detect  $NO_X$  (= $NO + NO_2$ ) perturbations by lightning<sup>27,28</sup>, which are produced in high amounts within the hot lightning channel and may impact the ozone budget in the UTLS region.

Ground-based lightning detection systems with global coverage have also been fielded during the past decades. They are based on measurements of pulses of electromagnetic waves in the VLF-LF – range (3–300 kHz) emitted by lightning propagating in the earth-ionosphere wave guide<sup>29–31</sup>. They appear to detect

less lightning activity than OTD/LIS, with a more variable cover of the Earth because of the uneven distribution of wave receivers. They may be more sensitive to low-altitude cloud-to-ground lightning, whereas OTD/LIS is relatively more sensitive to intracloud lightning in clouds at higher altitudes<sup>32</sup>. Lightning Mapping Arrays (LMAs) have also been installed at select locations. They measure lightning processes in three dimensions (also altitude) over regions of several hundred km radii with a time of arrival technique of pulses<sup>33</sup>.

The OTD and LIS were pathfinders for the first-generation lightning detectors on weather satellites in geostationary orbit. One is the Geostationary Lightning Mapper (GLM) on the U.S. GOES-R and GOES-S satellites (GOES-16 and GOES-17), launched in 2016 and 2018, covering the Americas<sup>34</sup>. China has the Lightning Mapper Imager (LMI) on Fengyun-4A and 4B, launched in 2015 and 2021, covering China and East Asia<sup>35</sup>. Finally, Europe has developed the Lightning Imager (LI), launched on the Meteosat Third Generation (MTG-I) satellite on 13. December 2022, to cover Europe and Africa<sup>36</sup>.

### Observations by instruments on the ISS

The LIS instrument was installed on the ISS in 2017 to continue monitoring lightning activity, providing the longest record of activity from LEO with the same instrument  $^{37,38}$ . It allowed for coverage of regions within  $\pm 70^{\circ}$  latitude by OTD,  $\pm$  35° by TRMM-LIS, and  $\pm 51.6^{\circ}$  by ISS-LIS, securing lightning observations that cover more than 25 years, allowing studies of climate–lightning relationships  $^{38}$ . Some recent publications based on ISS-LIS measurements discuss the cross-correlation of ground-based lightning detection systems to estimate the detection efficiencies of measurement techniques such as the LINET 3D system  $^{39}$  and a lightning detection network in China  $^{40}$ . Other papers study the quantification of NO<sub>X</sub> based on lightning and TROPOMI measurements  $^{41}$ , parametrization of NO<sub>X</sub>-production in continuing currents  $^{42}$ , and meteorological conditions for lightning-ignited fires  $^{43}$ .

Recent studies also compared ISS-LIS observations with those of ASIM, GLM, and an LMA system in Colombia<sup>44,45</sup>. ASIM was launched to the ISS 2. April 2018. The payload includes three

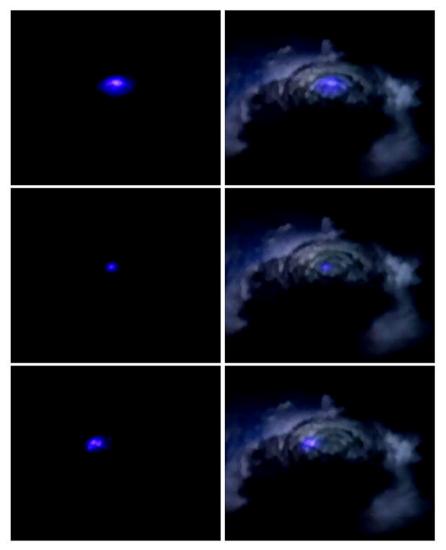


Fig. 2 Blue corona discharges at cloud tops observed by the THOR experiment. On the left, the discharges as observed, and on the right, superimposed on the cloud illuminated by internal lightning<sup>54</sup>.

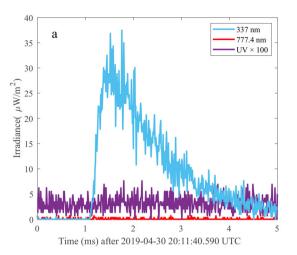
photometers at 180–235 nm (UV), 337 nm with 4 nm bandwidth (blue), and 777.4 nm with 5 nm bandwidth (red), and two cameras observe in the same blue (4 nm) and red bands (3 nm). The photometers sample at 100 kHz and the cameras take 12 frames per second at 1 Mpixels/frame and ~400 m resolution at cloud tops. The instruments point towards the nadir and measure only during the night<sup>46</sup>.

ASIM aims to study Transient Luminous Events (TLEs) in the atmosphere above thunderstorm clouds such as the sprites, jets, and gigantic jets<sup>47,48</sup>, the Terrestrial Gamma-ray Flashes (TGFs)<sup>47</sup> from thunderstorms, and conventional lightning<sup>50</sup>. One of the arguments for ASIM was that TLEs and TGFs, while being fascinating by themselves, also are a window into the inner workings of lightning, where the fast time scales and difficulties with in situ measurements in thunderclouds place severe challenges on the measurements. Thus, although studied since the time of Jacques de Romas and Benjamin Franklin, many aspects are unknown such as how lightning initiates in clouds, how high the electric fields in clouds must be to trigger lightning, and the microphysics of lightning<sup>51</sup>. Therefore, the blue band was selected for ASIM because it is a line of N<sub>2</sub> primarily emitted by streamers. Streamers are ionisation waves known to be launched from the tips of lightning leaders or pointy conductors in highvoltage potentials, where they form a streamer corona<sup>52</sup>, but they can also occur as isolated discharges in clouds<sup>53</sup>. The red band was chosen because it is a strong signature of the lightning leader and is used by LIS and other optical lightning detection imagers. Each spectral band is then sensitive to the two fundamental modes of atmospheric electrical activity: the hot leader (lightning) and the cold streamer (coronas). These modes also form most of the TLEs in the atmosphere above thunderstorms. They are easier to observe here because the temporal and spatial scales are approximately proportional to the mean free path of electrons, which can be orders of magnitude longer in the stratosphere and mesosphere.

