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### Special Section:

Carbon Weather: Toward the next generation of regional greenhouse gas inversion systems

### Key Points:

- This study uses ethane observations to quantify both ethane and methane emissions from the United States oil and gas (O&G) sector
- Ethane emissions in the southcentral and eastern United States are larger than the NEI 2017 inventory by more than a factor of 2
- Ethane results indicate that the US EPA methane inventory is underestimating leak rates from the O&G sector

### Supporting Information:

Supporting Information may be found in the online version of this article.

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## Analysis of Oil and Gas Ethane and Methane Emissions in the Southcentral and Eastern United States Using Four Seasons of Continuous Aircraft Ethane Measurements

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**Abstract** In the last decade, much work has been done to better understand methane (CH<sub>4</sub>) emissions from the oil and gas (O&G) industry in the United States. Ethane (C<sub>2</sub>H<sub>6</sub>), a gas that is co-emitted with thermogenic sources of CH<sub>4</sub>, is emitted in the US predominantly by the O&G sector. In this study, we perform an inverse analysis on 200 h of atmospheric boundary layer C<sub>2</sub>H<sub>6</sub> measurements to estimate C<sub>2</sub>H<sub>6</sub> emissions from the US O&G sector. Measurements were collected from 2017 to 2019 as part of the Atmospheric Carbon and Transport (ACT) America aircraft campaign and encompass much of the central and eastern United States. We find that for the fall, winter, and spring campaigns, C<sub>2</sub>H<sub>6</sub> data consistently exceeds values that would be expected based on EPA O&G leak rate estimates by more than 50%. C<sub>2</sub>H<sub>6</sub> observations from the summer 2019 data set show significantly lower C<sub>2</sub>H<sub>6</sub> enhancements in the southcentral region that cannot be reconciled with data from the other three seasons, either due to complex meteorological conditions or a temporal shift in the emissions. Combining the fall, winter, and spring C<sub>2</sub>H<sub>6</sub> posterior emissions estimate to an inventory of O&G CH<sub>4</sub> emissions, we estimate that O&G CH<sub>4</sub> emissions are larger than EPA inventory values by 48%–76%. Uncertainties in the gas composition data limit the accuracy of using C<sub>2</sub>H<sub>6</sub> as a proxy for O&G CH<sub>4</sub> emissions. These limits could be resolved retroactively by increasing the availability of industry-collected gas composition data.

**Plain Language Summary** Methane is a potent greenhouse gas responsible for a quarter of the warming the climate has experienced thus far. The oil and gas (O&G) sector is a significant source of methane through leaks in its infrastructure. Recent studies of individual basins have found emissions from O&G in the US to be greater than inventory estimates, but difficulties arise with source attribution in broader scale studies due to the numerous potential sources of methane. This study quantifies methane emissions from O&G by looking at atmospheric ethane, a gas whose emissions originate mostly from O&G in the US. Hundreds of hours of boundary layer ethane observations were collected via aircraft over the course of four seasons between 2017 and 2019. These observations are compared with model-projected ethane values based on our current knowledge of ethane emissions, and those emissions are adjusted to best match the observed data. We find ethane emissions are grossly underestimated in the US. Because ethane is co-emitted with O&G methane sources, this underestimation of ethane reflects similar underestimations in O&G methane emissions. We conclude that US inventories are underestimating methane emissions from O&G by 48%–76%.

## 1. Introduction

Methane (CH<sub>4</sub>) is an important greenhouse gas with 28–35 times the warming potential of carbon dioxide over a 100 year period (Myhre et al., 2013). Global CH<sub>4</sub> concentrations in the atmosphere have nearly tripled since pre-industrial times, mainly driven by anthropogenic activity, and are responsible for a fourth of the increased radiative forcing on the planet (Myhre et al., 2013). Although CH<sub>4</sub> concentrations stabilized for a

brief period in the early 2000s, global concentrations began increasing again by 2007 and have continued to increase through the present date (Dlugokencky, 2021; Nisbet et al., 2019).

Efforts to understand causes for increasing global CH<sub>4</sub> trends are hampered by difficulties related to source attribution which arise from complex emission processes, limited data availability, and disagreements between bottom-up inventories and top-down assessments (Saunio et al., 2020). CH<sub>4</sub> has numerous anthropogenic sources, including animal agriculture, fossil fuel extraction, and waste management. In addition to anthropogenic emitters, there are natural sources of CH<sub>4</sub> emissions that play a large role in the global CH<sub>4</sub> budget. Of particular importance to the global CH<sub>4</sub> budget are emissions from anaerobic respiration in wetlands, which create unique challenges to CH<sub>4</sub> source attribution. Unlike most anthropogenic sources, whose emissions can be spatially mapped out using inventory techniques and are relatively consistent in magnitude on an annual timeframe, wetland emissions are more difficult to describe, with an uncertain spatial pattern and large seasonal variability in emissions based on soil temperature and moisture (Yvon-Durocher et al., 2014). Many prior studies have examined seasonal emissions and magnitudes of local wetland sources using flux tower measurements (Grant & Roulet, 2002; Matthes et al., 2014), but extrapolating these results to a continental or global scale is challenging.

Difficulties with source attribution extend into the US, where both anthropogenic and natural sources play a large role in the country's CH<sub>4</sub> budget. In the US, greenhouse gas emissions from anthropogenic sources are quantified and aggregated on a national scale through a bottom-up inventory created by the United States Environmental Protection Agency (EPA). The largest sources of anthropogenic CH<sub>4</sub> emissions projected by this inventory are leaks in oil and gas (O&G) infrastructure, enteric fermentation and manure management related to livestock, and anaerobic respiration occurring in landfills (US Environmental Protection Agency, 2020b). These sources were responsible for 83% of US anthropogenic CH<sub>4</sub> emissions in the 2018 inventory estimate. Natural sources in the US are dominated by wetlands and are not tracked by the EPA's inventory. WetCHARTs, a global wetland emissions ensemble, is generally used as the prior for CH<sub>4</sub> emissions from wetlands in the US (Bloom et al., 2017; Maasakkers et al., 2016; Sheng et al., 2017). WetCHARTs ensemble members estimate the country's wetland emissions to vary from values that are insignificant to totals rivaling those from anthropogenic emissions, with seasonal dependence and spatial variability between ensemble members, resulting in large uncertainties in the overall US CH<sub>4</sub> budget.

Contrary to recent global increases in the atmospheric growth rate of CH<sub>4</sub>, the EPA's inventory estimates of anthropogenic CH<sub>4</sub> emission in the US show a 10% decrease in the last decade, from 28 Tg in 2008 to 25 Tg in 2018. This decrease is driven primarily by projected decreases in emissions from the energy sector, despite a 50% increase in gas production and a greater than 200% increase in oil production during the 10 year period (US Energy Information Administration, 2020b). Over the last several years, various atmospheric studies monitoring emissions from O&G from individual wellpads (Caulton et al., 2019; Rella et al., 2015; Robertson et al., 2017), basins (Barkley et al., 2017; Karion et al., 2013, 2015; Peischl et al., 2015, 2016, 2018; Pétron et al., 2012, 2014), and entire regions (Barkley, Davis, et al., 2019) have consistently found emission rates larger than the EPA inventory, raising concerns of a broad underestimation of leaks from the O&G sector (Alvarez et al., 2018). However, large-scale CH<sub>4</sub> inversion studies involving the US have not been as conclusive, with differing findings as to the accuracy of inventory emissions from O&G, animal agriculture, and wetlands (Maasakkers et al., 2019; Miller et al., 2013; Sheng et al., 2018; Yu et al., 2020). The large spread of uncertainty regarding the magnitude, spatial distribution, and seasonality of CH<sub>4</sub> emissions from wetlands, as well as the numerous other potential sources of CH<sub>4</sub>, poses a challenge to large scale studies attempting source attribution of detected CH<sub>4</sub> signals.

One common method to disaggregate certain sources is to measure both CH<sub>4</sub> and ethane (C<sub>2</sub>H<sub>6</sub>) mixing ratios (Barkley, Lauvaux et al., 2019; McKain et al., 2015; Pétron et al., 2020). Sources of C<sub>2</sub>H<sub>6</sub> coincide with thermogenic CH<sub>4</sub> emitters (O&G and C<sub>2</sub>H<sub>6</sub> supply chain, biomass burning), but not biogenic sources (wetlands, animal agriculture, and landfills) (Yacovitch et al., 2014). Thus, a correlation (or lack of) between C<sub>2</sub>H<sub>6</sub> and CH<sub>4</sub> can provide information on the sector responsible for observed CH<sub>4</sub> enhancements. While global C<sub>2</sub>H<sub>6</sub> emissions are split between the energy sector and biomass burning (Xiao et al., 2008), the National Emissions Inventory (NEI) 2017 C<sub>2</sub>H<sub>6</sub> inventory attributes more than two-thirds of C<sub>2</sub>H<sub>6</sub> emissions in the US to the O&G sector, making it an ideal tracer for identifying O&G emissions in the region (US Environmental Protection Agency, 2020a). This technique has been used in various facility and basin-scale

studies across individual O&G basins to verify that the measured enhancements are associated with O&G activity (Peischl et al., 2018; Roscioli et al., 2015). Similarly, C<sub>2</sub>H<sub>6</sub> concentrations from a network of flask measurements have been used to infer trends in emissions from O&G emissions on broader scales (Dalsøren et al., 2018; Helmig et al., 2016; Lan et al., 2019; Tzompa-Sosa et al., 2017; Xiao et al., 2008).

