# A Field of Small Pitted Cones on the Floor of Coprates Chasma: Volcanism inside Valles Marineris 

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#### Abstract

We present observations of a field of $>100$ pitted cones and mounds situated on the floor of Coprates Chasma (part of Valles Marineris (VM); Fig. 1), which display similarities to terrestrial and martian scoria cones. If these cones are indeed volcanic in origin, they will significantly expand our knowledge about the morphometry of pyroclastic cones on Mars. Moreover, a magmatic origin, which would necessarily post-date the opening of the main VM troughs, would contribute to our understanding of the volcano-tectonic evolution of VM.




Figure 1: Location of investigated cones (triangles) and mounds (circles) in eastern Coprates Chasma (CTX mosaic) The edifices are spread over the entire trough.

## 1. Introduction

It has long been suggested that volcanism played a role in the formation of Valles Marineris [e.g., 1], but unambiguous evidence has been rare. Recent images acquired by the CTX camera ( $\sim 5-6 \mathrm{~m} / \mathrm{px}$ ) reveal the existence of several fields of small pitted cones, mainly associated with chaotic terrain in the eastern part of Valles Marineris [2,3] and also in Coprates Chasma [3,4]. Coprates Chasma is a linear graben extending in west-east direction for $\sim 1000 \mathrm{~km}$. It probably formed as one of the most recent main depressions of the VM system. Based on morphological similarities, Harrison [3] suggested that these cones might represent scoria cones, but without providing further details. Meanwhile the whole area is covered by CTX images that enable analysis of the entire cone field. One
cluster of pitted cones is covered by a HiRISE stereo pair, allowing the production of a HiRISE Digital Elevation Model (DEM). Based on these new data, we studied the cones in unprecedented detail. Here we show preliminary results.

## 2. Data and methods

We used images from CTX, HRSC, and HiRISE. Topographic information is derived from single MOLA shots, HRSC DEMs, and HiRISE and CTX DEMs that were computed using the methods described in [5]. HiRISE and CTX DEMs have a grid spacing of $\sim 1 \mathrm{~m} /$ pixel and $\sim 10 \mathrm{~m} /$ pixel and a vertical accuracy of approximately a few decimeters and a few meters, respectively.

## 3. Observations

The cones and mounds are widely spread over a total area of about $155 \times 35 \mathrm{~km}$. Some cones stand alone, others are concentrated in clusters with up to ten edifices. In plan view, their morphology is characterized by circular to elongated shapes. Their flanks have slopes up to $25^{\circ}$, but are generally more shallow. Cone basal diameters vary from 0.5 km up to 2.2 km , with a mean of 1 km (based on 23 cones). Most of them have summit craters (Fig. 2a), which have diameters from 0.15 km up to 0.8 km (mean 0.3 km ). In some cases craters are superposed by other craters suggesting the lateral migration of explosion sites or feeder dikes. Typically, the cones are not breached, but there are two exceptions which seem to result from explosion and/or collapse of the cone. In some cases cones are superposed on units with a rough texture that forms local bulges (Fig. 2a). To compare the cones morphologically with martian and terrestrial analogues, we measured the basal diameters of the cones $\left(\mathrm{W}_{\mathrm{CO}}\right)$ and the crater diameters $\left(\mathrm{W}_{\mathrm{CR}}\right)$. The $\mathrm{W}_{\mathrm{CR}} / \mathrm{W}_{\mathrm{CO}}$ ratio ranges between 0.22 and 0.5 , with an average of 0.34 . Fresh terrestrial and Martian scoria cones have ratios of $\sim 0.4$ [6], and $\sim 0.27$ [7], respectively. The cones appear relatively pristine, and small impact craters do not change their shapes significantly.


Figure 2: Cluster of investigated cones in Coprates Chasma (a) and comparison to similar features on Mars and Earth (b-d). (a) HiRISE ESP 0341311670 , centered $12.73^{\circ} \mathrm{S}, 62.8^{\circ} \mathrm{W}$, (b) Hydraotes Chasma; CTX image G19_025493_1800, $0.2^{\circ} \mathrm{N}, 33.83^{\circ} \mathrm{W}$, (c) Ulysses Colles; CTX image G11_022582_1863, $5.81^{\circ} \mathrm{N}, 122.59^{\circ} \mathrm{W}$, (d) Andes (Earth); NASA, Digital Globe, Google Earth ${ }^{\mathrm{TM}}, 26.29^{\circ} \mathrm{S}, 67.35^{\circ} \mathrm{W}$.