The ASIM experiment was preceded by the THOR experiment of Danish astronaut Andreas Mogensen in 2015. With a NIKON D4 colour video camera, he captured profuse activity in the top layer of a thunderstorm cell on the Indian east coast<sup>54</sup>. Some examples are shown in Fig. 2. The observations underscored that the upper regions of thunderstorm clouds are particularly poorly researched because they are difficult to measure from the ground.

The ASIM measurements confirmed that blue streamer discharges are common in the upper regions of deep convective cloud cells<sup>55–57</sup>. Figure 3 shows examples of the two most commonly observed discharges. They consist of a single blue flash





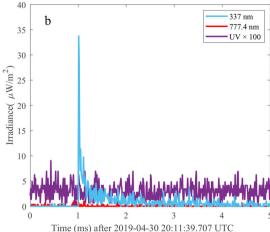
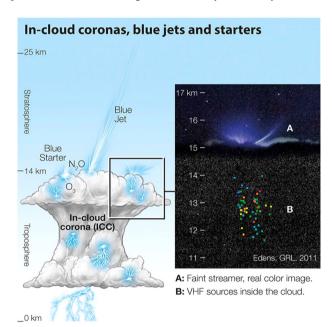


Fig. 3 Blue corona discharges measured by the ASIM photometers. a) With rise time ~50 μs and b) with rise time within one sample of 10 μs.



**Fig. 4 A schematic representation of corona discharges in thunderclouds**<sup>53</sup>**.** The insert is a true colour photo of blue corona emissions, including a blue starter with the altitude of concurrent LMA sources shown as coloured dots increasing in time from blue to red<sup>65</sup>.

with relatively faint emissions in the red band, suggesting the emissions are from corona streamers. In Fig. 3a, the rise time of the blue pulse is several samples, and in Fig. 3b, the pulse peaks within 1 sample (10  $\mu s$ ). Assuming the source of the emissions is short compared to the sampling period, the rise times are interpreted to be from the scattering of photons in cloud hydrometeors and then reflect the depth of the discharges in the clouds  $^{52}$ . The pulse with a fast rise time in Fig. 3b is then from a discharge at the top of a cloud, whereas the one with a longer rise time in Fig. 3a is from a source some kilometres below the cloud top  $^{57}$ .

Blue corona discharges appear to be the optical equivalent to the so-called Narrow Bipolar Events (NBE) in radio signals from thunderstorms. These are short (10–20 µs) bipolar pulses with very high power in the High Frequency (3–30 MHz) and Very High Frequency (3–300 MHz) bands<sup>58</sup>. They are thought to be fast breakdown and may occur at the initiation of lightning flashes<sup>59</sup>. Correlating the ASIM blue discharges with NBEs observed by

ground lightning networks, a consistent picture appears where negative NBEs, which tend to occur at the top of clouds, correlate with fast blue pulses, and positive NBEs, most common within the clouds, with the slower pulses 52.60. Multi-pulse blue discharges are also observed in relation to NBEs where the subsequent pulses could be caused by leader branching 61. In addition, ASIM observed a blue corona discharge at the initiation of lightning and at the onset of a blue jet into the stratosphere 63, suggesting that corona discharges and NBEs may be integral to the initiation of both lightning and upward lightning in the form blue jets.

The discovery of blue corona discharges in thunderclouds follows that of sprites and blue jets (1990, 1994)<sup>47</sup>, TGFs of radiation reaching several tens of MeV (1994)<sup>49</sup>, and gigantic jets from cloud tops to the ionosphere (2002)<sup>48</sup>. Blue emissions at cloud tops with little apparent leader activity were first observed from the ground and reported in 2003<sup>64</sup> and 2011<sup>65</sup>. Observations followed by the Imager for Sprites and Upper Atmospheric Lightning (ISUAL) on the FORMOSAT-2 satellite<sup>66,67</sup>. However, the measurements by ASIM, combined with simultaneous measurements by ground-based radio receivers, proved they are from corona streamers<sup>52</sup>.

### PERSPECTIVES ON LIGHTNING OBSERVATIONS FROM THE ISS

In the following, we first discuss the merits of continued observations of thunderstorms from the ISS in a slanted viewing geometry that allows for altitude resolution of discharges and cloud structure. Our main point is that the change to the atmosphere's radiative properties depends on the altitude of the perturbations <sup>68,69</sup>. It is important, therefore, to measure the altitude of electrical activity and cloud structure to understand the role of thunderstorms in a changing climate. From the ISS, we can get such measurements at high spatial resolution. The discussion is also valid for other LEO space platforms of the future. We end with comments on the synergy of current and planned Earth observation missions and the new lightning mappers on geostationary satellites.

### Perturbations to greenhouse concentrations by corona discharges

Lightning is one of the major sources of nitrogen oxides  $(NO_X = NO + NO_2)$  throughout most of the free troposphere and the most important contributor in the upper troposphere, where the radiative forcing of ozone maximises.  $NO_X$  is the primary precursor for ozone, and the production efficiency in the UTLS region is 2–5 times higher than in the lower troposphere<sup>68</sup>. Whereas the effects of the lightning leader on the atmosphere

have been discussed for several years, corona discharges represent a new pathway of chemical perturbations, which begs the question of their importance relative to lightning leaders. The ASIM studies have so far considered pure corona discharges (blue) and primarily in the upper levels of clouds where photons can escape the cloud in sufficient numbers to be detected by the instruments. Lightning flashes (red) are also commonly observed with intense blue emissions suggesting corona discharges are integral to the lightning processes, as expected (unpublished). Since thundercloud coronas occur in the vicinity of ordinary lightning or in isolation, they may represent a significant source of greenhouse gases. A schematic of corona discharges in a thundercloud is shown in Fig. 4.