In addition to its role in attributing CH<sub>4</sub> emissions to different source types, understanding C<sub>2</sub>H<sub>6</sub> emissions is important in atmospheric air quality studies. Oxidation of C<sub>2</sub>H<sub>6</sub> is involved in the formation of ozone and other photochemical constituents (peroxyacetyl nitrate, acetaldehyde, formaldehyde, and carbon monoxide) giving rise to poor air quality (Aikin et al., 1982; Hodnebrog et al., 2018). Furthermore, since the decay of C<sub>2</sub>H<sub>6</sub> in the troposphere occurs primarily via reactions with the hydroxyl (OH) radical, significant local enhancements of C<sub>2</sub>H<sub>6</sub> from O&G operations can affect the local concentrations and lifetimes of other organic gases. Additionally, other volatile organic compounds are co-emitted with CH<sub>4</sub> and C<sub>2</sub>H<sub>6</sub> in O&G basins, contributing to the formation of surface ozone (Edwards et al., 2014; Gilman et al., 2013). Therefore, C<sub>2</sub>H<sub>6</sub> in its production of ozone acts as a direct greenhouse gas, and at the same time, its effects on CH<sub>4</sub> and other organics act as an indirect greenhouse gas (Kort et al., 2016).

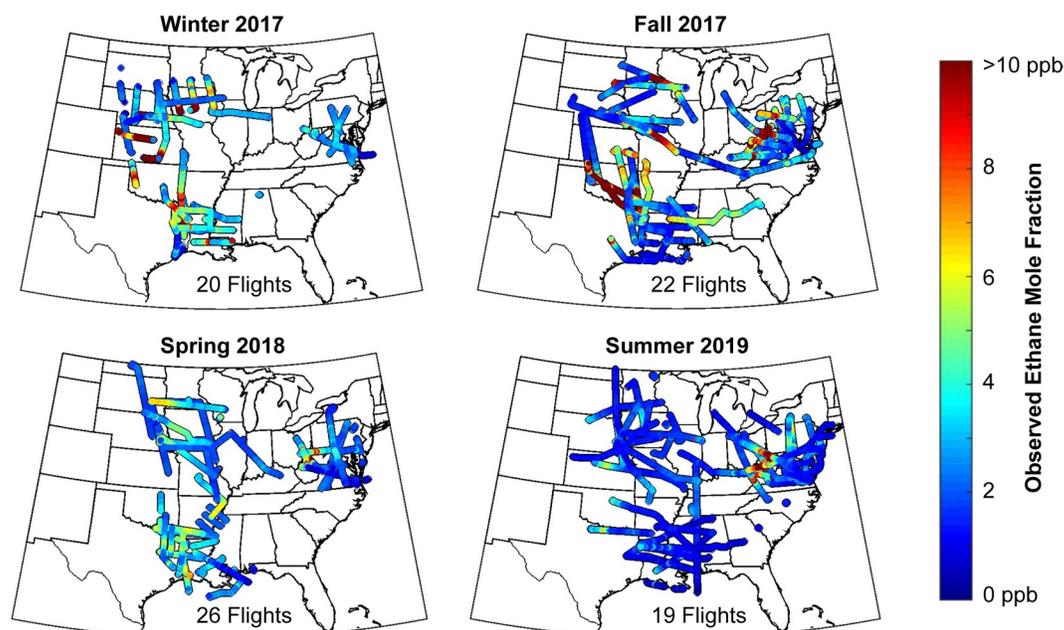
Advancements in technology have resulted in the increased availability of instrumentation capable of measuring C<sub>2</sub>H<sub>6</sub> mixing ratios precisely at high-temporal resolution (Kostinek et al., 2019; Weibring et al., 2020; Yacovitch et al., 2014), expanding our capabilities of studying C<sub>2</sub>H<sub>6</sub> emissions. The Atmospheric Carbon and Transport-America (ACT-America) mission was a 5-season aircraft campaign across the central and eastern US designed to examine various trace gases and their transport in the atmosphere. During four of the seasonal campaigns, continuous C<sub>2</sub>H<sub>6</sub> measurements were collected, producing more than 200 h of boundary layer C<sub>2</sub>H<sub>6</sub> data capable of capturing various plume structures related to O&G activity on a regional scale. In this study, we examine the characteristics of this unprecedented C<sub>2</sub>H<sub>6</sub> data set. From this analysis, we (1) provide up-to-date maps of C<sub>2</sub>H<sub>6</sub> distributions in the boundary layer from 2017 to 2019 spanning four seasons over the eastern and southcentral US; (2) quantify C<sub>2</sub>H<sub>6</sub> emissions from O&G sources in these regions using atmospheric inversion techniques and compare these values to existing bottom-up inventories; and (3) utilize these updated C<sub>2</sub>H<sub>6</sub> inventories in determining CH<sub>4</sub> emissions from the O&G sector.

## 2. Methods

### 2.1. ACT-America Aircraft Campaign and Observational Data Set

The ACT-America campaign was a NASA-funded Earth Venture suborbital aircraft mission designed to study the transport of trace gases throughout the US, with flight activities spanning summer 2016 to summer 2019 (Davis et al., 2021). During this time, five individual campaigns covering all four seasons (summer twice) were conducted using two aircraft, a NASA Langley Research Center Beechcraft B200 King Air and a NASA Wallops Flight Facility Lockheed C130 Hercules, collecting meteorological and trace gas data within fair weather and frontal weather patterns across the central and eastern US. Flights took place during afternoon hours and were typically composed of long horizontal transects (>500 km) at multiple heights, with designated boundary layers and free troposphere legs. Altogether, a total of 1,140.7 h of aircraft observations were collected across 121 flight days. For this study, boundary layer observations are primarily used in the analysis, identified using flags from the ACT-America data set that determine boundary layer height based on potential temperature profiles from the flights, and constitute one-third of the entire ACT-America data set (Pal et al., 2020). Data from the ACT-America campaign used in this study can be found at [https://daac.ornl.gov/cgi-bin/dataset\\_lister.pl?p=37](https://daac.ornl.gov/cgi-bin/dataset_lister.pl?p=37).

For all five seasonal campaigns, high frequency CH<sub>4</sub> dry air mole fraction data were collected on both aircraft using a commercial PICARRO G2401-m instrument adapted with a custom inlet system for drying and conditioning the sample air (DiGangi et al., 2018). During winter 2017, fall 2017, spring 2018, and summer 2019 campaigns, high frequency C<sub>2</sub>H<sub>6</sub> mole fraction data were collected on the B200 aircraft using the CAMS-2 (Compact Airborne Multi-Species Spectrometer) instrument (Weibring et al., 2020). The CAMS-2 C<sub>2</sub>H<sub>6</sub> measurements when averaged over time and linearly regressed versus NOAA portable flask packages collected during the flights yielded slopes in the 0.990 to 1.031 range across seasonal campaigns, with *r*<sup>2</sup> values between the two measurements of 0.996 (Baier et al., 2020; Weibring et al., 2020), providing high confidence in the accuracy of its measurements. Additionally, during the fall 2017 and summer 2019 campaigns,