## 4. Summary and Conclusions

The studied cones bear many morphological similarities to edifices in Hydraotes Colles [2] (Fig. 2b) and Ulysses Colles [7] (Fig. 2c) that were previously interpreted as scoria cones and with terrestrial scoria cones (Fig. 2d). The Coprates cones are smaller (WCO on average 1 km ) than the Hydraotes cones ( 1.5 km ) and the Ulysses cones ( 2.3 km ), but with similar WCR/WCO ratios ( 0.34 for cones in Coprates and 0.27 for cones in Ulysses). This might be caused by a higher atmospheric pressure at the floor of Coprates Chasma ( $\sim 5 \mathrm{~km}$ beneath Mars’ global datum) disabling a wider dispersion of ejected particles from the vent [8], by a smaller amount of erupted material or by smaller erosion. The associated elevated rough units around the edifices are similar to what is observed in Hydraotes Colles and Ulysses Colles. Similar rough textures around terrestrial scoria cones are associated with lava flows and possibly pyroclastic deposits.


Figure 3: Details of cone morphology. (a) ESP_033986_1670, $12.73^{\circ} \mathrm{S} / 297.2^{\circ} \mathrm{E}$. (b) ESP_$036109 \_1675,12.4^{\circ} \mathrm{S} / 297.21^{\circ} \mathrm{E}$. (c) ESP_036-254_1665, $13.28^{\circ} \mathrm{S} / 298.52^{\circ}$ E. (d) Parallel, flat-lying layers exposed in inner crater wall. (e) Similar fine layering in scarp of eroded cone. (f) Exhumed (?) surface adjacent to cone, reminiscent of a lava flow surface texture.

At HiRISE resolution, further details provide possible hints to the nature of the cones. Where scarps break the surface of the otherwise smooth-textured cones, series of fine parallel layers are visible in some places (Fig. 3 a and d, b and e). Although other interpretations are possible, this would be consistent with a volcanic origin of some cones, specifically as phreatomagmatic edifices, e.g., tuff cones. In another location, the surface texture near a cone is reminiscent of a lava flow (Fig. 3 c and f). The spatial distribution of cones ( 2 point-azimuth analysis; for details see ref [9]) reveals two main trends of a possible structural control. One is oriented parallel to the main VM trend $\left(\sim \mathrm{N} 110^{\circ}\right)$, while another is $\sim \mathrm{N} 75^{\circ}$. The first trend suggests that magma feeding the cones may have ascended (as dikes?) along weakness zones created by VM formation. Preliminary age determinations based on crater counting suggest that the cones were formed at 200 Ma to 400 Ma .

Our preliminary results, therefore, support previous suggestions $[3,4]$ that this field is probable volcanic in origin and consists primarily of scoria cones. Ongoing investigations will help to assess alternative formation scenarios (e.g., mud volcanism), extend our knowledge on scoria cone formation on Mars and will also help to provide further insight about the evolution of Valles Marineris.

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References: [1] Lucchitta, B.K. (1987) Science, 235, 565-567. [2] Meresse, S. et al. (2008) Icarus, 194, 487-500. [3] Harrison, T.N. (2012) LPSC, XLIII, Abstract \#1057. [4] Harrison, K.P. and Chapman, M.G. (2008) Icarus, 198, 351-364. [5] Moratto, Z.M. et al. (2010) LPSC, XLI, Abstract \#2364. [6] Wood, C.A. (1980) J. Volcanol. Geotherm. Res., 7, 387-413. [7] Brož, P. and Hauber, E. (2012) Icarus, 218, 1, 88-99. [8] Brož, P. et al. (2014) Earth Planet. Sci. Lett., 406, 14-23. [9] Brož, P. and Hauber, E. (2013) JGRPlanets, 118, 1656-1675.