In the channels of lightning strokes, neutral gas and electron temperatures reach up to  $\sim$ 40,000 K  $\sim$  3.5 eV<sup>70</sup>. On the other hand, corona discharges in thunderclouds are formed by millions of streamer filaments<sup>60,71,72</sup> that have non-uniform electric fields several times larger in the streamer head than in its body. The streamer head transmits energy to ambient electrons that become hotter (~8 eV) than in the lightning leader channels. The lightning leader is a hot, thermal plasma, whereas corona discharges have cold atoms, molecules and ions, with hot electrons that are not in thermal equilibrium. This leads to different efficiencies in the production of key chemical species. For example, laboratory studies show that ordinary lightning channels produce high amounts of nitric oxide (NO), while their direct ozone (O<sub>3</sub>) and nitrous oxide (N<sub>2</sub>O) production is negligible<sup>73</sup>. In contrast, coronas produce significant amounts of O<sub>3</sub> and N<sub>2</sub>O, but only small amounts of NO<sup>74-78</sup>.

It is a priority to understand the sources and sinks of  $N_2O$  because it is the third strongest greenhouse gas after  $CO_2$  and  $CH_4$  and can deplete ozone<sup>79</sup>. Since the early 1970s, reports have shown sudden, substantial ozone enhancements in or near thunderstorms that are thought to be produced in corona discharges;<sup>80</sup> however, their global contribution is still highly uncertain. In addition to coronas, fertilised soils of agricultural areas are significant sources of  $N_2O^{81}$ , which can be transported to the UTLS region by deep convection<sup>11</sup>. Thus, current research aims to identify and quantify  $N_2O$  sources and sinks and assess their importance.

We suggest that altitude-resolved spectral observations of coronas from the ISS and NBEs from the ground will allow the quantification of corona chemical activity. Spectral observations can also help unravel the underlying non-equilibrium kinetics of cloud corona streamers required for modelling perturbations to

greenhouse gas agents. This will be a first step towards understanding the possible effects of coronas on the atmosphere's radiation budget and the importance of coronas relative to lightning leaders.

#### Altitude-resolved observations

Thunderstorm convection may transport and mix species directly produced by lightning and coronas with urban pollution and smoke from biomass burning to the UTLS region, affecting the radiation budget 11–13,82. For example, the convective transport of SO<sub>2</sub> might trigger the production of new ultrafine particles that changes the aerosol composition and radiation budget in the region 83,84. And thunderclouds can enhance tropospheric O<sub>3</sub> by wrapping and shedding stratospheric air 85. The rapid spread of the anvil outflow in the horizontal direction may affect a large area (~100–1000 km in one direction) and its radiation balance 86. Even a few days after the thunderstorm dissipated, this effect is measurable by airborne in-situ trace gas instrumentation 11.

Clouds and aerosols are the sources of significant uncertainties in climate change predictions. In connection with thunderstorms, some outstanding questions relate, for instance, to the formation of deep convective events and to what extent they affect the moisture distributions within the UTLS region<sup>87–92</sup>. We suggest that these aspects can be addressed in slanted viewing geometry. The photos in Fig. 5 illustrate the potential of such measurements. At night, lightning inside the clouds give a spectacular view of the finer details of their structure and vertical extent. The smaller cloud turret is ~20-30 km across and reaches the tropopause at ~16 km altitude<sup>54</sup>. During the day, the structure can be detected from sunlight reflections. Slanted viewing from the ISS can then measure such phenomena as deep convective events and the formation of cloud turrets extending into the stratosphere, giving a full view of the clouds at a precision that exceeds that of the monitoring satellites in geostationary orbit.

Machine learning techniques applied to the lightning energy and spatial distribution measured by the GLM lead to an altitude within 1.5 km compared to measurements by an LMA<sup>93</sup>. Estimations of cloud top altitudes and location from stereo observations of lightning from GOES-R and -S, lead to ~5 km accuracy on the geographic location, ~240 m in mean altitude compared to the Advanced Baseline Imager (ABI) Cloud Height Algorithm (ACHA), with some kilometres spread in the estimation<sup>94</sup>. As with the stereo observations of GOES, a single instrument on the ISS can give position and altitude determinations comparable to the pixel resolution using the different perspectives of the clouds as the ISS passes over the storm.

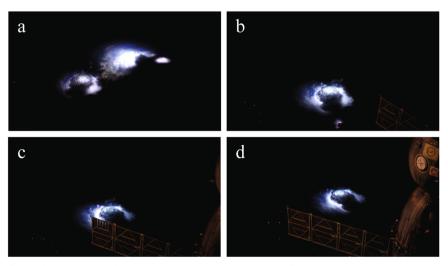
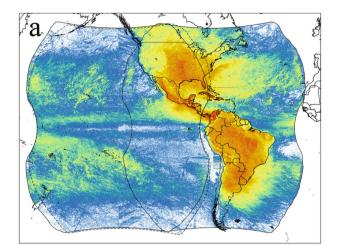
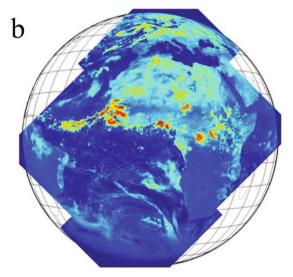
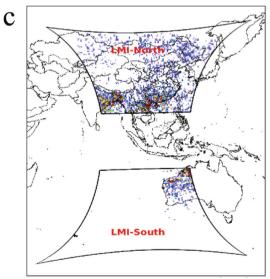


Fig. 5 A time sequence of thunderstorm clouds viewed in a slanted perspective from the ISS as it passes over the storm. The sequence progresses clockwise from (a) to (d) (Credit NASA, ESA; Photo: Andreas Mogensen).





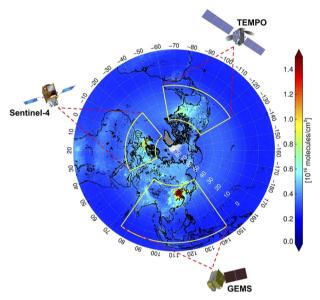




It requires that the storm cloud has recognisable structures observable in several instances during the overflight.