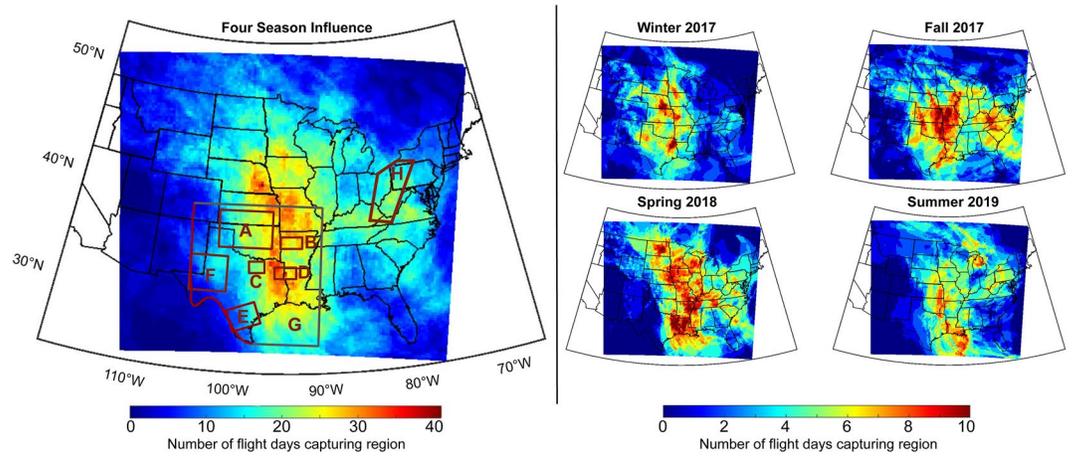
**Figure 1.** Continuous boundary layer  $C_2H_6$  measurements were collected by the B200 and C130 aircraft during each seasonal campaign. The total number of flight days are shown at the bottom of each seasonal figure.

the C130 aircraft was equipped with a quantum and interband cascade laser spectrometer (QCLS) capable of continuous in situ  $C_2H_6$  measurements (Kostinek et al., 2019). The QCLS instrument performed in-flight two-point calibrations every three to 10 minutes to ensure accurate measurements throughout the flights and, similar to the CAMS-2 instrumentation, measurements were compared to NOAA portable flask samples taken in flight (Kostinek et al., 2019). Altogether, the entire  $C_2H_6$  data set contains more than 500 h of continuous airborne  $C_2H_6$  observations, of which more than 200 h were within the atmospheric boundary layer (ABL), making it an ideal data set to study atmospheric  $C_2H_6$  plumes and structures throughout the central and eastern US.

From the observational data set, large  $C_2H_6$  plumes were consistently observed in the boundary layer downwind of Texas/Oklahoma/Louisiana in the southcentral US and along the western Appalachians in the northeast (Figure 1).  $C_2H_6$  plumes could be observed in the midwestern flights as far north as Minnesota when consistent southerly winds were present.  $C_2H_6$  mixing ratios within the plumes were largest during the fall and winter campaign. During the summer 2019 campaign,  $C_2H_6$  plumes greater than 7 ppb were detected in the western Appalachian similar to other seasons, but signals in the southcentral US were substantially smaller (<3 ppb). Additionally, boundary layer  $C_2H_6$  mixing ratios in the summer were often similar in value to their free tropospheric counterparts, a trait not observed during other seasons (Figure S1).

Influence functions for observations within the boundary layer were created using the Lagrangian particle dispersion model FLEXPART-WRF (Brioude et al., 2013) to provide information on the area captured by the ACT-America  $C_2H_6$  data set. The model was run at a 27 km resolution over the North American domain and was meteorologically driven using WRF-Chem simulations developed as part of the ACT-America campaign (Feng et al., 2019). Every 30 s of flight time, 5,000 particles were released from the aircraft location and traced back 10 days in time. Further details on the development of the influence functions can be found in Cui et al. (2015).

For each season, we sum up the number of flight days where a surface grid cell is in the top 80% of the boundary layer influence function for a given day to examine the amount of coverage of different O&G plays captured during the seasonal campaigns (Figure 2). Much of the midwestern and southcentral US is captured across all four seasons, encompassing many of the major O&G basins including Haynesville, Fayetteville, Anadarko, Woodford, and the Barnett. The Eagle Ford and the Permian plays, located on the extreme southern and western portions of the southcentral US, have less uniform seasonal coverage, with



**Figure 2.** (Left) Map showing the frequency that regions are captured by flights from the four season ACT-America data set used in this study. A flight day is considered to have influence from a grid if the grid is within the top 80% of the boundary layer influence function for that day. O&G basins and plays pertinent to this study are highlighted in boxes and are labeled as follows: (A) Anadarko. (B) Fayetteville. (C) Barnett. (D) Haynesville. (E) Eagle Ford. (F) Permian. (G) Southcentral U.S. (H) Appalachia. Basins are defined using an approximation of the location of their O&G emissions from the EPA 2012 Gridded Methane Inventory. (Right) Same as left, but broken down into seasonal campaigns.

few flights capturing the regions in the spring and summer campaign, but better coverage during the fall and winter campaign when winds are climatologically more westerly. In the northeast, many flight tracks were designed to fly downwind of the Marcellus and Utica O&G basins in the western Appalachian regions of West Virginia and western Pennsylvania, and thus this area has significant coverage in the influence maps for all four seasons.

## 2.2. Inversion Framework

Boundary layer  $C_2H_6$  mixing ratio data collected from the ACT-America campaign is used in this study to perform an inversion and learn about  $C_2H_6$  emissions from O&G sources in the southcentral and eastern US. The basic inversion framework is similar to previous work (Lauvaux et al., 2012; Sheng et al., 2018) and is expressed by cost function,

$$J(\mathbf{x}) = \frac{1}{2}(\mathbf{y} - \mathbf{H}\mathbf{x})^T \mathbf{R}^{-1}(\mathbf{y} - \mathbf{H}\mathbf{x}) + \frac{1}{2}(\mathbf{x} - \mathbf{x}_0)^T \mathbf{B}^{-1}(\mathbf{x} - \mathbf{x}_0) \quad (1)$$

In this equation, we solve for a posterior emissions grid  $\mathbf{x}$  that minimizes the cost function  $J$  using influence functions ( $\mathbf{H}$ ) that translate the flux field to a modelled enhancement ( $\mathbf{H}\mathbf{x}$ ). In the cost function, two terms control the solution. The first term is a cost related to the mismatch between the posterior modelled enhancements versus observed mixing ratio enhancement ( $\mathbf{y}$ ), with greater discrepancies resulting in a larger cost term. Here,  $\mathbf{R}$  is the observation error covariance matrix and weights the first term based on the confidence in the observations and model transport. The second term in the cost function equation is a cost related to the change between the posterior flux ( $\mathbf{x}$ ) and the prior flux ( $\mathbf{x}_0$ ), with larger changes resulting in a greater cost. Here,  $\mathbf{B}$  is the flux error covariance matrix and weights the second term based on the confidence in the prior flux field. Minimizing the cost function with respect to  $\mathbf{x}$  yields

$$\mathbf{x} = \mathbf{x}_0 + \mathbf{B}\mathbf{H}^T(\mathbf{H}\mathbf{B}\mathbf{H}^T + \mathbf{R})^{-1}(\mathbf{y} - \mathbf{H}\mathbf{x}_0) \quad (2)$$

and solving for  $\mathbf{x}$  yields the posterior flux field (Brasseur & Jacob, 2017), assuming Gaussian errors. For this study, observations from each season are grouped together and a posterior flux map is solved for each of the seasonal flight campaigns.

A best guess  $C_2H_6$  emissions map was created to serve as a prior for the inversion. To create this prior,  $CH_4$  emissions from the O&G sector of the EPA Gridded 2012  $CH_4$  Emissions Inventory (Maasakkers

et al., 2016) were multiplied by expected molar  $C_2H_6/CH_4$  ratios of each basin (Table S2), resulting in an  $C_2H_6$  emissions map. For Eagle Ford, Haynesville, Fayetteville, Barnett, Permian, Denver-Julesburg, and the Bakken O&G basins, atmospheric measurements from NOAA aircraft studies were available to derive observed basin-wide  $C_2H_6/CH_4$  ratios (Peischl et al., 2015, 2018; Tzompa-Sosa et al., 2017). For the Anadarko and Appalachian region where representative atmospheric measurements were not available, data from the United States Geological Survey were used to create a spatial map of  $C_2H_6/CH_4$  ratios to apply to these regions (Kitanidis, 1997; US Geological Survey, 2018) (See Section S1 and Figure S1 for additional info). For all emissions related to transmission, storage, and distribution, an  $C_2H_6/CH_4$  ratio of 0.027 was applied (Plant et al., 2019). For  $C_2H_6$  emissions unrelated to O&G operations, as well as sources in Canada and Mexico, values from the US National Emissions Inventory 2017 (NEI2017) for  $C_2H_6$  are used. These other sources when multiplied by the influence functions from the flight campaign represent 31% of the signal in the prior inventory, and are generally uncorrelated spatially with O&G sources. Overall, the described  $C_2H_6$  emissions map represents our best guess as to representing  $C_2H_6$  emissions based on the EPA's 2012 bottom-up O&G  $CH_4$  emissions inventory and will be referred to as the "Default" map henceforth. Although O&G emissions from the Default map are based on EPA estimates from 2012, nationwide bottom-up O&G  $CH_4$  emission estimates from the EPA have decreased by only 4% between 2012 and 2018 (US Environmental Protection Agency, 2020b). For this reason, we do not expect the 2012 EPA Gridded 2012  $CH_4$  Emissions Inventory to be unrepresentative of the EPA estimates of total O&G emissions during the years of the flight campaign (2017–2019), though individual basins may be more susceptible to higher uncertainty.