Estimates of chemical perturbations by lightning are based on, for instance, energy deposition in the atmosphere by strokes, observations from aircraft, balloons and the ground, laboratory experiments, and from different lightning parameterisations in

**Fig. 6** The coverage of geostationary lightning imagers. Flashes detected by the GLM on GOES-R and -S during 1 Dec 2018–31 May 2020 with the colour bar from 0,003 to 30 flashes/km²/month<sup>103</sup> © American Meteorological Society. Used with permission (a), coverage of the four identical detectors of LI on MTG covering Africa and Europe<sup>36</sup> © American Meteorological Society. Used with permission (b), and the LMI on FY4A with flashes per 0.5° × 0.5° bin/year from 0 to 1000. The instrument cover either one or the other hemisphere selected by satellite attitude changes<sup>104</sup> (c).



**Fig. 7 The Geostationary Air Quality Constellation covering the most polluted regions in the Northern Hemisphere.** The background image is 10-yr average NO<sub>2</sub> column densities observed by the Ozone Monitoring Instrument (OMI) on board the Aura satellite from 2005 to 2014, showing spatial coverage of GEMS over Asia, TEMPO over North America, and Sentinel-4 over Europe<sup>98</sup> © American Meteorological Society. Used with permission.

chemistry-climate models<sup>95</sup>. Some models of lightning-generated NO<sub>X</sub>, like the NASA Lightning Nitrogen Oxides Model, can estimate the unmixed and otherwise environmentally unmodified vertical profile of lightning NO<sub>X</sub>, provided the altitude of lightning flashes and channel lengths are measured or assumed<sup>96</sup>. Such a model is well suited to ingest altitude-resolved activity. Whereas global climate models are not the best option to simulate thunderstorm impacts because their horizontal resolution leads to averages of regions that are large compared to the typical size of thunderstorms. Mesoscale models are better suited<sup>97</sup>.

# Opportunities with the research infrastructure of the coming years

With the launch of the LI on MTG-I, there will be almost global and continuous coverage of lightning from geostationary orbit with sensors based on the same principles as OTD/LIS. The range of GLM on GOES-R, S, LI on MTG and LMI on FY4A is shown in Fig. 6. Not shown is LMI on FY4B stationed at 105°E.

Many trace gases will also be measured from geostationary orbit with the Air Quality Constellation. It consists of South Korea's Geostationary Environment Monitoring Spectrometer on the GEO-KOMPSAT-2B satellite<sup>98</sup>, NASA's Tropospheric Emissions: Monitoring of Pollution (TEMPO)<sup>99</sup> on Intelsat 40e, expected to be launched in 2022, and ESAs Sentinel-4 instruments in 2023 to be carried by MTG-S<sup>100</sup>. Their coverage is shown in Fig. 7. The measurements will include, for instance, the total column densities of O<sub>3</sub>, NO<sub>2</sub>, SO<sub>2</sub>, HCHO, vertically resolved measurements of O<sub>3</sub>

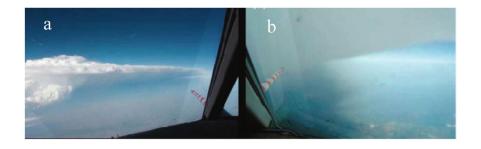


Fig. 8 View from the cockpit of Falcon flight on 29–30 May 2012 over Oklahoma. On 29 May at 23:54 UTC (a) and 8 min later on 30 May at 00:02 UTC inside the lower boundary of the anvil outflow (b)<sup>11</sup>. (Photo R. Welser and A. Minikin, DLR).

and NO<sub>2</sub>, cloud optical thickness, fraction, altitude, aerosol column, optical thickness, type, layer height and absorbing index. Although aimed at detecting pollution in industrial regions of the Northern Hemisphere, the measurements also cover the major thunderstorm regions of the hemisphere and offer an opportunity to measure lightning perturbations to trace gas concentrations.

The upcoming launch of MTG-S with an Infra-red Sounder and an Ultraviolet Visible Near-infra-red spectrometer (Sentinel-4 instrument) will detect several trace species ( $O_3$ ,  $NO_2$ ,  $SO_2$ , HCHO, CO,  $H_2O$ , and aerosols)<sup>100</sup>. The MTG-I and -S offer an unprecedented opportunity to study the processes outlined above. The spatial and temporal resolution is  $4 \times 4 \, \mathrm{km}^2$  and 30 min, which will allow for observations on the scale of a thunderstorm cell. Furthermore, the temporal evolution of  $O_3$  production from lightning  $NO_X$  might even be monitored in long-lived thunderstorms like Mesoscale Convective Systems.

Simultaneous observations by MTG and limb/slant pointing instruments on the ISS of lightning and corona activity, and other measurements with resolution in altitude, offers for the first time an opportunity to simultaneously monitor how thunderstorms and lightning can influence the chemical composition of the UTLS region. In addition to the temporal evolution of the column density of species, vertically resolved MTG-S measurements will be available for some of these species (SO<sub>2</sub>, H<sub>2</sub>O, formaldehyde (HCHO) and aerosols). They can be combined with the altitude-resolved measurements from the ISS, giving essential information for testing chemical models to study the climate effect of thunderstorms and lightning on the climate-sensitive UTLS region.

# Validation of measurements by geostationary lightning and trace species mappers

Planned and current space missions that observe thunderstorm clouds and electrical activity do so with nadir-pointing instruments. For instance, in the case of MTG, a slant-viewing instrument in LEO could be valuable for characterising the detection efficiency, the influence of cloud cover on radiance and geographic location accuracies. The highest flash rate storms, which are of significant interest for nowcasting, tend to have flashes with relatively weak optical emissions <sup>101</sup>. Capturing these weaker lightning discharges may prove essential. A slant-viewing instrument could assist in objectively characterising how MTG LI performs with respect to such flashes.

Airborne, in-situ measurements in the vicinity of thunderstorms must provide ground truth of satellite measurements and models. Needed are measurements of trace species, estimates of specific trace species' vertical column densities, and the vertical distribution of some species, such as O<sub>3</sub> and NO<sub>2</sub>. Flights near deep convection are challenging and rare but have successfully been performed in the past in several studies around the world, as in the field experiment illustrated in Fig. 8. Combining measurements from aircraft with satellite observations of trace gases will help to interpret observations from either platform and allow more accurate comparisons between

thunderstorm electrical activity and convection to the effects on the atmosphere's composition.

### **Outlook and Summary**

The next decade promises to be a renaissance for studies of electric storms. The geostationary spacecraft constellations with instrumentation for detecting lightning, clouds and trace gas perturbations will allow long-duration monitoring of lightning activity and its effect on the atmosphere's trace gas composition. We discuss how instruments in a slanted viewing geometry on spacecraft in low Earth orbit may resolve the altitude with a high spatial resolution that will help to quantify the occurrence and conditions for the generation of coronas and overshooting clouds. Improved models of the chemical perturbations by the electrical activity can be evaluated from satellite and simultaneous aircraft observations. Understanding the relationship between thunderstorm characteristics and atmospheric perturbations gained by instruments on the ISS or other low Earth orbit satellites can be applied to the global characterisation of thunderstorm activity by the geostationary lightning mappers to estimate the global effects. It will allow us to advance in the guest to understand how thunderstorms influence the atmosphere's radiative properties and how climate change affects thunderstorm activity.

### **Reporting summary**

Further information on research design is available in the Nature Portfolio Reporting Summary linked to this article.

Received: 11 May 2022; Accepted: 20 January 2023; Published online: 04 February 2023

### **REFERENCES**

- Mulholland, J. P., Peters, J. M. & Morrison, H. How does LCL height influence deep convective updraft width? *Geophys. Res. Lett.* 48, e2021GL093316 (2021).
- Holton, J. R. et al. Stratosphere-troposphere exchange. Rev. Geophys. 33, 403–439 (1995).
- Gamelin, B. L., Carvalho, L. M. V. & Jones, C. Evaluating the influence of deep convection on tropopause thermodynamics and lower stratospheric water vapor: a RELAMPAGO case study using the WRF model. *Atm. Res.* 267, 105986 (2022).
- Solomon, S. et al. Contributions of stratospheric water vapor to decadal changes in the rate of global warming. Science 327, 1219–1223 (2010).
- Rosenfeld, D., Sherwood, S., Wood, R. & Donner, L. Climate effects of aerosolcloud interactions. Science 343, 379–380 (2014).
- Banerjee, A. et al. Lightning NO<sub>x</sub>, a key chemistry–climate interaction: impacts of future climate change and consequences for tropospheric oxidizing capacity. *Atmos. Chem. Phys.* 14, 9871–9881 (2014).
- Trapp, R. J., Diffenbaugh, N. S. & Gluhovsky, A. Transient response of severe thunderstorm forcing to elevated greenhouse gas concentrations. *Geophys. Res. Lett.* 36, L01703 (2009).
- Schumann, U. & Huntrieser, H. The global lightning-induced nitrogen oxides source. Atmos. Chem. Phys. 7, 3823–3907 (2007).