Two additional  $C_2H_6$  emission maps were created to test the sensitivity of the inversion to different priors. The first alternative map was created by taking the EPA  $CH_4$  emissions map used in the creation of the Default inventory and applying a flat  $C_2H_6/CH_4$  ratio of 0.085. This ratio preserves the total  $C_2H_6$  emissions from the "Default" map but redistributes them in a way that removes the knowledge of the unique gas compositions of different basins, and is referred to henceforth as the "Flat Ratio" inventory. The second alternative map is based on O&G  $C_2H_6$  emissions provided by the NEI2017 inventory. In addition to having a slightly different spatial distribution compared to the Default inventory, total emissions from the O&G sector in the NEI2017 inventory are roughly half of the comparable emissions from the Default inventory.  $C_2H_6$  emissions from the Emissions Database for Global Atmospheric Research (EDGAR) v4.3.2 were originally considered as well for this analysis, but total emissions were 8 times lower than our Default inventory and were decided to be too inaccurate to serve as a useful prior (Huang et al., 2017). For each of the three inventories used, the prior is regridded to a  $27\text{ km}^2$  field to match the mapping of the influence functions and are solved for at that resolution. All fluxes in both the prior and posterior are treated as a constant value for the season they are being evaluated. All three priors and their respective posterior solutions can be seen in Figure 4.

Approximately 200 h of continuous boundary layer  $C_2H_6$  measurements from the ACT-America campaign were available and used as observational input for this study. For this study, we restrict the domain of our inversion to the area within the four corners (23.7°N, 110.7°W), (23.0°N, 77.5°W), (49.9°N, 67.3°W), and (51.1°N, 119.0°W), (domain shown as the colored region in Figure 2). Because the influence functions only provide a local enhancement inside the study domain, for each flight date a background value is determined to represent the  $C_2H_6$  mixing ratios entering the domain. This value is chosen by taking the fifth percentile of the observed boundary layer  $C_2H_6$  mixing ratios on a given flight and subtracting it from the observations, producing an observed  $C_2H_6$  enhancement. The fifth percentile of model-projected enhancements along the flight track is then added onto the observed enhancements in order to align the modeled and observed background values. This final step is necessary for rare scenarios where modeled O&G enhancements are influencing the entire aircraft transect, thus impacting the observed background mixing ratios (see Barkley, Davis, et al., 2019 for further details). In calculating the modeled  $C_2H_6$  enhancements for this study, we treat  $C_2H_6$  as an inert gas rather than a reactive one due to its long average lifetime (weeks to months) relative to the length of time the local plumes travel from the source to the aircraft (a few hours to 1–2 days). For more information on the possible influence of  $C_2H_6$  loss rates, see Section S2.

To run an inversion, values must be assigned to the **R** and **B** matrices related to the uncertainty in the observation/transport and prior flux fields respectively. For the observational/transport uncertainty matrix **R**, we use a method similar to the residual error method discussed in Sheng et al. (2017). For each flight, modeled

enhancements are first scaled by a constant to have the same mean as the mean observed enhancement. This step is performed to remove any existing mean bias that may exist in the prior inventory for the calculation of  $\mathbf{R}$ . After removing this bias, the residual errors are calculated between the model and observations and the standard deviation of this error is used to represent the  $\mathbf{R}$  value along its diagonal for a given flight, with no value assigned for the off-diagonals. This process results in flights with large observational and transport uncertainty on days with large enhancements and poor correlations between the observed and modeled values, thus giving these flights less weight in the inversion solution. Similarly, flights where observed and modeled plume structures align have a smaller  $\mathbf{R}$  value assigned and thus are given greater weight in the overall solution. Values for the diagonal elements of  $\mathbf{R}$  varied across flights (Figure S3), but seasonal averages for the standard deviation of the error ranged from 0.7 ppb in the spring to 1.8 ppb in the fall. This method for classifying transport uncertainty is particularly effective for a  $\text{C}_2\text{H}_6$  inversion study since the general locations of the sources (i.e., O&G infrastructure) are known with high confidence such that misaligned plumes would most likely be caused by errors in the transport rather than problems with the spatial mapping of the flux.

For the flux uncertainty matrix  $\mathbf{B}$ , there lacks a clear answer on what the uncertainty of the prior fluxes should be. Since the primary source of  $\text{C}_2\text{H}_6$  in the US is from the O&G production sector, the location of  $\text{C}_2\text{H}_6$  emitters in the US should be accurate. However, leak rates from O&G activity are beyond the uncertainty bounds of the EPA 2012 Gridded  $\text{CH}_4$  Inventory (Alvarez et al., 2018; Maasakkers et al., 2016). For this reason, a value for the flux uncertainty  $\mathbf{B}$  is selected based on mathematical constraints of the inversion rather than an understanding of the prior flux map. Specifically, all flux grids are assigned an error as a percentage of their prior, where the percent uncertainty is selected using a chi-square metric, defined as

$$\chi^2 = \frac{1}{m}(\mathbf{y} - \mathbf{H}\mathbf{x}_0)^T(\mathbf{H}\mathbf{B}\mathbf{H}^T + \mathbf{R})^{-1}(\mathbf{y} - \mathbf{H}\mathbf{x}_0) \quad (3)$$

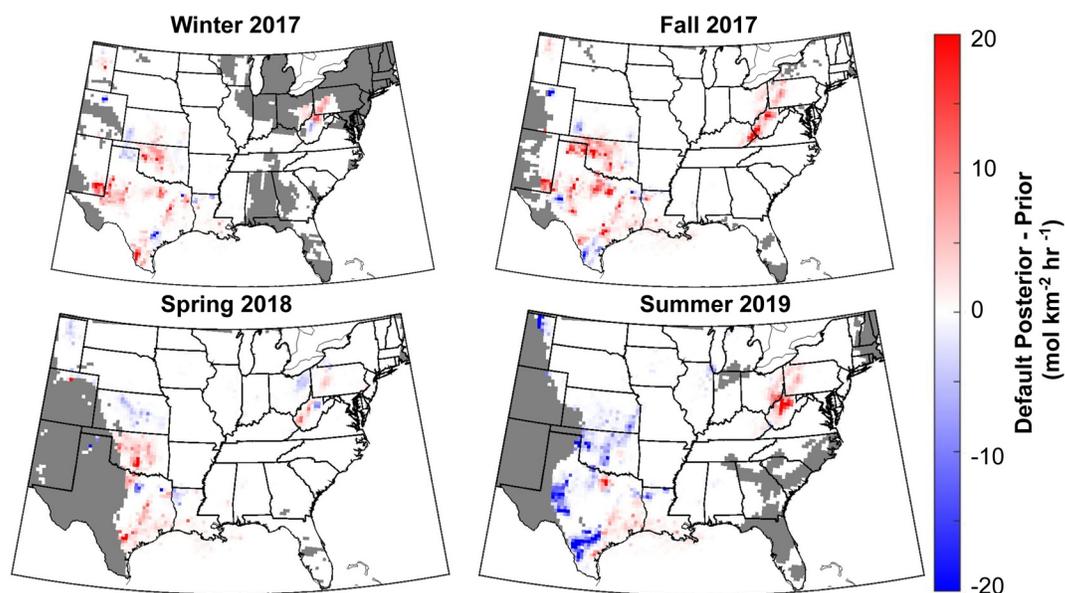
where  $m$  is the number of observations and  $\chi^2$  is the chi-square metric used to assess whether the inversion errors satisfy a Gaussian distribution. Here, we select a percent error for the flux uncertainty in  $\mathbf{B}$  that brings  $\chi^2$  close to 1. For the Default, Flat Ratio, and NEI2017 inventory, the assigned flux errors along the diagonal of  $\mathbf{B}$  were 50, 50, and 80% of the total grid emissions, respectively. Additionally, a correlation length with an e-folding decay length of 50 km is assigned to the off-diagonal elements in  $\mathbf{B}$ . This added correlation provides consistent shifts in the posterior emissions within a basin while allowing changes across basins to behave uniquely. Implications of the  $\mathbf{R}$  and  $\mathbf{B}$  matrices selected for this study are explored further in a sensitivity analysis. For sources unrelated to the O&G sector, no flux uncertainty was applied, such that all changes in the posterior are reflective of changes to O&G emissions.

To test the robustness of the inversion results, a sensitivity analysis is performed for each prior inventory, adjusting various parameters of the inversion set up to examine how these changes impact the posterior solution. Among the tests performed were sensitivity to errors in the model boundary layer depth and wind speed, sensitivity to a low and high biased prior, sensitivity to the background choice, sensitivity to the correlation length scale used in the inversion, sensitivity to the chosen transport uncertainty matrix  $\mathbf{R}$ , and sensitivity to potential errors in the non-O&G sources. Details on how each of these tests was performed can be found in Section S3.