- Boer, M. M., Resco de Dios, V. & Bradstock, R. A. Unprecedented burn area of Australian mega forest fires. Nat. Clim. Chang. 10, 171–172 (2020).
- Knorr, W., Dentener, F., Lamarque, J.-F., Jiang, L. & Arneth, A. Wildfire air pollution hazard during the 21st century. Atmos. Chem. Phys. 17, 9223–9236 (2017).
- Huntrieser, H. et al. Injection of lightning-produced NO<sub>X</sub>, water vapor, wildfire emissions, and stratospheric air to the UT/LS as observed from DC3 measurements. J. Geophys. Res. Atmos. 121, 6638–6668 (2016).
- Hirsch, E. & Koren, I. Record-breaking aerosol levels explained by smoke injection into the stratosphere. Science 371, 1269–1274 (2021).
- Stocker, M., Ladstädter, F. & Steiner, A. K. Observing the climate impact of large wildfires on stratospheric temperature. Nat. Sci. Rep. 11, 22994 (2021).
- Finney, D. L. et al. A projected decrease in lightning under climate change. Nat. Clim. Change 8, 210–213 (2018).
- Bieniek, P. A. et al. Lightning variability in dynamically downscaled simulations of Alaska's present and future summer climate. J. Appl. Meteorol. Clim. 59, 1139–1152 (2020).
- Romps, D. M., Seeley, J. T., Vollaro, D. & Molinari, J. Projected increase in lightning strikes in the United States due to global warming. Science 346, 851–854 (2014).
- Holzworth, R. H. et al. Lightning in the Arctic. Geophys. Res. Lett. 48, e2020GL091366 (2021).
- 18. Aich, V. et al. Lightning: a new essential climate variable. Eos 99, 07 (2018).
- Christian, H. J. et al. Global frequency and distribution of lightning as observed from space by the Optical Transient Detector. J. Geophys. Res. 108, 4005 (2003).
- Christian, H.J. et al., The lightning imaging sensor, In: Proc. 11th International Conference of Atmospheric Electricity. https://www.academia.edu/22402745/ The\_lightning\_imaging\_sensor (1999).
- Zhang, D., Cummins, K. L., Bitzer, P. & Koshak, W. J. Evaluation of the performance characteristics of the Lightning Imaging Sensor. J. Atm. Ocean. Tech. 36, 1015–1031 (1999).
- Boccippio, D. J., Goodman, S. J. & Heckman, S. Regional differences in tropical lightning distributions. J. Appl. Meteor. Clim. 39, 2231–2248 (2000).
- Cecil, D. J., Buechler, D. E. & Blakeslee, R. J. Gridded lightning climatology from TRMM-LIS and OTD: Dataset description. Atmos. Res. 135, 404–414 (2014).
- Albrecht, R. I., Goodman, S. J., Buechler, D. E., Blakeslee, R. J. & Christian, H. J. Where are the lightning hotspots on Earth? *Bull. Am. Met. Soc.* 97, 2051–2068 (2016)
- 25. Koshak, W. J. et al. Variability of CONUS lightning in 2003–12 and associated impacts. *J. Appl. Meteorol. Clim.* **54**, 15–41 (2015).
- Petersen, W. A., Christian, H. J. & Rutledge, S. A. TRMM observations of the global relationship between ice water content and lightning. *Geophys. Res. Lett.* 32, 114819 (2005)
- Bucsela, E. J., Pickering, K. E., Allen, D. J., Holzworth, R. H. & Krotkov, N. A. Midlatitude lightning NO<sub>X</sub> production efficiency inferred from OMI and WWLLN data. *J. Geophys. Res. Atmos.* 124, 13475–13497 (2019).
- Allen, D. J., Pickering, K. E., Bucsela, E., Krotkov, N. & Holzworth, R. Lightning NO<sub>X</sub> production in the tropics as determined using OMI NO<sub>2</sub> retrievals and WWLLN stroke data. *J. Geophys. Res. Atmos.* 124, 13498–13518 (2019).
- Kaplan, J. O. & Lau, K. H.-K. The WGLC global gridded lightning climatology and time series. Earth Syst. Sci. Data 13, 3219–3237 (2021).
- Rudlosky, S. D., Peterson, M. J. & Kahn, D. T. GLD360 performance relative to TRMM LIS. J. Atm. Ocean. Technol. 34, 1307–1322 (2017).
- 31. Zhu, Y., Stock, M., Lapierre, J. & DiGangi, E. Upgrades of the earth networks total lightning network in 2021. *Remote Sens* 14, 2209 (2022).
- Bürgesser, R. E. Assessment of the World Wide Lightning Location Network (WWLLN) detection efficiency by comparison to the Lightning Imaging Sensor (LIS). Q. J. R. Meteorol. Soc. 143, 2809–2817 (2017).
- 33. Thomas, R. J. et al. Accuracy of the lightning mapping array. *J. Geophys. Res. Atm.* **109**, 1–34 (2004).
- Goodman, S. et al. The GOES-R Geostationary Lightning Mapper (GLM). Atmos. Res. 125–126, 34–49 (2013).
- Yang, J. et al. Introducing the new generation of Chinese geostationary weather satellites—Fengyun 4 (FY-4). Bull. Am. Meteorol. Soc. 98, 1637–1658 (2016).
- Holmlund, K. et al. Meteosat Third Generation (MTG): continuation and innovation of observations from geostationary orbit. *Bull. Am. Meteorol. Soc.* 102, E990–E1015 (2021).
- Blakeslee, R. J. et al. Lightning Imaging Sensor (LIS) for the International Space Station (ISS): mission description and science goals. In *Poc. XV International Conference on Atmospheric Electricity*, 15–20 June 2014, Norman, Oklahoma, U.S.A. https://ntrs.nasa.gov/api/citations/20140011608/downloads/20140011608.pdf (2014)
- Blakeslee, R. J. et al. Three years of the Lightning Imaging Sensor onboard the International Space Station: expanded global coverage and enhanced applications. J. Geophys. Res.: Atm. 125, e2020JD032918 (2020).

- Tiberia, A. et al. A Joint LINET and ISS-LIS view of lightning distribution over the Mt. Cimone area within the GAMMA-FLASH program. *Remote Sens.* 14, 3501 (2022).
- Zhang, Y. et al. Advances in lightning monitoring and location technology research in China. Remote Sens. 14, 1293 (2022).
- Pérez-Invernón, F. J. et al. Quantification of lightning-produced NO<sub>X</sub> over the Pyrenees and the Ebro Valley by using different TROPOMI-NO<sub>2</sub> and cloud research products. Atmos. Meas. Tech. 15, 3329–3351 (2022).
- Pérez-Invernón, F. J., Huntrieser, H., Jöckel, P. & Gordillo-Vázquez, F. J. A parameterization of long-continuing-current (LCC) lightning in the lightning submodel LNO<sub>X</sub> (version 3.0) of the Modular Earth Submodel System (MESSy, version 2.54). Geosci. Model Dev. 15. 1545–1565 (2022).
- Pérez-Invernón, F. J., Huntrieser, H. & Moris, J. V. Meteorological conditions associated with lightning-ignited fires and long-continuing-current lightning in Arizona. New Mexico and Florida. MDPI 5. 96 (2022).
- van der Velde, O. et al. Comparison of high-speed optical observations of a lightning flash from space and the ground. Earth Space Sci. 7, e2020EA001249 (2020).
- Montanyá, J. et al. A simultaneous observation of lightning by ASIM, Colombia-Lightning Mapping Array, GLM, and ISS-LIS. J. Geophys. Res. 126, e2020JD033735 (2021).
- Chanrion, O. et al. The Modular Multispectral Imaging Array (MMIA) of the ASIM payload on the International Space Station. Space Sci. Rev. 215, 28 (2019).
- 47. Neubert, T. On sprites and their exotic kin. Science 300, 747-748 (2003).
- Pasko, V. P., Stanley, M. A., Mathews, J. D., Inan, U. S. & Wood, T. G. Electrical discharge from a thundercloud top to the lower ionosphere. *Nature* 416, 152–154 (2002).
- Fishman, G. J. et al. Discovery of intense gamma-ray flashes of atmospheric origin. Science 164, 1313–1316 (1994).
- Neubert, T. et al. The ASIM mission on the International Space Station. Space Sci. Rev. 215, 26 (2019).
- Dwyer, J. A. & Uman, M. A. The physics of lightning. Phys. Rep. 534, 147–241 (2014).
- Soler, S. et al. Blue optical observations of narrow bipolar events by ASIM confirm streamer activity in thunderstorms. J. Geophys. Res. Atm. 125, e2020JD032708 (2020).
- Gordillo-Vázquez, F. J. & Pérez-Invernón, F. J. A review of the impact of transient luminous events on the atmospheric chemistry: past, present, and future. Atm. Res. 252, 1–48 (2021).
- Chanrion, O. et al. Profuse activity of blue electrical discharges at the tops of thunderstorms. Geophys. Res. Lett. 44. https://doi.org/10.1002/2016GL071311 (2017).
- Dimitriadou, K. et al. Analysis of blue corona discharges at the top of tropical thunderstorm clouds in different phases of convection. *Geophys. Res. Lett.* 49, e2021GL095879 (2022).
- Soler, S. et al. Global frequency and geographical distribution of nighttime streamer corona discharges (BLUEs) in thunderclouds. *Geophys. Res. Lett.* 48, e2021GL094657 (2021).
- Husbjerg, L. S. et al. Characteristics of blue corona emissions observed by ASIM. Geophys. Res. Lett. 49, e2022GL099064 (2022).
- Wu, T. et al. Discharge height of lightning narrow bipolar events. J. Geophys. Res. Atm. 117, D05119 (2012).
- Rison, W. et al. Observations of narrow bipolar events reveal how lightning is initiated in thunderstorms. *Nat. Commun.* 7, 10721 (2016).
- Li, D. et al. Blue flashes as counterparts to narrow bipolar events: the optical signal of shallow in-cloud discharges. J. Geophys. Res. Atm. 126, e2021JD035013 (2021).
- Li, D. et al. Multi-pulse corona discharges in thunderclouds observed in optical and radio bands. Geophys. Res. Lett. e2022GL098938 (2022).
- López, J. A. et al. Initiation of lightning flashes observed from space and the ground: Narrow bipolar events and continuous cloud corona discharges. Atmos. Res. 268, 105981 (2022).
- 63. Neubert, T. et al. Observations of the onset of a blue jet into the stratosphere. *Nature* **589**, 371–375 (2021).
- Lyons, W. A., Nelson, T. E., Armstrong, R. A., Pasko, V. P. & Stanley, M. A. Upward electrical discharges from thunderstorm tops. *Bull. Am. Meteorol. Soc.* 84, 445–454 (2003).
- Edens, H. E. Photographic and lightning mapping observations of a blue starter over a New Mexico thunderstorm. *Geophys. Res. Lett.* 38, L17804 (2011).
- Kuo, C. L., Su, H. T. & Hsu, R. R. The blue luminous events observed by ISUAL payload on board FORMOSAT-2 satellite. *J. Geophys. Res. Space Phys.* 120, 9795–9804 (2015).
- 67. Liu, F. et al. Observations of blue discharges associated with negative narrow bipolar events in active deep convection. *Geophys. Res. Lett.* **45**, 2842–2851 (2018).