### 3. Results and Discussion

#### 3.1. $\text{C}_2\text{H}_6$ Inversion Results

For the seasonal inversion results, we focus specifically on changes to grid cells that had at least two flights with overlapping influence functions in a season (area shown in Figure 3). For each season, posterior  $\text{C}_2\text{H}_6$  flux maps were successful in reducing both the mean absolute error and bias and increasing the correlation between the observed and modeled signal for that season (Table 1). In the fall, winter, and spring campaigns there was an overall positive bias between the modeled prior and observed  $\text{C}_2\text{H}_6$  enhancements (observed enhancements larger than modeled enhancement), resulting in seasonal posterior maps that generally increased  $\text{C}_2\text{H}_6$  emissions in order to compensate (Figure 3). Regionally, observed  $\text{C}_2\text{H}_6$  plumes were consistently underestimated in the western Appalachian region during the fall, winter, and summer



**Figure 3.** Change in emissions between the Default posterior and prior  $C_2H_6$  flux map for each individual season. Areas with less than two flights that captured their influence in a season (see Figure 2) are greyed out.

season, and posterior  $C_2H_6$  emissions were increased by 50%–110% more than the prior to correct for the underestimation. In the southcentral US, a similar negative bias was observed in the winter, fall, and spring campaigns. This led to posterior solutions for these three seasons that show a consistent increase to the total posterior flux between 30% and 50%. The increase in the southcentral region is largest and most con-

**Table 1**  
Table Describing the Performance of Each of the Seasonal Posteriors Relative to the Observations From That Season

	Default Inventory Winter 2017 Performance		Default Inventory Fall 2017 Performance		Default Inventory Spring 2018 Performance		Default Inventory Summer 2019 Performance	
	Prior	Posterior	Prior	Posterior	Prior	Posterior	Prior	Posterior
Southcentral O&G $C_2H_6$ ( $mol\ s^{-1}$ )	917 (100%)	1246	905 (98%)	1375	413 (45%)	574	681 (74%)	239
Western Appalachia O&G $C_2H_6$ ( $mol\ s^{-1}$ )	88 (70%)	131	124 (99%)	262	125 (100%)	128	125 (100%)	265
Mean Absolute Error (ppb)	1.25	0.98	1.49	1.14	0.64	0.55	1.05	0.69
Mean Bias (ppb)	-0.75	-0.33	-1.12	-0.53	-0.19	-0.15	0.22	0.05
y, Hx Correlation	0.82	0.89	0.74	0.82	0.53	0.65	0.37	0.71

*Note.* Definitions for the regions encompassed in “Southcentral” and “Western Appalachia” can be seen in boxes G and H of Figure 2. Seasonal total emissions calculations for the Southcentral and Appalachia totals only include emissions from grid cells that had at least 2 flights that captured their influence in a season (shown in Figure 3). The percentage of regional emissions captured in each season is given in parenthesis under the prior column for each region. The total prior  $C_2H_6$  emissions contained in the entire Southcentral and Western Appalachia regions are 917 and 125  $mol\ s^{-1}$ , respectively.

**Table 2**  
Table Describing the Performance of Each of the Individual Seasonal Posterior Fluxes and the Three Season Combined Posterior When Applied to all Four Seasons of Observations

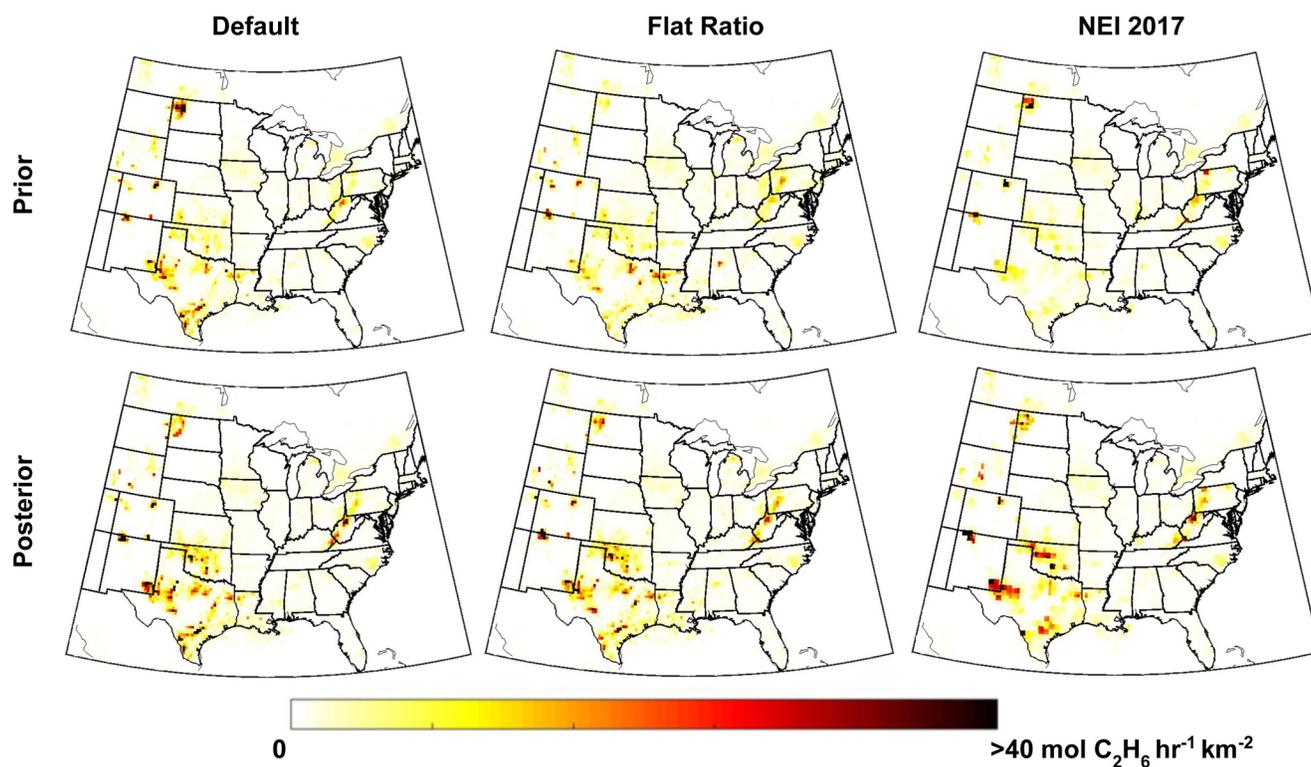
	Default Inventory Prior	Default Inventory Winter 2017 Posterior	Default Inventory Fall 2017 Posterior	Default Inventory Spring 2018 Posterior	Default Inventory Summer 2019 Posterior	Default Inventory 3 Season Posterior
Four Season Mean Absolute Error (ppb)	1.13	1.07	1.05	1.09	1.2	1.01
Four Season Mean Bias (ppb)	-0.44	-0.17	0.07	-0.36	-0.75	-0.14
Four Season $\rho_{y,Hx}$ Correlation	0.67	0.71	0.76	0.67	0.46	0.76
Percent of Flights with Reduced Mean Absolute Error	N/A	65	59	64	45	77

Note. Green areas highlight statistical improvement compared to the prior, whereas red boxes show degradation.

sistent in the western Anadarko region and northwestern Permian basin for seasons with coverage in those regions. A more detailed breakdown of the seasonal posterior performance and emissions can be found in Figures S4–S11.

Observations from the southcentral and midwestern US during the Summer 2019 campaign behave like outliers compared to the rest of the data set. Of the 87 flights used in the inversion analysis, only 12 had an overall negative bias when comparing the observed  $C_2H_6$  to the model prior (observed enhancement less than modelled enhancement), eight of which occurred in the southcentral and midwestern portion of the Summer 2019 campaign. These low observations have a profound effect on the posterior emissions in the southcentral US for the Summer 2019 campaign, with total emissions decreasing by 65% from the prior compared to the 30% or greater increase in all other seasonal campaigns. To demonstrate how unrepresentative the Summer 2019 results are compared to the rest of the data set, we take the posterior inventory derived for each season and apply it to all entire four seasons of observations (Table 2). In doing this, we find that the winter, fall, and spring posteriors all produce similar statistical improvements to the overall data set compared to the prior. However, when applying the summer  $C_2H_6$  posterior to the four season data set, not only does it perform substantially worse than each of the other seasonal posteriors, it also performs worse than the prior, with increases to the absolute error, bias, and a substantial decrease in the model-obs correlation coefficient from 0.67 to 0.46.

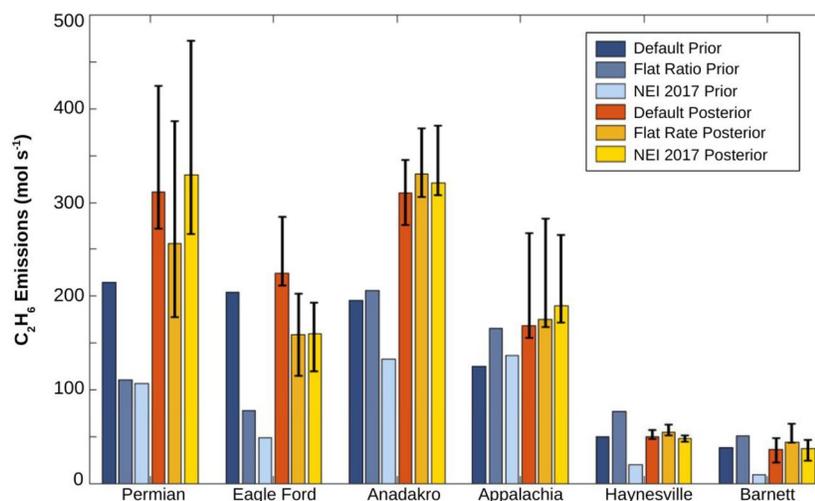
Differences between observed and modeled  $C_2H_6$  enhancements can occur for various reasons, including both errors in the prior emissions as well as errors in the imperfect model transport used to create the influence functions. One possible explanation for the discrepancy in summer could be related to errors in the model transport simulation. Daily errors in the modeled boundary layer wind direction compared to flight observations were largest during the Summer 2019 campaign, likely related to the slower wind speeds observed during summer (Tables S3–S6). These directional transport errors can create misalignments in observed versus modeled plumes, leading to an overall reduction in the posterior emissions relative to the truth. Compounding on these wind issues is the possibility of increased convective activity in summer. While boundary layer statistics are tracked in the model and compared to observations, it is more difficult to assess how much of the signal is lost out the top of the boundary layer due to summertime convective processes. This possibility is supported by the fact that the gradient between the free troposphere and boundary layer  $C_2H_6$  mixing ratios was smallest during the Summer 2019 campaign (Figure S2), and dates with the largest positive model bias occurred on days with southerlies bringing hot and humid air from the Gulf of Mexico into the flight path. Another, nontransport-related possibility is that there could be a non-trivial loss of  $C_2H_6$  due to an increase in the OH chemical sink during the summer months. In this study, the impacts of  $C_2H_6$  loss were not considered, as  $C_2H_6$  has an average lifetime of two months and most plumes were captured within 48 h of release from the source (Burkholder et al., 2015; Thompson et al., 2003). However,



**Figure 4.** Prior  $C_2H_6$  inventories used in this study and their respective three season mean posteriors. “Default” represents the best guess prior from this study based on multiplying the O&G sector of the Gridded EPA 2012  $CH_4$  Emissions Inventory by best-available  $C_2H_6/CH_4$  ratios of emissions in individual basins. “Flat Ratio” multiplies the same  $CH_4$  inventory by a uniform rate  $C_2H_6/CH_4$  ratio of 0.085, producing a similar total as “Default” with a different spatial representation. “NEI 2017” comes from the NEI 2017  $C_2H_6$  inventory.

in conditions with excessive heat and large OH mixing ratios that can occur in the summer, the lifetime of  $C_2H_6$  can be reduced to as little as 4 days in the most extreme conditions (see Section S2 for more details). Even so, a lifetime of 4 days would only have minor impacts on local plumes, and the extreme conditions leading to high  $C_2H_6$  loss would only exist for a short period in the afternoon hours. Furthermore, summer  $C_2H_6$  observations from the northeast were elevated to levels similar to those observed during the other seasons, seemingly unaffected by summertime conditions. For these reasons, it is unlikely that a chemical loss could be solely responsible for the low values observed in the southcentral US. One final possibility is that the low  $C_2H_6$  enhancements observed in the Summer 2019 campaign are due to a real and significant temporal change in  $C_2H_6$  emissions in the southcentral US during this period. However, the near-zero emission rate solution provided by the inversion posterior in the various southcentral plays lacks any sort of real-world explanation for such a large shift compared to previous seasons. Production data shows the Texas and Oklahoma were at peak O&G production rates during the period and well counts were stable (US Energy Information Administration, 2020a). Regardless of the reason, the low  $C_2H_6$  observations measured in the southcentral US during the Summer 2019 months are not representative of data from the winter, fall, and spring campaigns, which show better consistency in the location and expected magnitudes of regional  $C_2H_6$  enhancements. For this work, we choose to discard summer data in evaluating our best estimate of US  $C_2H_6$  emissions but consider it a source of uncertainty and intrigue for future research.

To create our best guess regarding  $C_2H_6$  emissions across the southcentral and eastern US, we combine the observations from the winter, fall, and spring campaigns and rerun the inversion using the Default prior to coming up with a posterior solution that best describes data from all three seasons. In combining the seasonal data sets, influence functions for available flights provide coverage over the entire southcentral US and the western Appalachian region, allowing the inversion to solve for posterior fluxes for all grid cells in each region. This combined posterior solution provides consistent improvement compared to the prior, reducing the absolute error between the model and observations in 86% of flights from those three seasons.

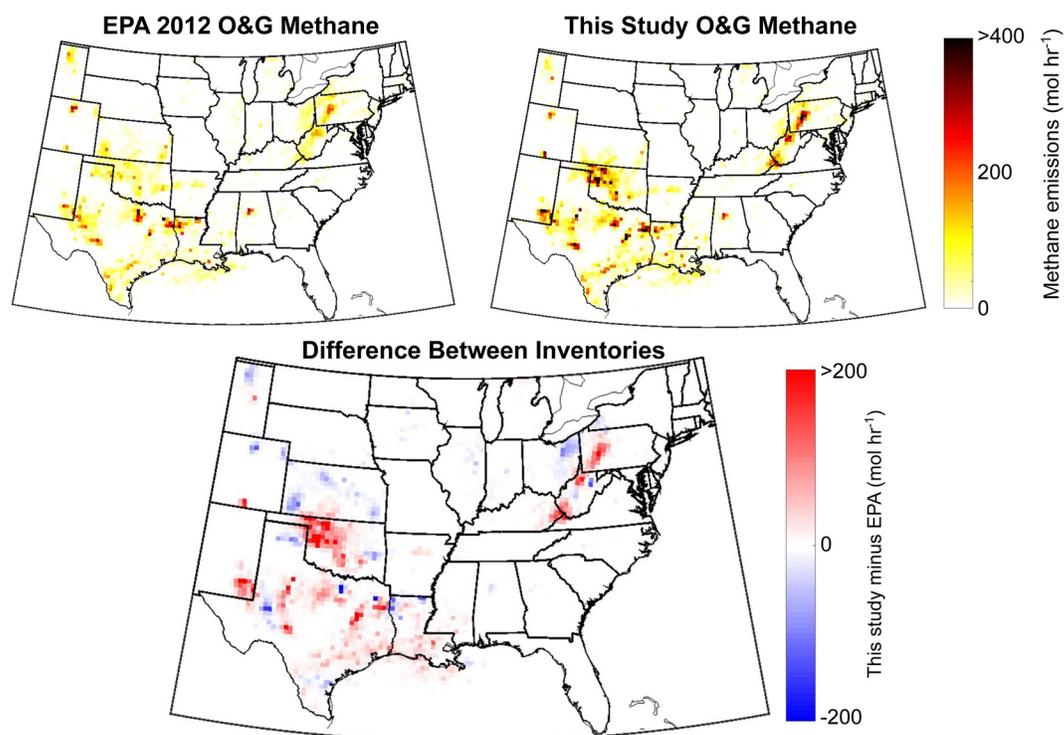


**Figure 5.** Priors and their respective 3 season posterior C<sub>2</sub>H<sub>6</sub> emissions for major basins and plays in this study. Error bars on the posteriors show the maximum and minimum values derived for each region from the sensitivity analysis (Tables S7–S12).

C<sub>2</sub>H<sub>6</sub> emissions from the 3 season posterior are almost universally increased compared to the prior, with a 42% increase overall in the southcentral US and a 34% increase in the Appalachia. With these large increases in the emissions, the mean model-obs bias is improved from  $-0.80$  to  $-0.41$  ppb. The inability to eliminate this bias is a natural result of errors in the transport preventing the inversion from solving for misaligned observed plumes, and thus the emission increases in the 3 season posterior are likely still an underestimation of the true emissions.