- Dahlmann, K., Grewe, V., Ponater, M. & Matthes, S. Quantifying the contributions of individual NO<sub>x</sub> sources to the trend in ozone radiative forcing. *Atmos. Environ.* 45, 2860–2868 (2011).
- Zhang, F., Tie, X. & Bond, D. W. Impacts of anthropogenic and natural NO<sub>x</sub> sources over the U.S. on tropospheric chemistry. *Proc. Nat. Acad. Sci.* 100, 1505–1509 (2003).
- Walker, T. D. & Christian, H. J. Triggered lightning spectroscopy: 2. A quantitative analysis. J. Geophys. Res. Atmos. 124, 3930–3942 (2019).
- Liu, N. et al. Understanding the radio spectrum of thunderstorm narrow bipolar events. J. Geophys. Res. Atm. 124, 10134–10153 (2019).
- Cooray, V., Rubinstein, G. & Rachidi, M. F. Modeling compact intracloud discharge (CID) as a streamer burst. Atmosphere 11, 549 (2020).
- Wang, Y., DeSilva, A. W., Goldenbaum, G. C. & Dickerson, R. R. Nitric oxide production by simulated lightning-dependence on current energy and pressure. *J. Geophys. Res.* 103, 19149–19159 (1998).
- 74. Donohoe, K. G., Shair, F. H. & Wulf, O. R. Production of  $O_3$ , NO and  $N_2O$  in a pulsed discharge at 1 atm. *Ind. Eng. Chem., Fundam.* **16**, 208 (1977).
- Brandvold, D. K., Martinez, P. & Hipsh, R. Field measurements of O<sub>3</sub> and N<sub>2</sub>O produced from corona discharge. Atm. Environ. 30, 973–976 (1996).
- Simek, M. & Clupek, M. Efficiency of ozone production by pulsed positive corona discharge in synthetic air. J. Phys. D: Appl. Phys. 35, 1171 (2002).
- Hill, R. D., Rahmim, I. & Rinker, R. G. Experimental study of the production of nitric oxide, nitrous oxide, and ozone in a simulated atmospheric corona. *Ind. Eng. Chem. Res.* 27, 1264–1269 (1988).
- Zahn, A. et al. Electrical discharge source for tropospheric "ozone-rich transients". J. Geophys. Res. Atm. 107, ACH–16 (2002).
- Portmann, R. W., Daniel, J. S. & Ravishankara, A. R. Stratospheric ozone depletion due to nitrous oxide: influences of other gases. *Philos. Trans. R. Soc. B: Biol. Sci.* 367, 1256–1264 (2012).
- Minschwaner, K. et al. Observation of enhanced ozone in an electrically active storm over Socorro, NM: Implications for ozone production from corona discharges. J. Geophys. Res.: Atm. 113, D17208 (2008).
- 81. Tian, H. et al. A comprehensive quantification of global nitrous oxide sources and sinks. *Nature* **586**, 248–256 (2020).
- and sinks. *Nature* **586**, 248–256 (2020).

  82. Rosenfeld, D. et al. Global observations of aerosol-cloud-precipitation-climate
- Hamed, A. et al. Changes in the production rate of secondary aerosol particles in Central Europe in view of decreasing SO<sub>2</sub> emissions between 1996 and 2006. Atmos. Chem. Phys. 10, 1071–1091 (2010).

interactions. Rev. Geophys. 52, 750-808 (2014).