To better understand the sensitivity of our three season posterior to the prior, the inversion was rerun using the two alternative prior maps discussed in Section 2.2 and a three season posterior was created from each set of priors (Figure 4). For every basin and play, posteriors from the three inventories converged towards a similar solution compared to their priors, despite starting from very different initial states. An example of this is the Haynesville basin, where the two alternative prior inventories have a factor of 4 spread between their emissions, but their posteriors converge to within 15% of the Default posterior solution. Ultimately, the two alternative posteriors produced solutions for the entire southcentral US that were within 45 mol/s of each other, despite starting 375 mol/s apart. In the northeast, discrepancies between the three inventory priors were much smaller, and all three posteriors converged to similar values (168–190 mol/s). Overall the NEI2017 prior, whose total C<sub>2</sub>H<sub>6</sub> emissions are 88% less than the Default prior, leads to a total posterior C<sub>2</sub>H<sub>6</sub> emissions estimate that is within 10% of the Default posterior.

Though the posteriors generated from the Default and NEI2017 prior inventories converge to a similar emissions map, the interpretation of the values they converge to are vastly different. The NEI2017 prior is an official product provided by the EPA and specifically provides C<sub>2</sub>H<sub>6</sub> emissions for the year 2017, whereas the Default inventory from this study was created using CH<sub>4</sub> emissions data from O&G sources for the year 2012 combined with knowledge of C<sub>2</sub>H<sub>6</sub>/CH<sub>4</sub> gas composition data of different basins. Despite both priors attempting to represent the same C<sub>2</sub>H<sub>6</sub> emissions, their representations of the southcentral US in particular are starkly different. The posterior solution of 1,313 mol/s from the Default posterior is a 43% increase compared to its prior of 917 mol/s, but the posterior solution of 1,176 mol/s from the NEI2017 posterior is a 182% increase compared to its prior of 417 mol/s. The convergence of solutions provides confidence in the ability of the ACT data set to constrain C<sub>2</sub>H<sub>6</sub> emissions but reveals a large discrepancy between the NEI2017 C<sub>2</sub>H<sub>6</sub> product and atmospheric data which points towards significantly larger emissions, particularly in the southcentral US. Given that the Default inventory prior from this study is based on EPA (2012) O&G CH<sub>4</sub> emissions data but contains more than double the O&G C<sub>2</sub>H<sub>6</sub> emissions compared the NEI2017 prior, the low bias in the NEI2017 C<sub>2</sub>H<sub>6</sub> inventory may be partially explained by some methodological flaw in its creation that underestimates C<sub>2</sub>H<sub>6</sub> emissions from documented O&G sources rather than some unknown or missing sources.



**Figure 6.** Top-left: Oil and gas CH<sub>4</sub> emissions from the EPA (2012) Gridded CH<sub>4</sub> Inventory. Top-right: Oil and gas CH<sub>4</sub> emissions estimated from the C<sub>2</sub>H<sub>6</sub> posterior in this study. Bottom: Difference between the two inventories.

In addition to testing the sensitivity of the inversion posterior to different priors, we also test its sensitivity to numerous other conditions, including adjustments to model meteorology errors, adjustments to the magnitude of the prior, changes to the selection of the background C<sub>2</sub>H<sub>6</sub> term, elimination of the correlation lengths applied to the prior flux uncertainty, adjustment of the observation error covariance matrix term to a constant (giving equal uncertainty to all observations), and allowing for the adjustment of all C<sub>2</sub>H<sub>6</sub> sources rather than only O&G sources. Details and tables with values from this sensitivity test can be found in Section S3 and Tables S7–S12, and the range of solutions is shown in Figure 5. Most posteriors produce similar results regardless of the methodology, providing further confidence in the solution. One particular concern was the inherent low bias in all of the prior inventories compared to the posterior solutions. However, multiplying the Default prior by 3 (making it much larger than the original posterior solution) and rerunning the inversion resulted in a total posterior solution that was only 2% higher than the original posterior, reducing concerns that using a prior with a low bias could be restricting the posterior solution. Of all the sensitivity tests, the one that had the largest impact for all three priors was adjusting the observation error covariance matrix **R**. Based on its original setup, larger uncertainty is given on days with larger C<sub>2</sub>H<sub>6</sub> plumes, as any transport error on these days could potentially create larger mismatches between modeled and observed enhancements. This leads to less weight being given to days with the biggest signals, particularly in the fall and winter seasons. By assigning a constant value to **R** and giving equal weight to all flights, the overall Default posterior emissions increase by 21% compared to the original posterior, or 72% compared to the prior.

One final source of uncertainty that is difficult to address in our sensitivity tests are errors in the transport used to create the influence functions. Observed minus modelled boundary layer wind speed and direction biases were small during the winter, fall, and spring aircraft campaigns. However, modelled boundary layer heights were shallow compared to observations, particularly during the spring campaign where observed boundary layer heights were on average 46% higher than modeled heights. A negative bias in the modeled boundary layer would result in a proportionally higher value in the influence functions and model-projected enhancements. As a way of testing these impacts, we rerun the inversion applying a unique correction

factor to the influence functions on each day to account for biases in the modeled versus observed wind speed and boundary layer heights from each flight day. Doing this correction produces an overall Default posterior that is 16% higher than the original Default posterior. Consistent with our past work that has applied corrections to simulated errors in ABL depth and winds (Barkley, Davis, et al., 2019; Barkley, Lauvaux, et al., 2019), we consider the posterior created using the ABL depth correction to be an equally plausible solution in our best estimate of the  $C_2H_6$  emissions, as the boundary layer bias is a source of error with a known and somewhat correctable bias on the solution. The boundary layer bias, along with a negative enhancement bias still present in the model versus observational comparison of the posterior  $C_2H_6$  enhancements, is both potential reasons to suspect the Default posterior on its own may still be underestimating overall  $C_2H_6$  emissions despite the large increase in the emissions relative to the prior.

### 3.2. Interpretation of $CH_4$ Emissions From the O&G Sector

The Default prior inventory developed in this study is created by multiplying the EPA (2012) Gridded  $CH_4$  Emissions Inventory for O&G sources by the assumed mean  $C_2H_6/CH_4$  ratio of each grid's emissions. If the applied ratios are correct, and emissions primarily occur in situations where the gas content is unaltered (i.e., processes unrelated to gas separation), then changes between the posterior and prior Default  $C_2H_6$  inventory should proportionally reflect changes in the EPA (2012)  $CH_4$  inventory for O&G. Thus, we can use our Default posterior  $C_2H_6$  inventory to create our best interpretation of O&G  $CH_4$  emissions in the southcentral and eastern US.

Using the converted  $C_2H_6$  posterior as a proxy for O&G  $CH_4$  emissions, our inventory projects that O&G  $CH_4$  emissions are almost universally larger than the 2012 EPA Inventory estimates (Figure 6). In the southcentral US, we estimate emissions to be 48% higher than inventory estimates (77% using the meteo-corrected posterior). Of this increase, more than half of it is driven by increases in the Anadarko and Permian plays, both of which had the largest proportional changes in the  $C_2H_6$  posterior and are some of the largest sources of  $CH_4$  emissions in the southcentral US. Of the remaining southcentral basins captured in this study, the Haynesville and Barnett basins are the only ones that did not see an increase in their emissions relative to the prior.

In the Appalachians, we estimate  $CH_4$  emissions to be 32% higher than EPA (2012) inventory estimates (84% using the meteo-adjusted posterior). Part of the discrepancy between inventory results and the posterior may be related to the increased presence of unconventional natural gas activity in the Marcellus shale. Between 2012 and 2018 Pennsylvania and West Virginia underwent some of the largest gas production growth in the US, with annual production tripling during the period (US Energy Information Administration, 2020a), a change that would not be captured in the EPA (2012) inventory. However, Pennsylvania state inventories, which provide annual inventory estimates of unconventional natural gas activity in the state using methodologies similar to the EPA, show  $CH_4$  emissions from unconventional activity only increased by 20% during this period, and that these unconventional wells only represent a small fraction of O&G emissions in the region (Omara et al., 2016), with much of the emissions coming from pre-existing conventional activity. Thus, it is unlikely that changes in unconventional activity between 2012 and the time of this study would be responsible for the increase in regional emission rates found from the inversion results relative to the EPA (2012) Gridded Inventory, and that the discrepancy would still be present in an updated EPA inventory.

The three-season  $CH_4$  inventory estimates for individual basins from this study generally align with top-down estimates of corresponding basins from the NOAA SONGNEX campaign (Peischl et al., 2018) and other studies. In the Haynesville basin, we calculate an O&G emission rate in our Default posterior ensemble of 50–59 Mg/hr, compared to 42 Mg/hr from Peischl et al. (2018) (year of study: 2015) and 76 Mg/hr from Cui et al. (2019) (year of study: 2013), which includes non-O&G values in its total as well. In the Barnett, we calculate emissions to be 29–42 Mg/hr, close to the 46 Mg/hr found in Peischl et al. (2018) (year of study: 2015) but less than the 60 Mg/hr found in Karion et al. (2015) (year of study: 2013). In Eagle Ford, both the western and eastern basin in this study had a combined emission rate of 54–72 Mg/hr versus 83 Mg/hr in Peischl et al. (2018) (year of study: 2015). In the Appalachian, our results support findings that show emissions from unconventional O&G infrastructure in the Marcellus are greatly underestimated by

EPA inventory estimates (Barkley, Lauvaux et al., 2019; Caulton et al., 2019). One area where we see strong disagreement is in the Permian, where this study's emissions estimate of 117–182 Mg/hr is substantially less than top-down estimates from Zhang et al. (2020) of 308 Mg/hr (years of study: 2018–2019) and is much closer to their updated bottom-up estimate of 114 Mg/hr. Overall, this study finds emissions from the southcentral and eastern US to be 48%–76% greater than the EPA 2012 Gridded Methane Inventory for O&G sources, agreeing with national estimates from Alvarez et al. (2018). However, in all of these comparisons, we note that the year each study took place varies, such that differences in the solution may include changes due to temporal variability in the O&G activity of a basin. The comparison of emissions in the Permian is an example of where one could expect large discrepancies. Analysis of the Permian play in this study is mostly representative of the year 2017 due to a lack of coverage from influence functions during the Spring 2018 campaign, whereas in Zhang et al. (2020) the analysis is from 2018 to 2019. Between 2017 and 2019 the Permian was undergoing rapid development, and both O&G production doubled in the 2 year period (US Energy Information Administration, 2021). This change may be partially responsible for the discrepancy between emission estimates in the Permian from this work and Zhang et al. (2020).

The interpretation of O&G CH<sub>4</sub> emissions using C<sub>2</sub>H<sub>6</sub> observations has a unique advantage compared to more traditional methodologies that rely on CH<sub>4</sub> measurements due to the simplicity of C<sub>2</sub>H<sub>6</sub> sources. In the US where CH<sub>4</sub> emissions have near equal contributions from fossil fuels, agriculture, and wetlands, each of which has their own uncertainties, C<sub>2</sub>H<sub>6</sub> emissions are dominated by the O&G sector. Furthermore, there is high confidence in the spatial mapping of O&G sources in the US due to extensive documentation of the various components associated with O&G extraction, simplifying the interpretation of atmospheric C<sub>2</sub>H<sub>6</sub> data. As an example, in the ACT-America campaign, the majority of aircraft flight tracks were hundreds to thousands of kilometers away from O&G basins. Despite this, the model prior was able to consistently track C<sub>2</sub>H<sub>6</sub> enhancements from these sources in the winter, fall, and spring, with correlations between the model versus observed boundary layer C<sub>2</sub>H<sub>6</sub> enhancements of 0.82, 0.74, and 0.53 respectively. The high skill in tracking enhancements from a single sector with well-defined locations creates a scenario where a stable posterior solution can be generated through various inverse methodologies (Table S4).

Despite high confidence in the C<sub>2</sub>H<sub>6</sub> posterior, the conversion of this posterior to O&G CH<sub>4</sub> emissions is entirely dependent on the quality and availability of information related to the C<sub>2</sub>H<sub>6</sub>/CH<sub>4</sub> ratio for each basin. During the time observations from this study were collected, there were numerous recent flights available from a separate study that quantified the mean C<sub>2</sub>H<sub>6</sub> and CH<sub>4</sub> emissions of various basins (Peischl et al., 2018), providing more confidence in a mean C<sub>2</sub>H<sub>6</sub>/CH<sub>4</sub> ratio to apply to those locations. However, geographically broad plays, such as the Anadarko and the Appalachia, are more difficult to characterize a mean ratio for using atmospheric data. Furthermore, the average gas composition of a basin can change over time, making ratios found in older studies less applicable (Lan et al., 2019). Additionally, sources such as C<sub>2</sub>H<sub>6</sub> cracker plants that do not adhere to any of the assigned C<sub>2</sub>H<sub>6</sub>/CH<sub>4</sub> ratios could create additional uncertainties in this conversion, but we assume emissions from these facilities are small relative to emissions from O&G production activity and mostly uncorrelated with the locations of specific basins. Publicizing upstream and midstream gas composition data collected by major O&G companies would be one immediate solution to reduce uncertainty introduced in a C<sub>2</sub>H<sub>6</sub> to CH<sub>4</sub> inventory conversion.

One other concern arises from the age of the inventory used to perform this C<sub>2</sub>H<sub>6</sub> to CH<sub>4</sub> inventory conversion. Comparisons in this study are made relative to the EPA bottom-up estimate from 2012, the most recent year in which gridded emissions data is available, whereas flights for this study took place between 2017 and 2019. A theoretically updated EPA-gridded inventory may show differences compared to the existing 2012 gridded inventory, particularly in basins that have undergone substantial changes to their O&G activity. However, total O&G emissions from the EPA bottom-up inventory from 2012 to 2018 have been stable, with only a small decrease from 6.6 Tg CH<sub>4</sub> in 2012 to 6.4 Tg CH<sub>4</sub> in 2018 (US Environmental Protection Agency, 2020b). Thus, while it can be argued that increases to CH<sub>4</sub> emissions in certain basins observed in this study compared to the 2012 Gridded Methane Inventory may be related to changes in activity since 2012, the 48%–76% increase in overall CH<sub>4</sub> emissions across the southcentral and eastern US relative to the 2012 inventory is not compatible with current EPA bottom-up inventory estimates and indicates that current inventory estimates are undercounting CH<sub>4</sub> emissions from O&G activity.

#### 4. Conclusions

Using the largest collection of airborne boundary layer C<sub>2</sub>H<sub>6</sub> data to date, an inversion was performed to estimate C<sub>2</sub>H<sub>6</sub> and CH<sub>4</sub> emissions from various O&G basins across the southcentral and eastern US. From this data set, we estimate that a large portion of C<sub>2</sub>H<sub>6</sub> emissions cannot be explained using O&G emission data from the EPA (2012) Gridded CH<sub>4</sub> Inventory and existing C<sub>2</sub>H<sub>6</sub>/CH<sub>4</sub> ratio data. We conclude that the EPA CH<sub>4</sub> emissions inventory underestimates CH<sub>4</sub> emissions overall from O&G sources in the southcentral and Appalachian regions by 48%–76%, an offset similar in magnitude to national estimates calculated in Alvarez et al. (2018). In particular, this study finds the largest discrepancies occurring in the Anadarko, western Appalachian, and Permian O&G plays. Though some of these discrepancies may be related to regional changes in O&G activity since 2012, bottom-up O&G inventory emission estimates nationally have decreased by 4% from 2012 to 2018 and cannot explain the higher overall emissions observed in this study. We also find C<sub>2</sub>H<sub>6</sub> emissions from O&G are more than two times higher than the NEI2017 C<sub>2</sub>H<sub>6</sub> inventory, a fact that should be taken into account in any future efforts to investigate O&G CH<sub>4</sub> emissions using atmospheric C<sub>2</sub>H<sub>6</sub> observations, as well as studies relying on existing C<sub>2</sub>H<sub>6</sub> inventories to account for certain chemical reactions.

This study reveals the potential to use broad-scale continuous C<sub>2</sub>H<sub>6</sub> data to constrain CH<sub>4</sub> emissions from the O&G sector on a nationwide scale. Our confidence in the spatial distribution of emissions associated with the O&G sector and the dominant role of O&G in US C<sub>2</sub>H<sub>6</sub> emissions allows for modeling and interpretation of observed signals without large concerns for source attribution. Furthermore, misalignment in observed versus modeled plumes may serve as a useful diagnostic of model transport errors due to the well-documented spatial knowledge of C<sub>2</sub>H<sub>6</sub> emitters in the US. One current weakness with relying on C<sub>2</sub>H<sub>6</sub> observations to understand CH<sub>4</sub> emissions from O&G is the requirement of knowing the C<sub>2</sub>H<sub>6</sub>/CH<sub>4</sub> ratios of various O&G fields in order to convert C<sub>2</sub>H<sub>6</sub> emissions into CH<sub>4</sub> emissions. Gas composition data is collected by individual companies but not shared publicly. Should some form of this information be made available for the public, it would greatly enhance the confidence of CH<sub>4</sub> emission estimates of the O&G industry using C<sub>2</sub>H<sub>6</sub> and other trace gases.

#### Conflict of Interest

The authors declare no conflicts of interest relevant to this study.

#### Data Availability Statement

Aircraft data from the ACT-America campaign used in this study can be found at <https://doi.org/10.3334/ORNLDAAC/1556>.

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