- Fan, J. et al. Substantial convection and precipitation enhancements by ultrafine aerosol particles. Science 359, 411–418 (2018).
- 85. Pan, L. L. et al. Thunderstorms enhance tropospheric ozone by wrapping and shedding stratospheric air. *Geophys. Res. Lett.* **41**, 7785–7790 (2014).
- Gasparini, B., Blossey, P. N., Hartmann, D. L., Lin, G. & Fan, J. What drives the life cycle of tropical anvil clouds? *J. Adv. Model Earth Syst.* 11, 2586–2605 (2019).
- Price, C. Evidence for a link between global lightning activity and upper tropospheric water vapour. *Nature* 406, 290–293 (2000).
- Dauhut, T., Chaboureau, J., Haynes, P. H. & Lane, T. P. The mechanisms leading to a stratospheric hydration by overshooting convection. *J. Atmos. Sci.* 75, 4383–4398 (2018).
- 89. O'Neill, M. E., Orf, L., Heymsfield, G. M. & Halbert, K. Hydraulic jump dynamics above supercell thunderstorms. *Science* **373**, 1248–1251 (2021).
- Smith, J. B. Convective hydration of the stratosphere. Science 373, 1194–1195 (2021).
- Corti, T. et al. Unprecedented evidence for deep convection hydrating the tropical stratosphere. Geophys. Res. Lett. 35, L10810 (2008).
- 92. Homeyer, C. R. et al. Convective transport of water vapor into the lower stratosphere observed during double tropopause events. *J. Geophys. Res. Atmos.* **119**, 10941–10958 (2014).
- Peterson, M., Light, T. E. L. & Mach, D. The illumination of thunderclouds by lightning:
   Retrieving optical source altitude. Earth Space Sci. 9, e2021EA001944 (2022).
- Mach, D. & Virts, K. A technique for determining three-dimensional storm cloudtop locations using stereo optical lightning pulses observed from orbit. *J. Atm. Ocean. Technol.* 38, 1993–2001 (2021).
- Gordillo-Vázquez, F. J., Pérez-Invernón, F. J., Huntrieser, H. & Smith, A. K. Comparison of six lightning parameterizations in CAM5 and the impact on global atmospheric chemistry. *Earth Space Sci.* 6, 2317–2346 (2019).
- Koshak, W. J., Peterson, H. S., Biazar, A. P., Khan, M. N. & Wang, L. The NASA Lightning Nitrogen Oxides Model (LNOM): application to air quality modeling. *Atmos. Res.* 135–136, 363–369 (2014).

- Evans, C., Weiss, S. J., Jirak, I. L., Dean, A. R. & Nevius, D. S. An evaluation of paired regional/convection-allowing forecast vertical thermodynamic profiles in warmseason, thunderstorm-supporting environments. Weather Forecast. 33, 1547–1566 (2018).
- Kim, J. et al. New era of air quality monitoring from space: geostationary environment monitoring spectrometer (GEMS). Bull. Am. Meteorol. Soc. 101, 1–22 (2020).
- Zoogman, P. et al. Tropospheric emissions: Monitoring of pollution (TEMPO). J. Quant. Spectrosc. Radiat. Transf. 186, 17–39 (2017).
- Riedl, S. et al. Sentinel 4 UVN: a geostationary imaging UVN spectrometer for airquality monitoring: performance, measurement modes and model philosophy. Proc. SPIE11180 International Conference on Space Optics — ICSO 2018, 1118005 (2019).
- Rutledge, S. A., Hilburn, K. A., Clayton, A., Fuchs, B. & Miller, S. D. Evaluating Geostationary Lightning Mapper flash rates within intense convective storms. *J. Geophys. Res. Atm.* 125, e2020JD032827 (2020).
- 102. NASA Earth Observatory, Image of the Day for June 23. https://earthobservatory.nasa.gov/images/6679/patterns-of-lightning-activity (2006).
- Rudlosky, S. & Virt, K.S. Dual Geostationary Lightning Mapper observations. Monthly Weather Rev. 149, 979–998, https://doi.org/10.1175/MWR-D-20-0242.1.
- 104. Cao, D., Lu, F., Zhang, X. & Yang, J. Lightning Activity Observed by the FengYun-4A Lightning Mapping Imager. *Remote Sens* **13**, 3013 (2021).

### **ACKNOWLEDGEMENTS**

ASIM is a mission of the European Space Agency (ESA) and is funded by ESA and by national grants of Denmark, Norway and Spain. THOR and ASIM are supported by DTU Space. F.J.G.V. acknowledges support from the Spanish Ministry of Science and Innovation under project PID2019-109269RB-C43 and the FEDER programme.

#### **AUTHOR CONTRIBUTIONS**

T.N. wrote the Introduction and the Observation sections and provided Figs. 1–3, 5 and 7. F.J.G.V. wrote the subsection on chemistry in the Perspective section and provided Fig. 4. H.H. wrote the subsection on LEO satellite and aircraft observations in the Perspective section and provided Figs. 6 and 8. All authors provided additional inputs and corrections to all sections.

### **COMPETING INTERESTS**

The authors declare no competing interests.

### **ADDITIONAL INFORMATION**

**Supplementary information** The online version contains supplementary material available at https://doi.org/10.1038/s41526-023-00257-4.

Correspondence and requests for materials should be addressed to Torsten Neubert.

**Reprints and permission information** is available at http://www.nature.com/reprints

**Publisher's note** Springer Nature remains neutral with regard to jurisdictional claims in published maps and institutional affiliations.



**Open Access** This article is licensed under a Creative Commons Attribution 4.0 International License, which permits use, sharing,

adaptation, distribution and reproduction in any medium or format, as long as you give appropriate credit to the original author(s) and the source, provide a link to the Creative Commons license, and indicate if changes were made. The images or other third party material in this article are included in the article's Creative Commons license, unless indicated otherwise in a credit line to the material. If material is not included in the article's Creative Commons license and your intended use is not permitted by statutory regulation or exceeds the permitted use, you will need to obtain permission directly from the copyright holder. To view a copy of this license, visit http://creativecommons.org/licenses/by/4.0/.

© The Author(s) 2023